Fundamentals of Satellite Precipitation Estimation







































International Precipitation Working Group

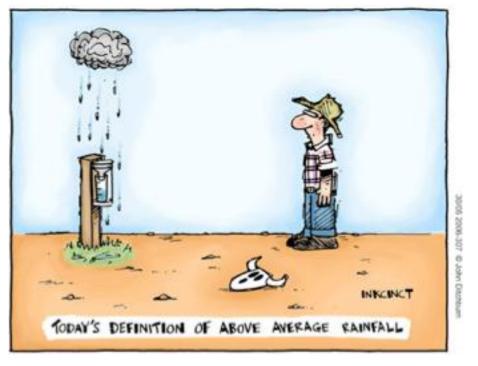






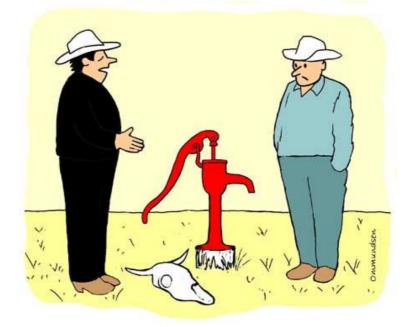
Precipitation...what are we talking about?





...meteorology

Is it humour or not?



...hydrology

"It ain't nothing that a century of rainfall won't fix."



©John Hambrock.

...climate



...common sense...

Rainfall Requirements Depend Upon the Application

OMonthly average rainfall, global (±60° latitude), pentad-type (e.g. 2.5-degree boxes), over land/ocean (Climatic shifts? Desertification?)

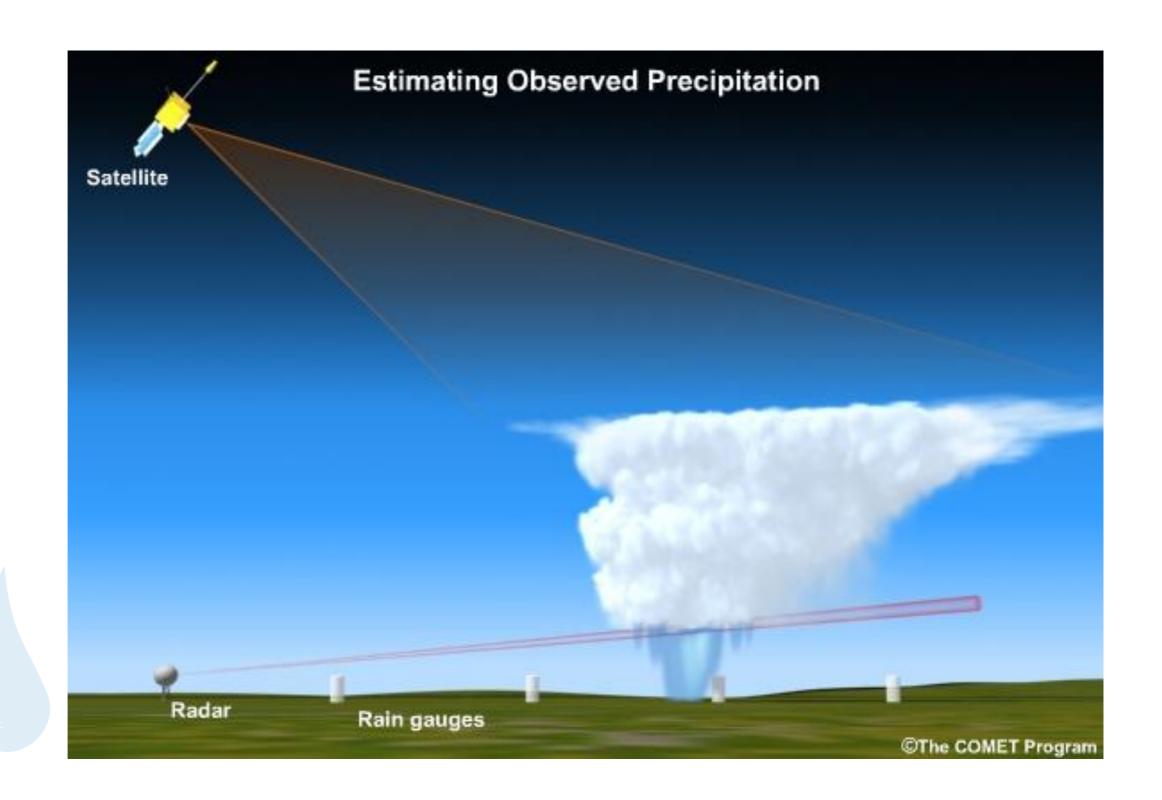
ODaily accumulated rainfall and snowpack, many stations over a watershed (When do I release water from a reservoir? Allocate water distribution?)

