



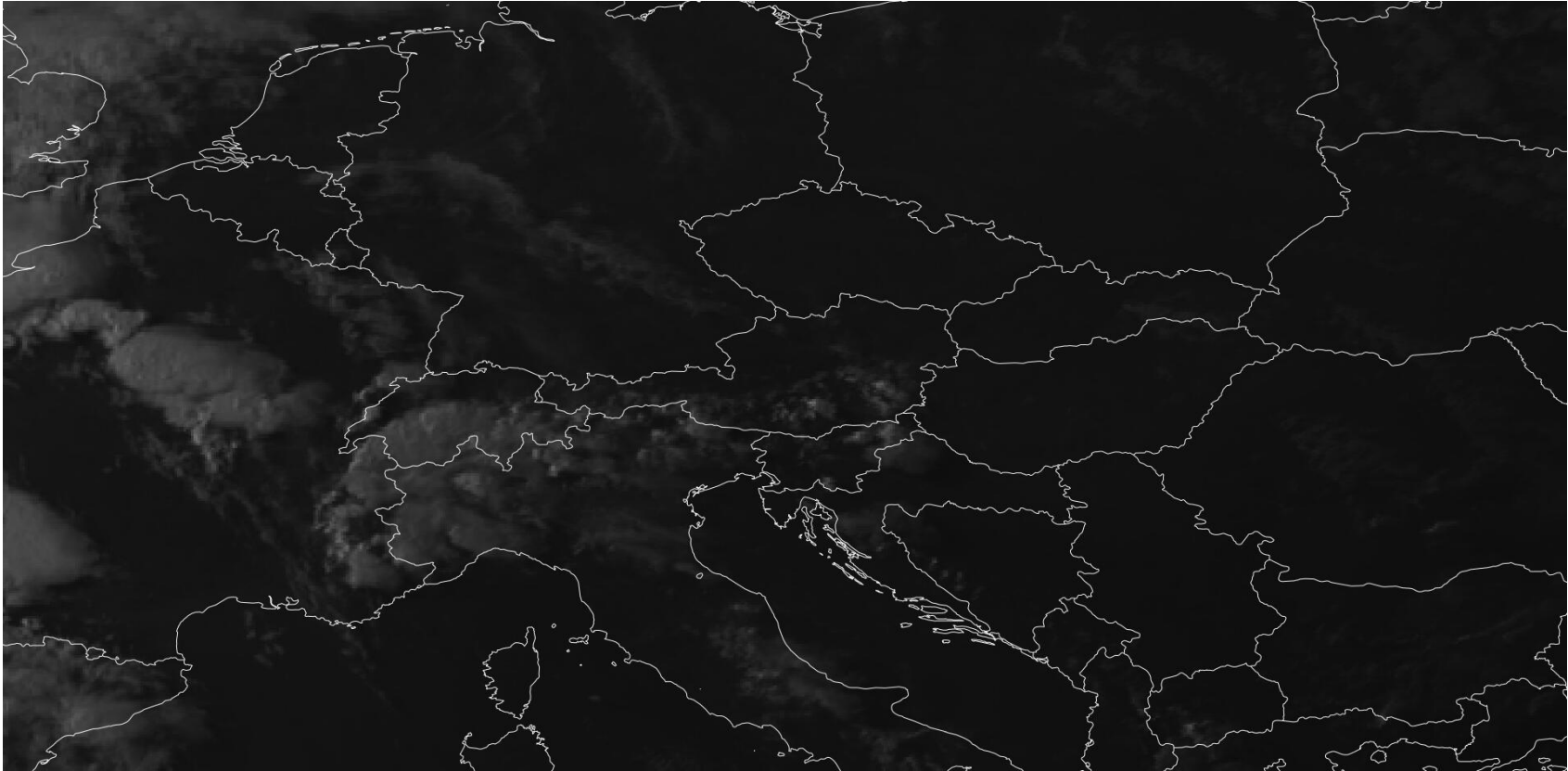
# Are Upper Tropospheric Moisture Gradients relevant for the Development of Deep Moist Convection?

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Zentralanstalt für  
Meteorologie und  
Geodynamik

# Forecasting single cell convection



- Conditions: presence of orography, absence of fronts, desired forecast lead time >3h
- Challenging task: predicting exact location, intensity and propagation of thunderstorms
- Single-cell or pulse convection, without fronts make ~30% of the summertime convection in the Eastern Alpine region, Krennert et al. (2003).

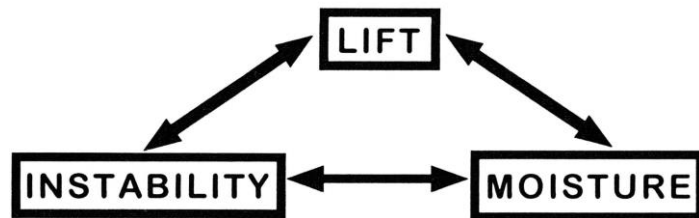


MSG RGB: IR 8.7 $\mu$ m (red), WV 7.3 $\mu$ m (green), WV 6.2 $\mu$ m (blue)

Ch07 (IR 8.7 $\mu$ m) | 07.07.2014 12:00  
Ch06 (WV 7.3 $\mu$ m) | 07.07.2014 12:00  
Ch05 (WV 6.2 $\mu$ m) | 07.07.2014 12:00  
AT - Schwechat: 07.07.2014 12:00  
AT - Salzburg: 07.07.2014 12:00  
AT - Patscherkofel: 07.07.2014 12:00  
AT - Valuga: 07.07.2014 12:00  
AT - Zirbitzkogel: 07.07.2014 12:00