O"Single point forecasts" (Will it rain on my wedding day? e.g., mid-afternoon of September 14, 2014 at location 36.84451N and 121.53481W)

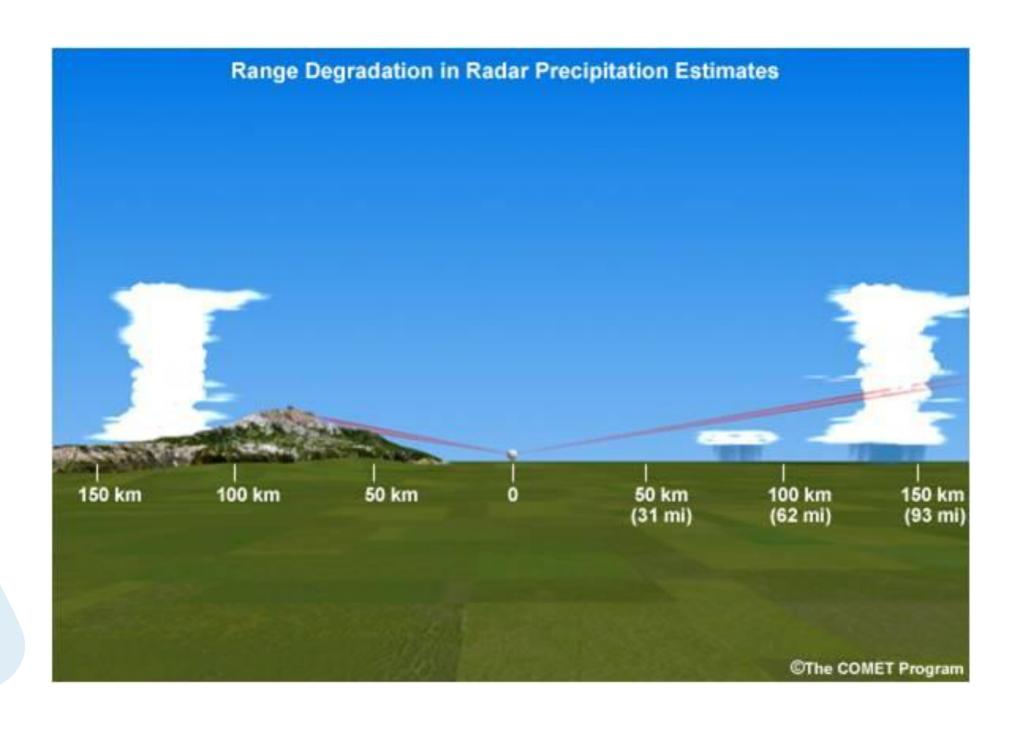
O"Realtime" global or regional analysis of rainrate at the best possible horizontal resolution (hydrological models)

O5-minute updates of point rainfall inside an area (e.g., 10⁵ km²) during the lifetime of a thunderstorm or landfalling hurricane away from coastal radars (Should coastal or low-lying areas be evacuated? Temporarily relocate naval fleet to safe harbor?)

OAny indications that this winter is associated with El Niño conditions? (An energy company, a tree removal company, emergency services)



Is the radar the perfect instrument for precipitation?



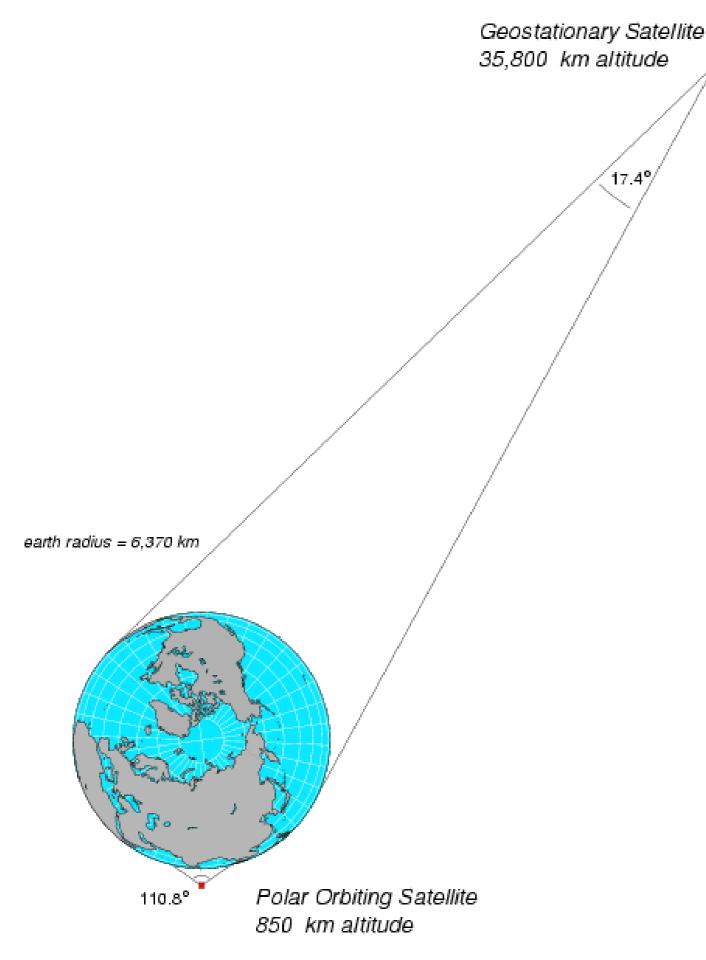
Density of rain gauges



Taking all of the rain gauges combined would equal about the surface area of two basketball courts!

Satellites, orbits,....





Two Main Orbits for Earth Orbiting Weather Satellites

Geostationary (GEO)

As the satellite orbits, the Earth appears in a fixed position (relative to an observer on the satellite)

Low Earth-Orbiting (LEO)

Earth rotates under the satellite as the satellite orbits. Satellites are much closer to Earth.