# Ingredients-based methodology to forecast convection

## INGREDIENTS-BASED METHODOLOGY FOR CONVECTION



## Ingredients for deep moist convection DMC

### Instability:

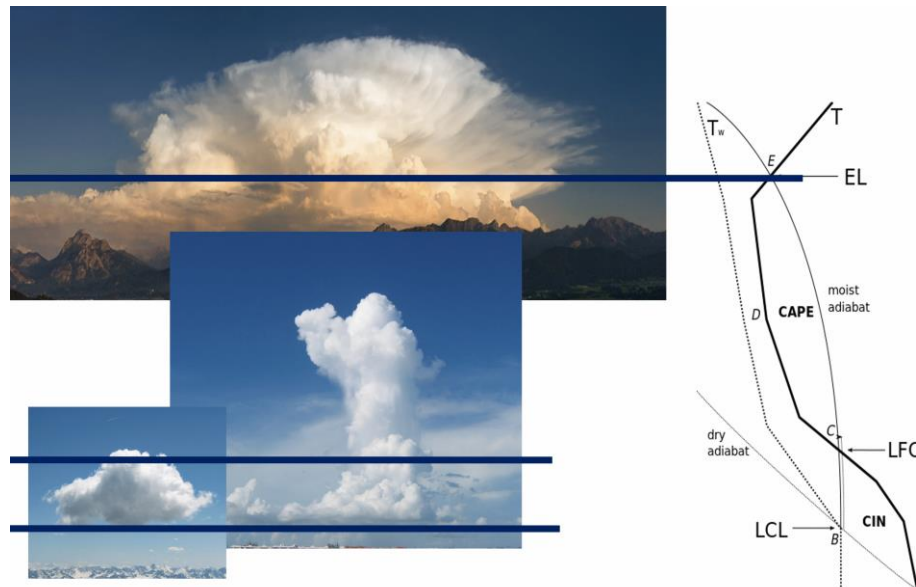
i.e. Conditional Instability, vertical lapse rate, needed to develop CAPE

### Lift:

To bring air parcels to their level of free convection LFC

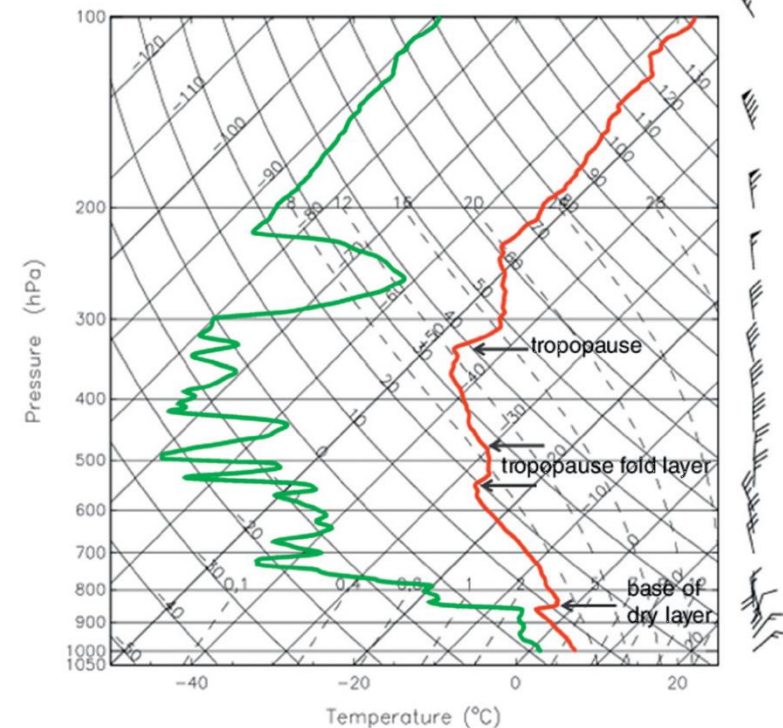
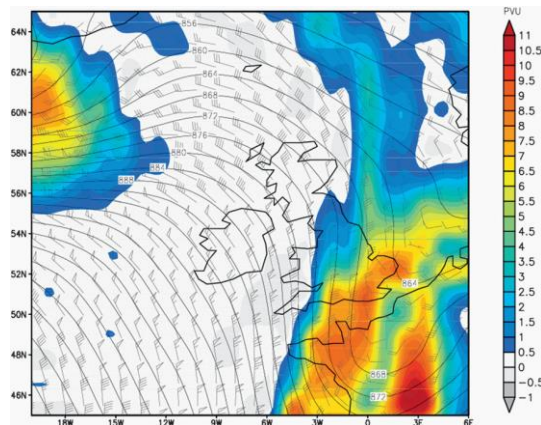
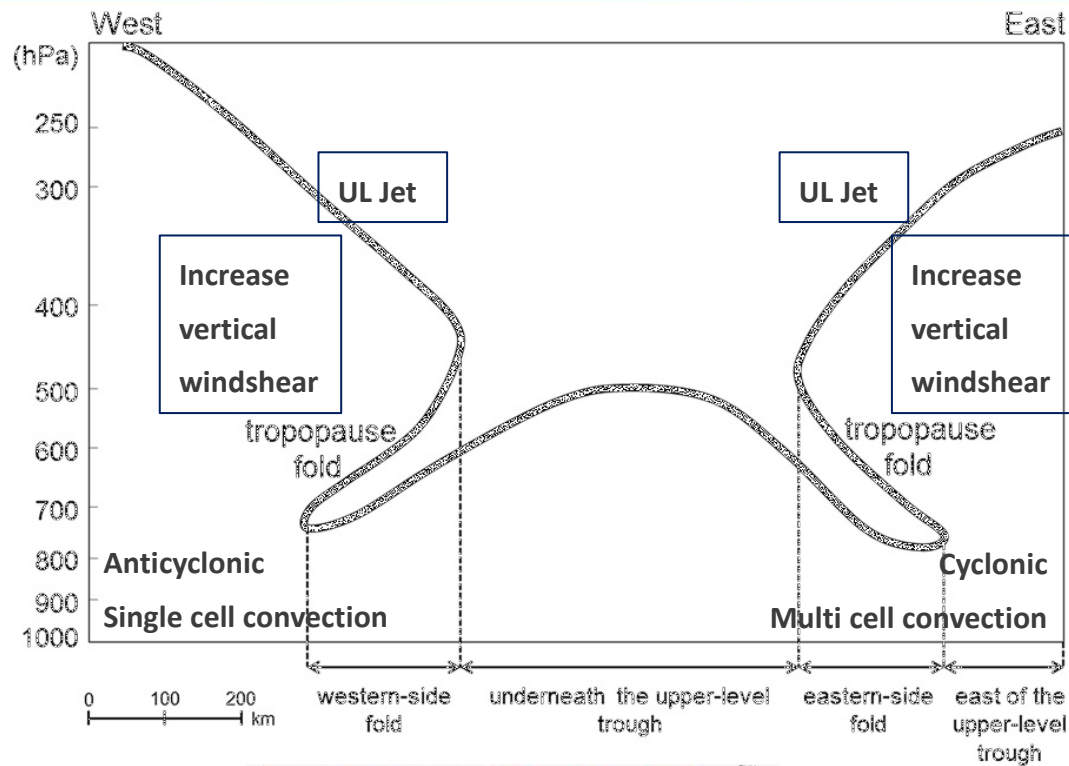
### Moisture:

Latent heat source, to drive convection





# Tropopause Folds and Deep Convection, Antonescu, et al., 2013



# Atmospheric lids beneath an UL PV anomaly, Russel, et al., 2008

Atmospheric lids beneath an upper-level PV anomaly and their impact on the initiation of DMC

Definition of a lid / capping inversion:

- Below the lid: moist layer and high  $\Theta_w$
- Above lid: a break in humidity with a thermal inversion, low  $\Theta_w$  in the middle and upper troposphere

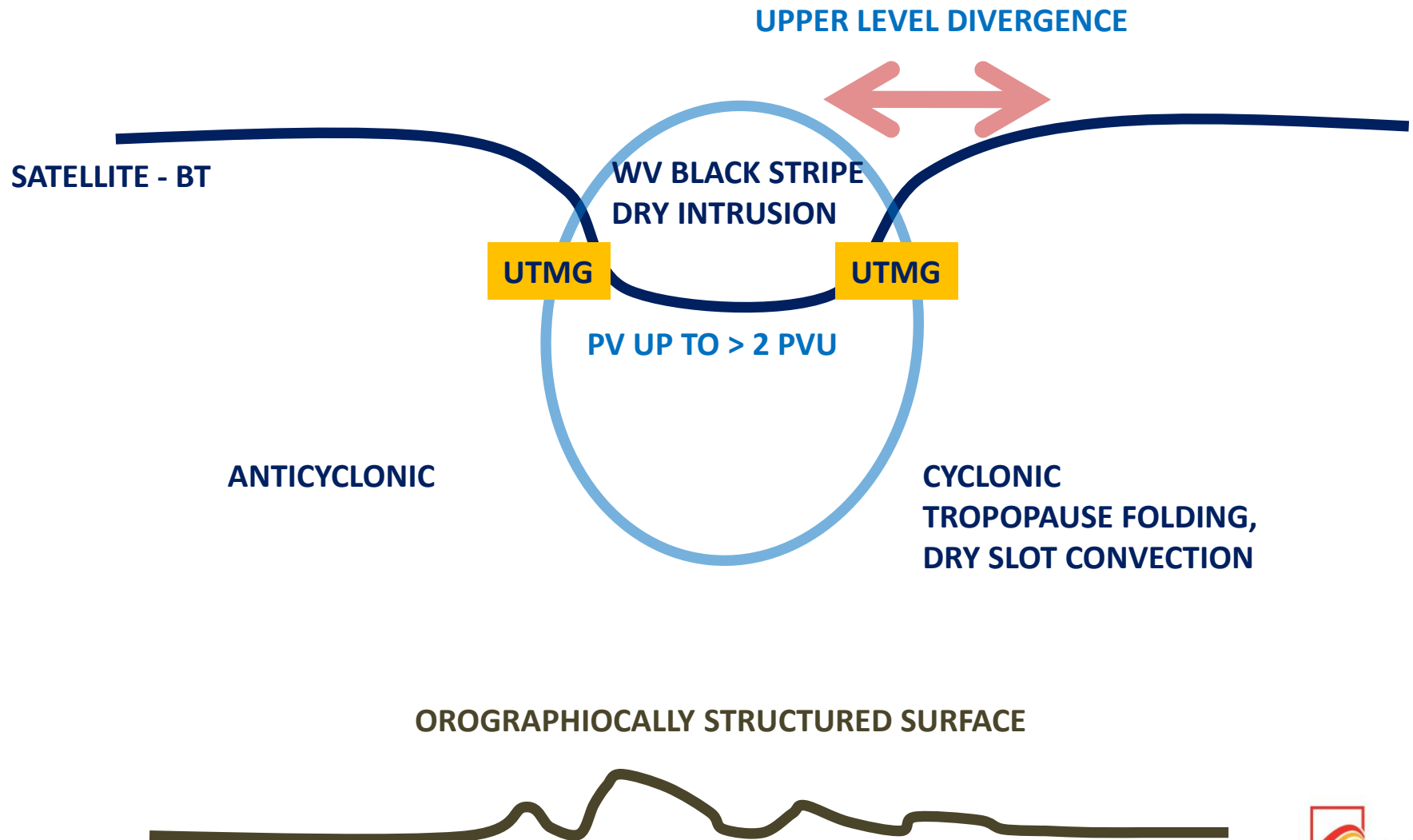
A lid can thus

- suppress convection
- promote convection  
in case of accumulated heat and moisture beneath, therefore creating conditional instability (loaded gun), convection tends to be more severe in the presence of a lid, although it is more likely to occur in its absence