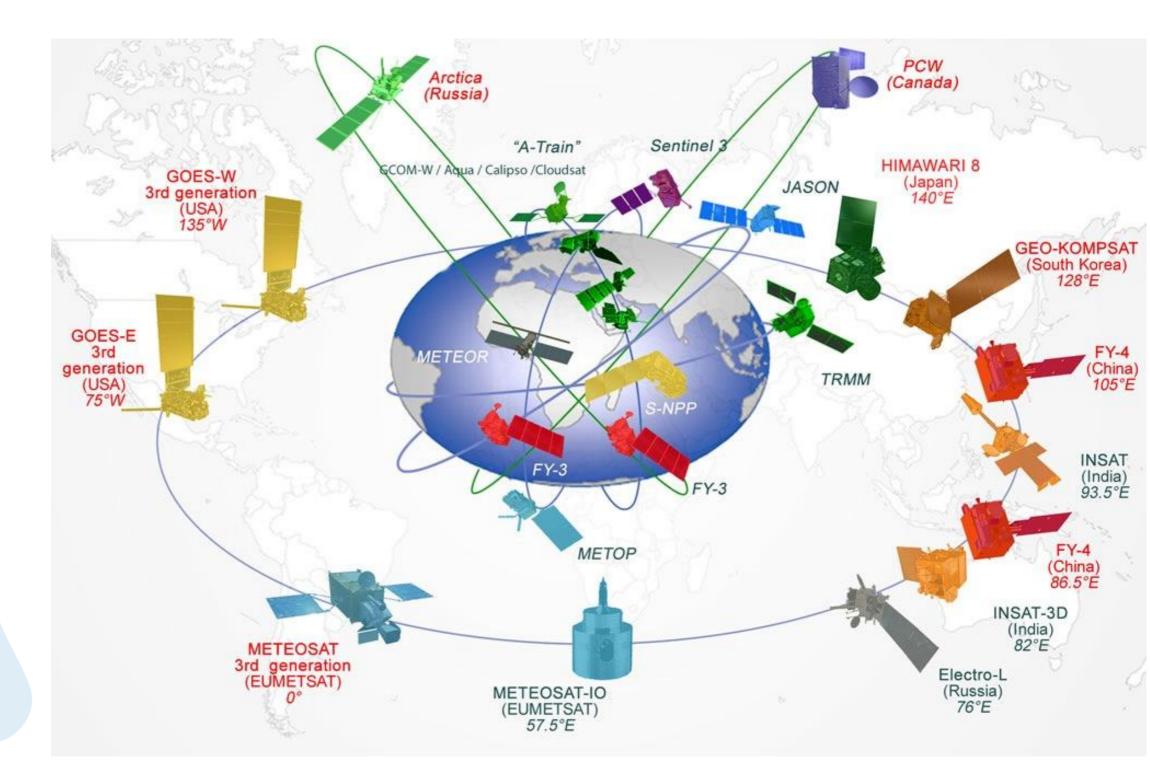
LEO orbit characteristics

- 1. Global coverage with 2 daily overpasses over the same area at very high spatial resolution (1 m 50 km).
- 2. Great variety of sensors on the same platform, both passive (radiometers) and active (radar, lidar).
- 3. Frequent monitoring of polar regions (once per orbit).
- 4. Heliosynchronous satellites guarantee an overpass at the same local time.

GEO orbit characteristics

- 1. High resolution spatial sampling of the observed area (≈1 km).
- 2. Quasi-real-time sampling (15-30 min) allows for a continuous monitoring of rapidly evolving events.
- 3. Imagery is distorted with increasing latitude.
- 4. The technology for active sensor is not available at present while microwave passive radiometers are foreseen in a few years.