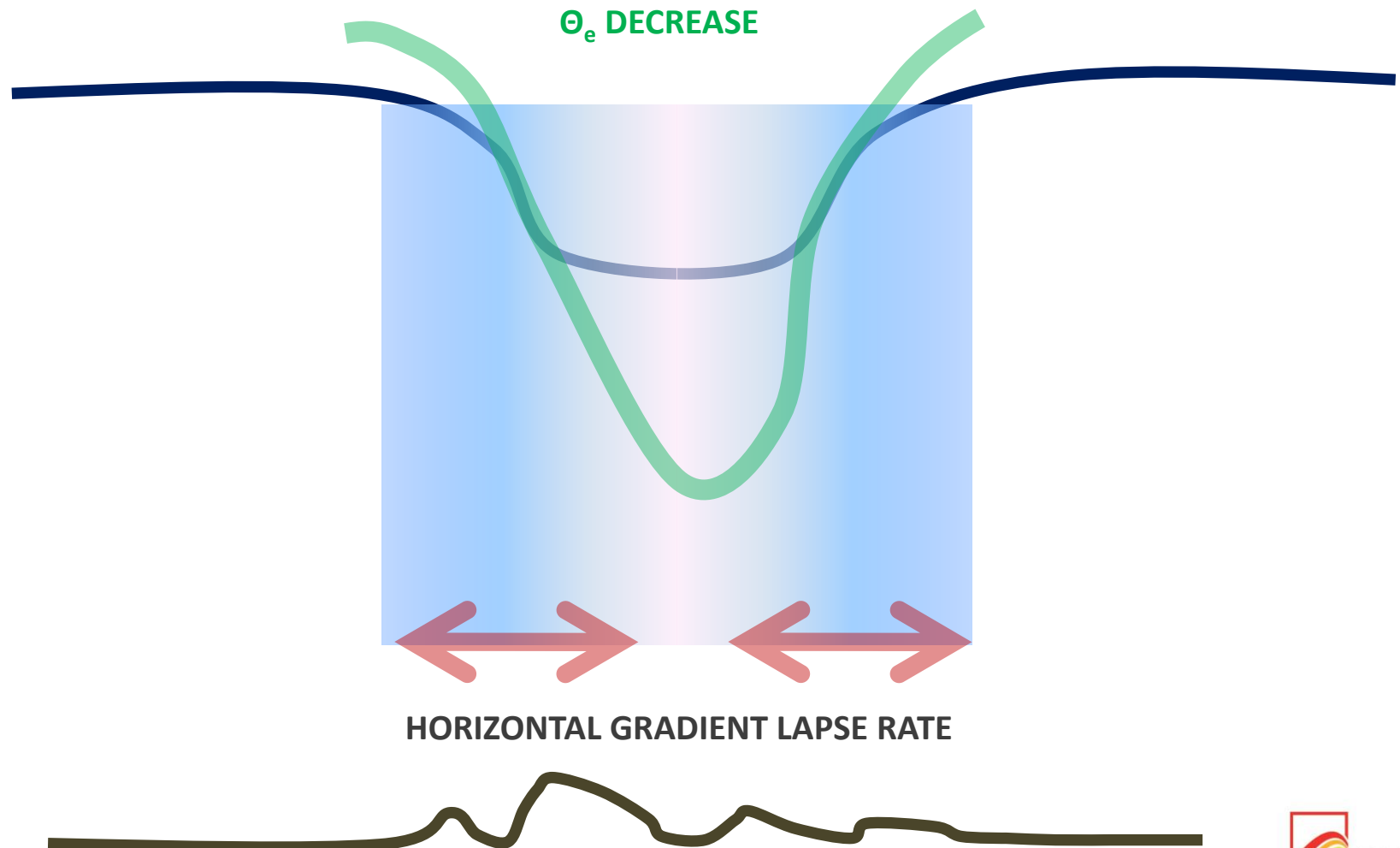
Lid-collapse by:

- sufficient boundary-layer forcing  
convergence, orographic lift is needed to overcome the lid and release CI
- large-scale uplift  
may contribute to weaken the lid
- *Increase of vertical wind speed, vertical shear, additional release of symmetric or inertial instabilities?  
“tip the scales?”*

## Assumption: characteristics at UTMG 1

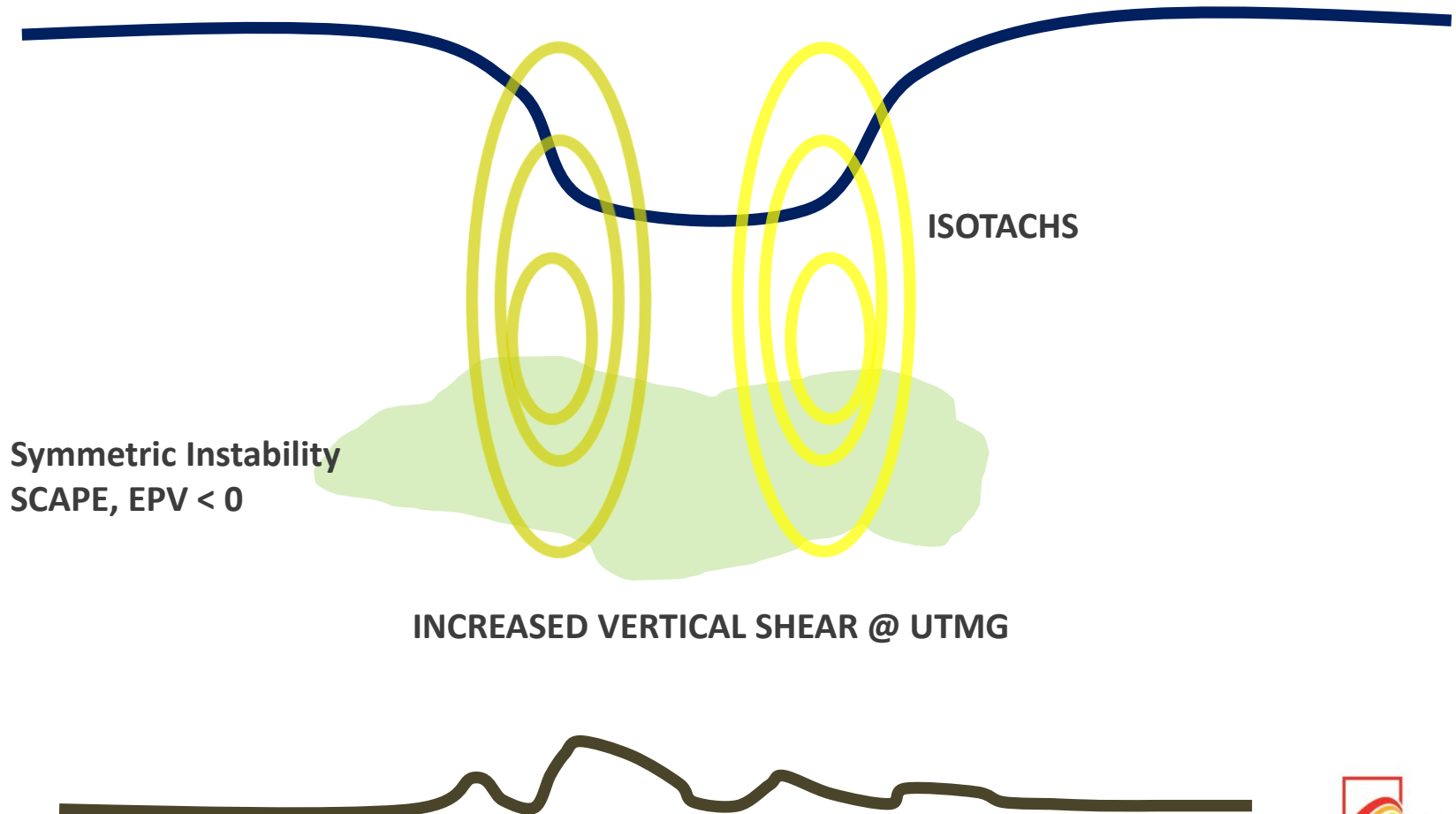


## Assumption: characteristics at UTMG 2





## Assumption: characteristics at UTMG 3





## Types of Mesoscale Instabilities

Static instabilities PI / CI:

- Air parcel is displaced and accelerated vertically

Inertial instability II:

- Air parcel is horizontally accelerated, away from an equilibrium position

Symmetric instability SI / PSI / CSI:

- May occur in an atmosphere that is statically and inertially stable
- Air parcel may be accelerated away from its original position if it is displaced, along a slantwise path

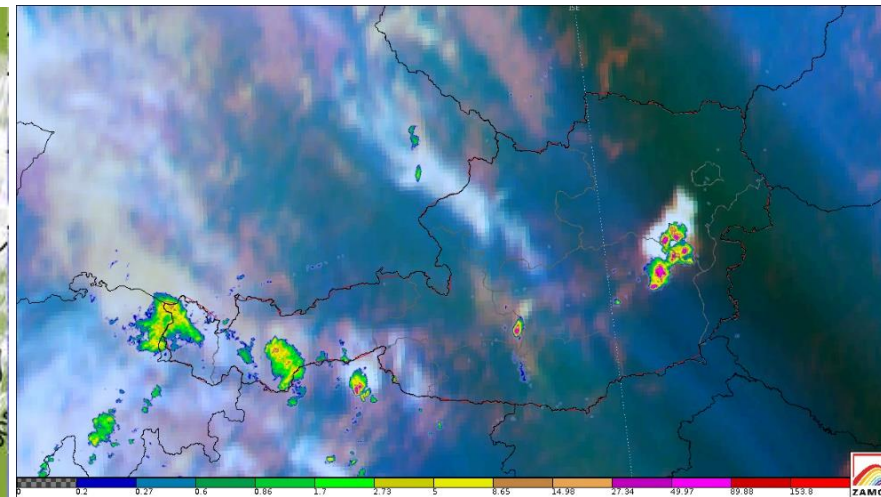
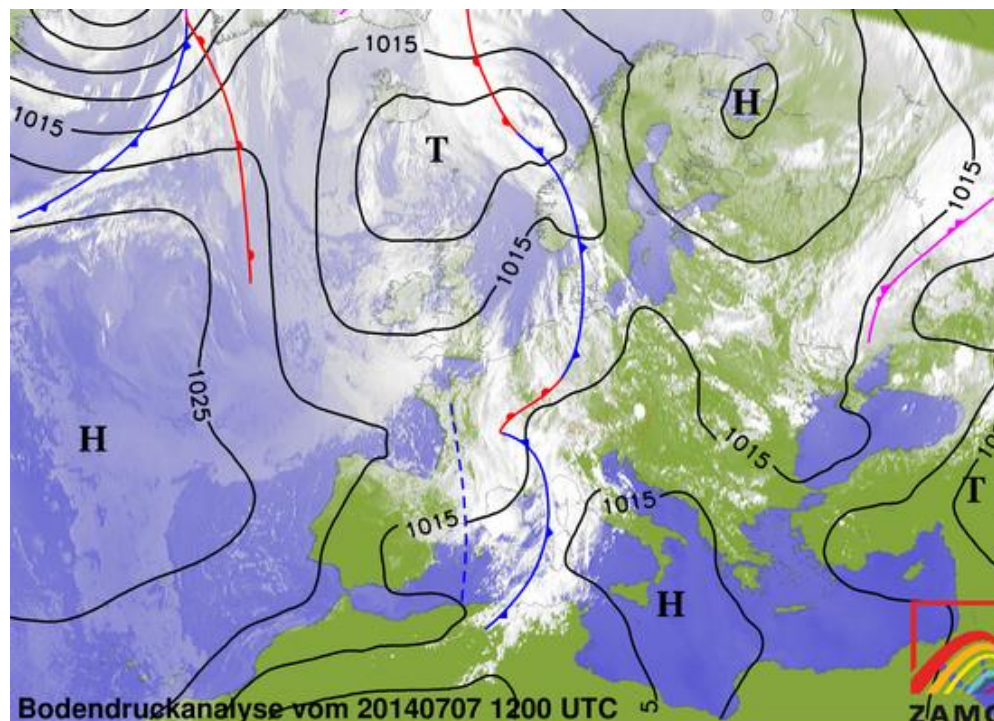
Schultz and Schumacher, 1999:

- Coexistence of CI/PI and CSI/PSI may lead to both gravitational and slantwise convection
- Gravitational convection:  $\text{ms}^{-1}$
- Slantwise convection:  $\text{cms}^{-1}$  -> "tip the scales?"

- Indicators for conditional (CI), inertial (II), and conditional symmetric instability (CSI) can be found along the UTMG:
  - Horizontal and vertical gradients of relative humidity RH,
  - Horizontal and vertical gradients of  $\Theta_e$
  - Horizontal and vertical wind shear, gradients of horizontal wind speed
  - negative Equivalent Potential Vorticity EPV (Mc Cann, 1995), between 900 and 700 hPa
- Existence of EPV minima and SCAPE maxima near the first onset of DMC

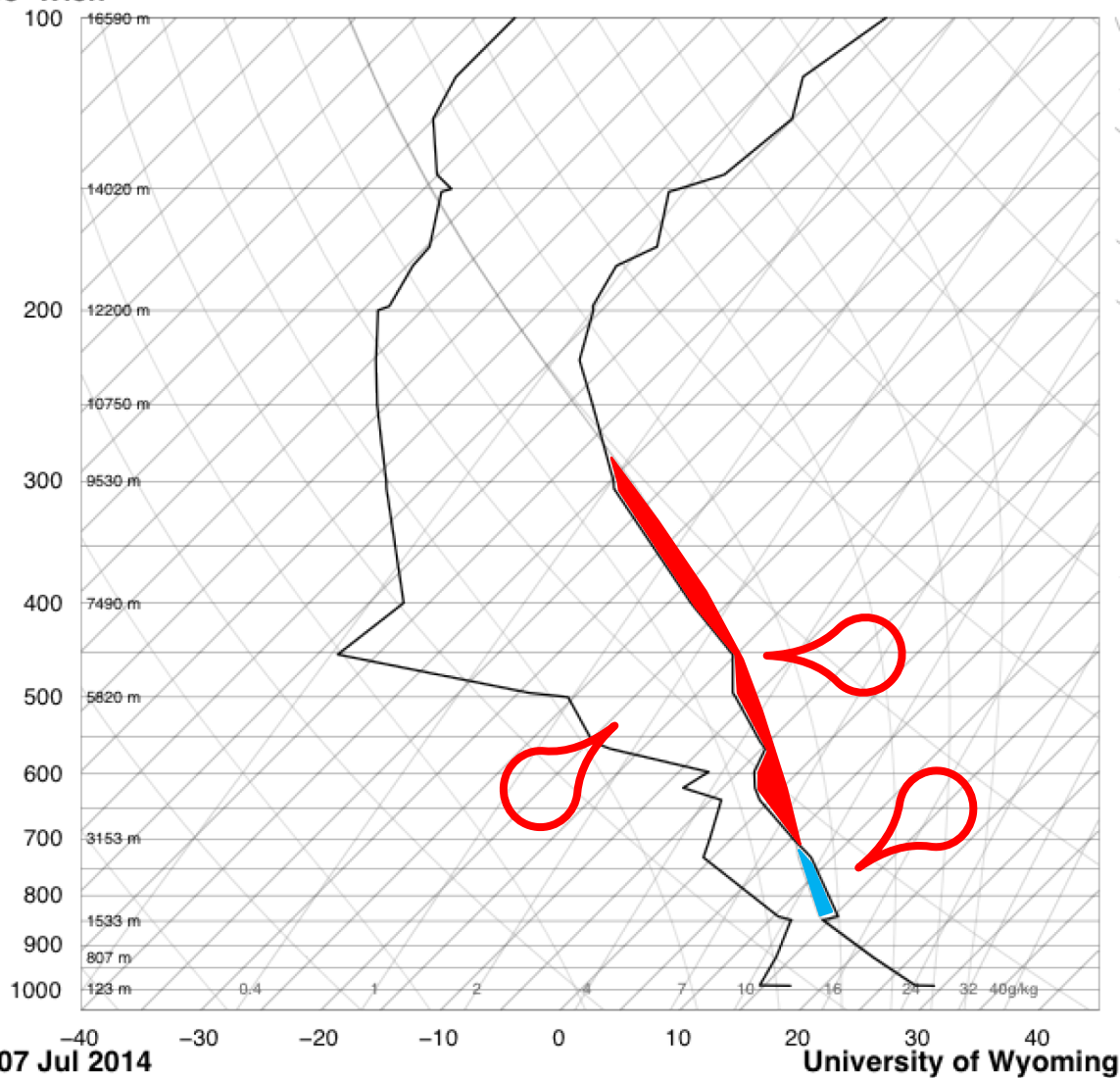
## Case Study, Eastern Austria: 7 July 2014, 0800-1200 UTC, Kainz 2016

- WRF Simulations:
  - 2 nested (ECMWF) domains,
  - resolution 20 / 0,8 km (D1 / D3),
  - PBL: MYJ; MP: WSM6; LS: RUC-3;  $\omega$ :Grell3D (only in D1)



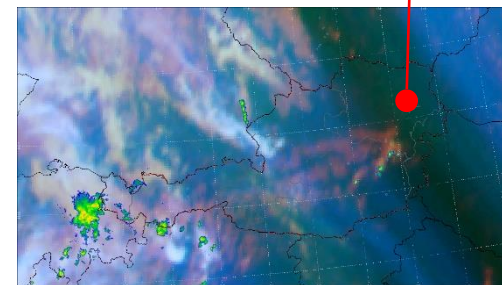
# Proximity Sounding

11035 Wien



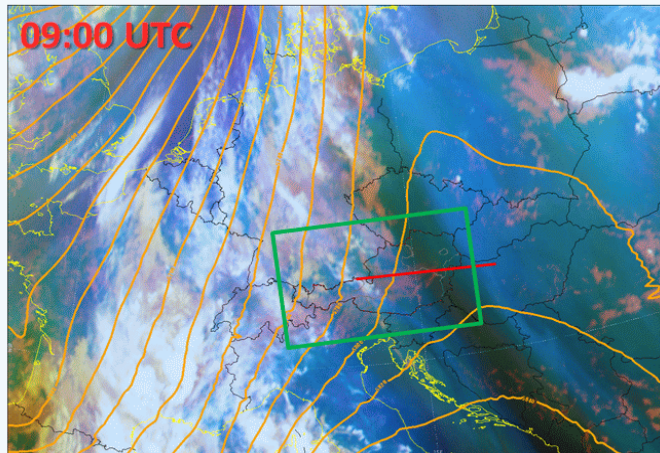
SLAT 48.25  
 SLON 16.36  
 SELV 200.0  
 SHOW -1.18  
 LIFT -1.72  
 LFTV -2.15  
 SWET 170.9  
 KINX 30.90  
 CTOT 23.30  
 VTOT 26.10  
 TOTL 49.40  
 CAPE 395.0  
 CAPV 479.5  
 CINS -78.1  
 CINV -53.4  
 EQLV 273.5  
 EQTV 273.0  
 LFCT 703.6  
 LFCV 719.1  
 BRCH 180.8  
 BRCV 219.5  
 LCLT 284.8  
 LCLP 819.9  
 MLTH 301.5  
 MLMR 10.69  
 THCK 5697.  
 PWAT 32.35

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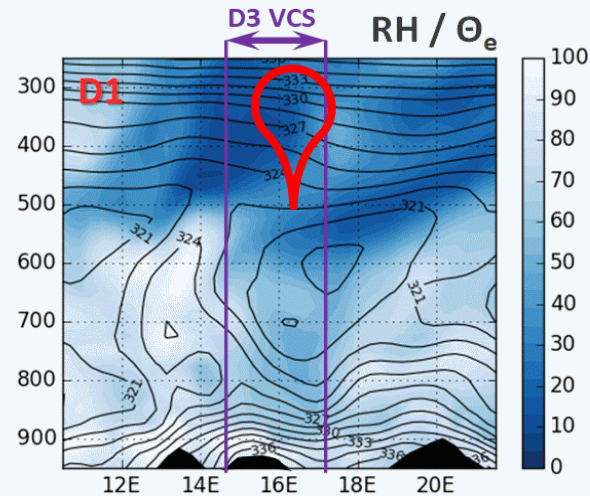