Operational environmental satellites



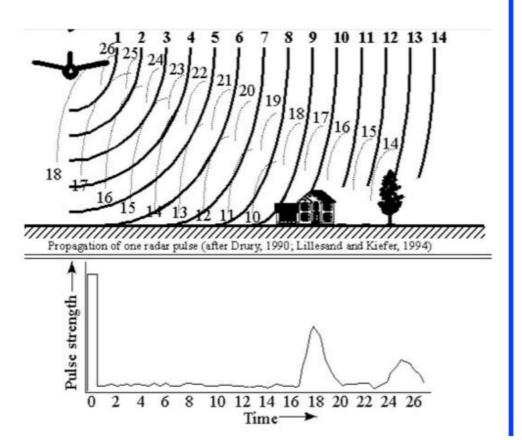


Remote Sensing Fundamentals

Active Remote Sensing

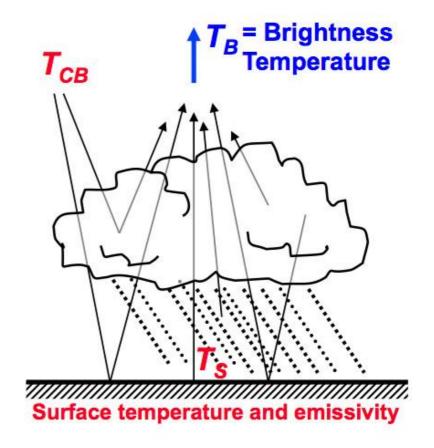
Source: Instrument pulse,

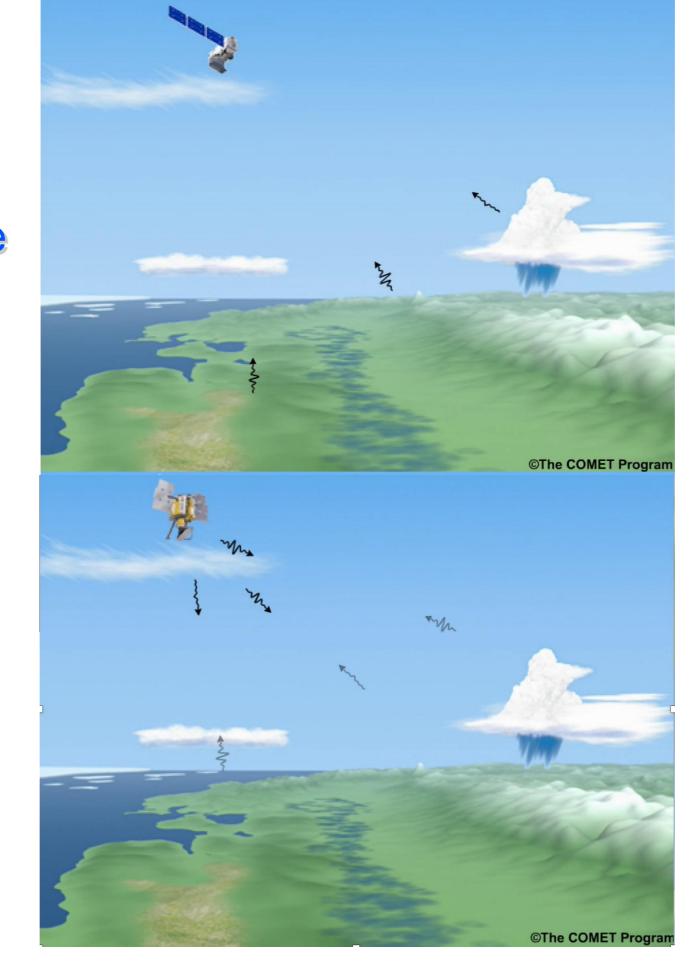
Needs power to operate



Passive Remote Sensing

Sources: surface emission, cosmic background, rain emission

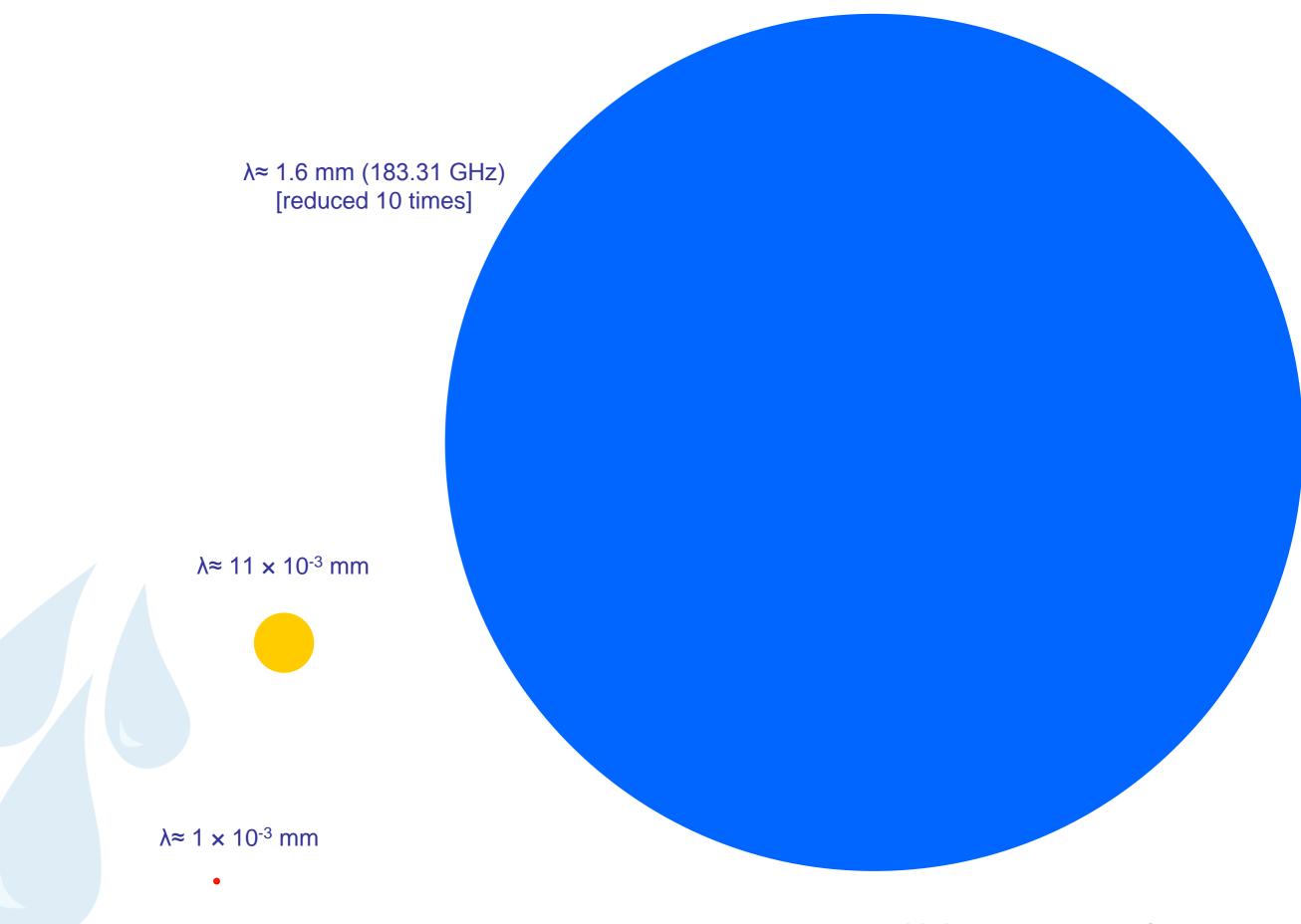




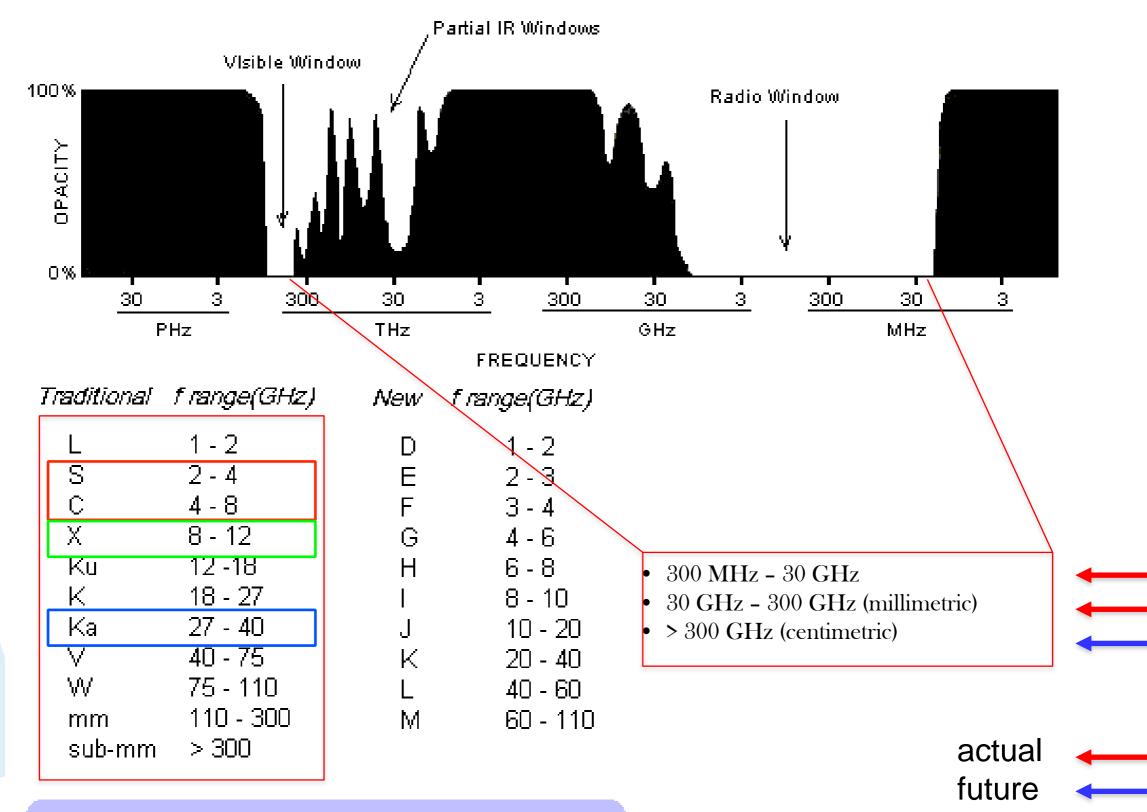
Passive

Active

From visible (VIS) to microwave (MW)



ATMOSPHERIC WINDOWS IN THE ELECTROMAGNETIC SPECTRUM.