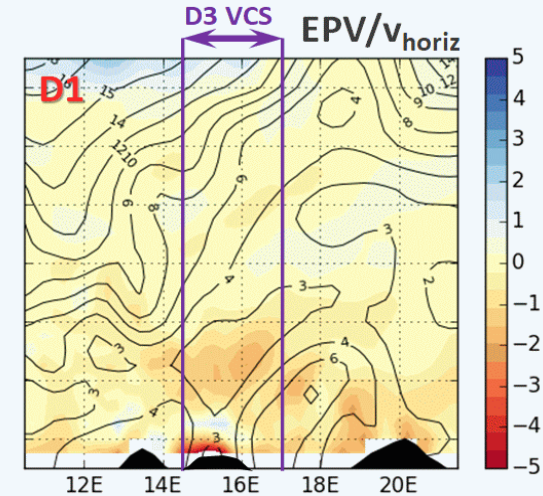
# Indicators for symmetric instabilities 1



7 July 2014, 0900 UTC,  
MSG RGB: IR 8.7 $\mu$ m (red)/ WV 7.3 $\mu$ m  
(green) / WV 6.2 $\mu$ m (blue);  
orange lines: 500 hPa Geopotential;  
red line: VCS D1;  
green square: D3, D1 bigger than  
image plane

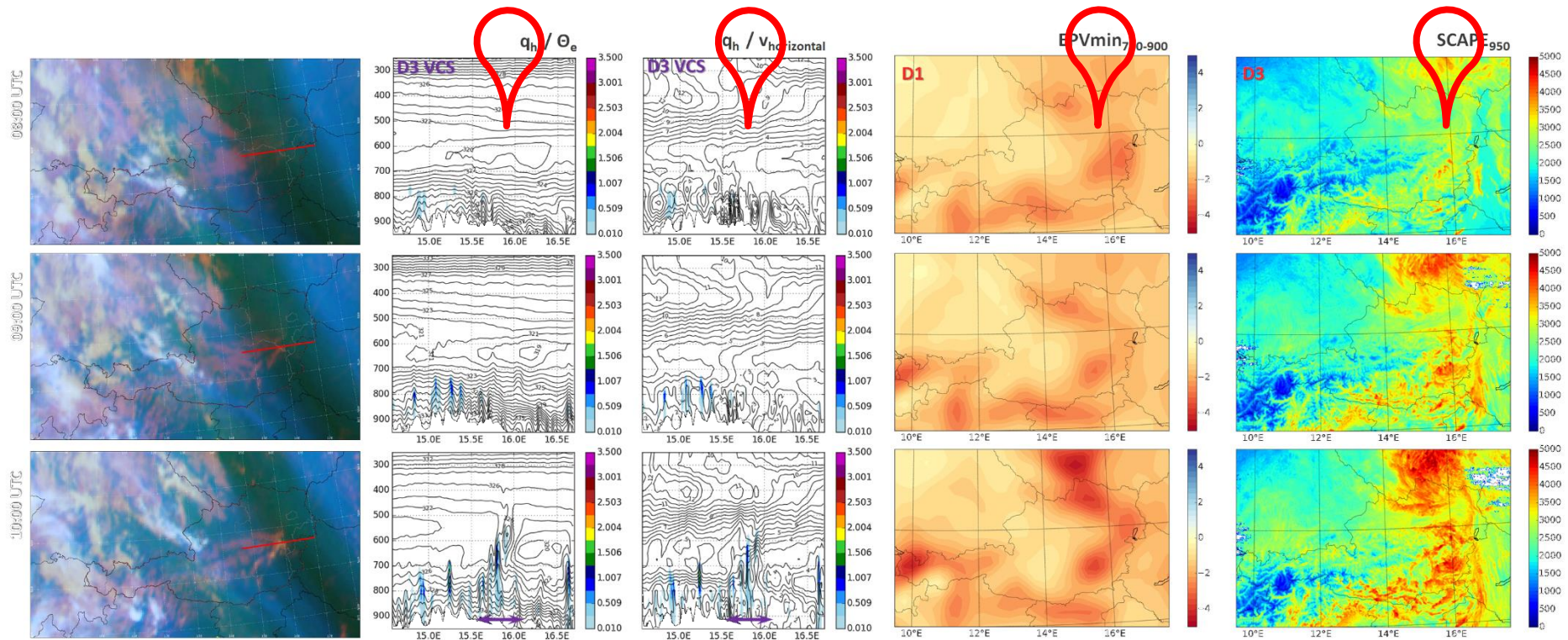


D1 (3), resolution 20 km,  
Relative Humidity [%] (shaded),  
 $\Theta_e$  [K] (lines)



D1 (3), resolution 20 km,  
EPV [PVU] (shaded),  
Horizontal Wind Component [m/s]

# Indicators for symmetric instabilities 2

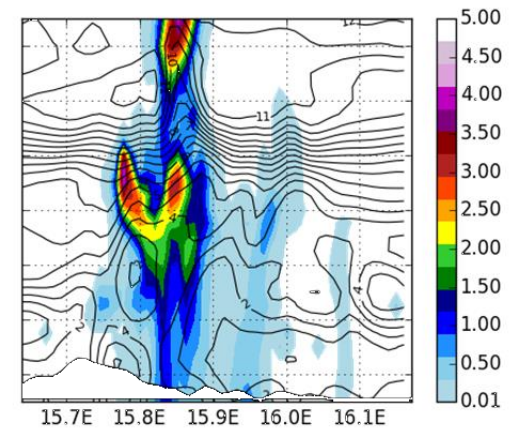
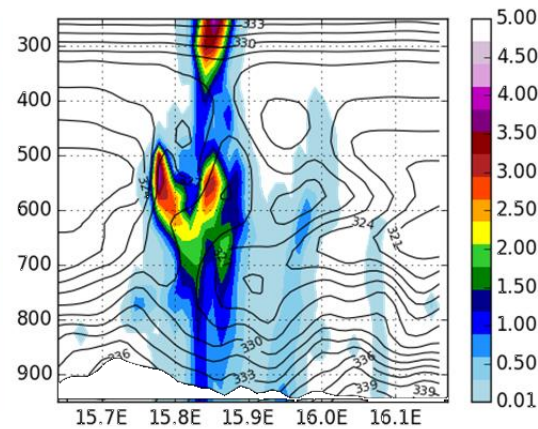
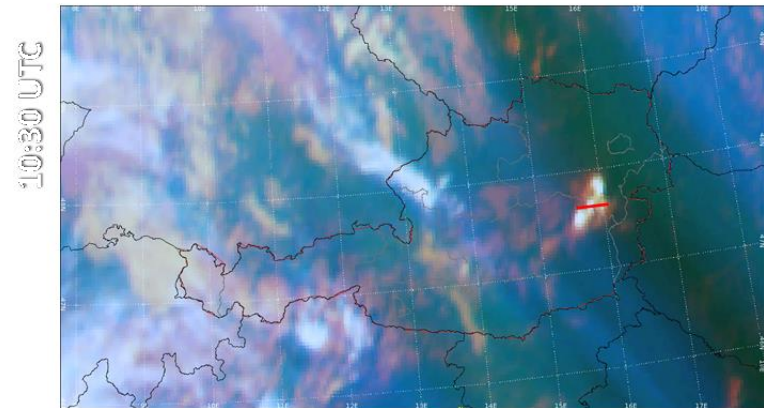


WRF simulation D3, VCS:  
 Total hydrometeor mixing ratio [g/kg];  
 $q_h$  (shaded)  
 Left: Equivalent Potential Temperature  
 $\Theta_e$  [K], (lines, left side)  
 Right: Horizontal component of true wind  
 speed  
 [ms<sup>-1</sup>](lines, right side)

Left: Negative EPV, 900-700 hPa  
 [PVU=10<sup>-6</sup>\*K\*m<sup>2</sup>/kg\*s],  
 Right: Symmetric Available Potential Energy SCAPE [J/kg],  
 integrated from 950 hPa, after Dixon (2000)

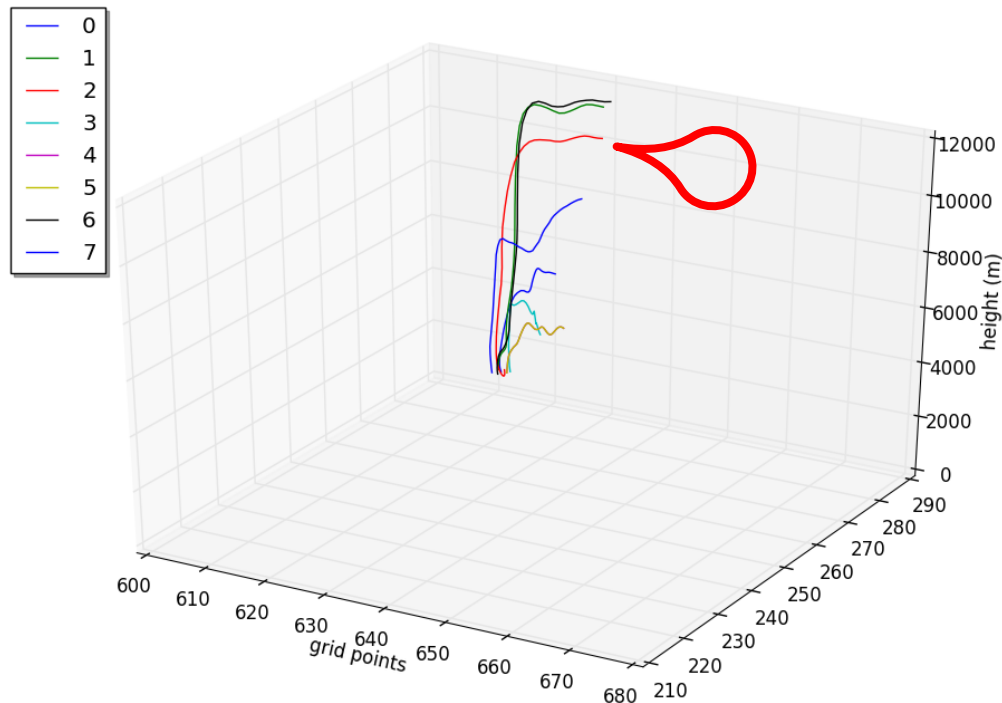


# Resulting DMC



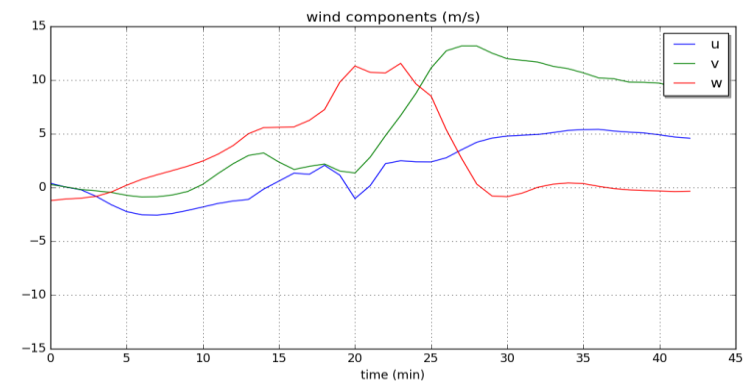
# Parcel Trajectory Analysis

Calculation of eight 45-min forward trajectories, initiated at 2000 m above ground



Despite existing indicators for symmetric instabilities, no clear signal for slantwise convection in the model

- Wind components along trajectory 2 (red) as time series (below)
- vertical component seems to dominate, indicating upright convection
- After approx. 25 minutes, decrease of vertical component, increase of horizontal components while updraft reaches tropopause





# Conclusions

- DMC development from pre-existing shallow convection in the vicinity of pronounced horizontal and vertical  $\Theta_e$  gradient and wind shear
- Distinct indication of SCAPE and lower-mid-level negative EPV, WRF simulations showed, that initial shallow convective towers reaching these levels before DMC evolution succeeds
- Coexistence of CI, II and CSI along the UTMG seems plausible
- Release of SCAPE indicates potential for slantwise convection seen in model layers between mid- and low levels of the troposphere
- Slantwise convection mechanism not visible in model trajectories, gravitational convection predominates trajectory path

## Remaining Questions:

- Impact of (advected) increasing wind shear in pre convective environments on lift and / or lid removal
- MTG application facilities will certainly lead to new approaches and results

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[https://journals.ametsoc.org/view/journals/mwre/127/12/1520-0493\\_1999\\_127\\_2709\\_tuamoc\\_2.0.co\\_2.xml](https://journals.ametsoc.org/view/journals/mwre/127/12/1520-0493_1999_127_2709_tuamoc_2.0.co_2.xml)

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