MICROWAVE RADIO BANDS

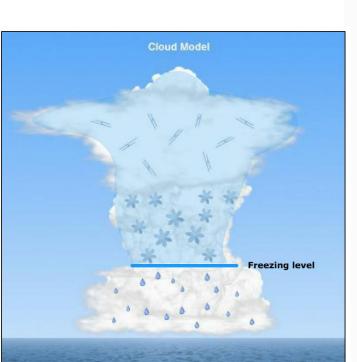


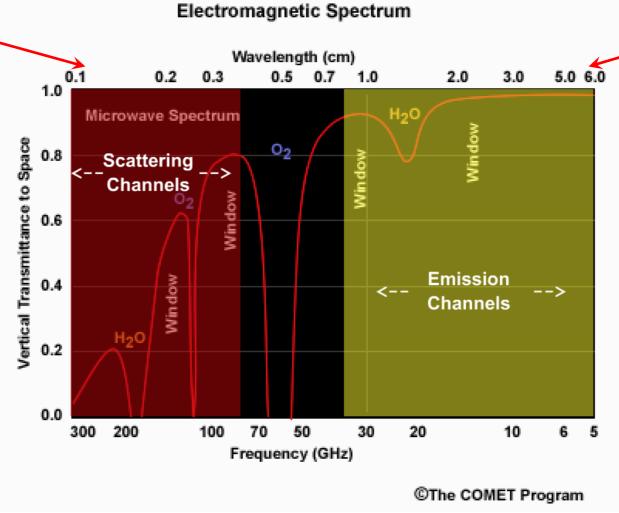
Infrared channels

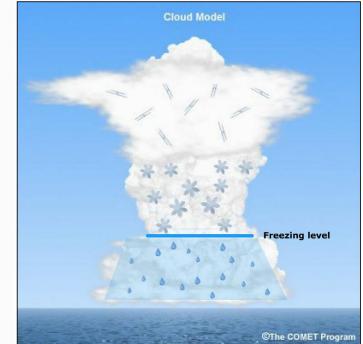
Infrared instruments sense energy emitted from cloud tops. IR-based cloud-top temperatures can be used to approximate precipitation rates, but no information comes from the layers below the cloud top.

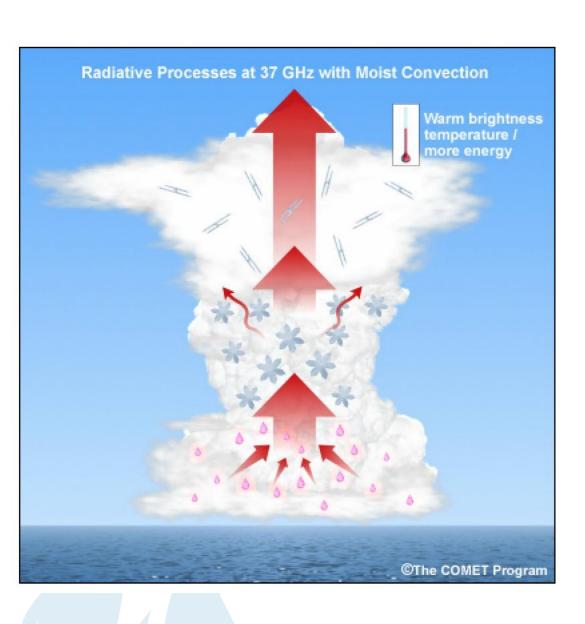
Higher-frequency channels
Emission from
small raindrops, and scattering
and attenuation by precipitation-sized ice hydrometeors

Lower-frequency channels
Emission from
medium-to-large-sized raindrops

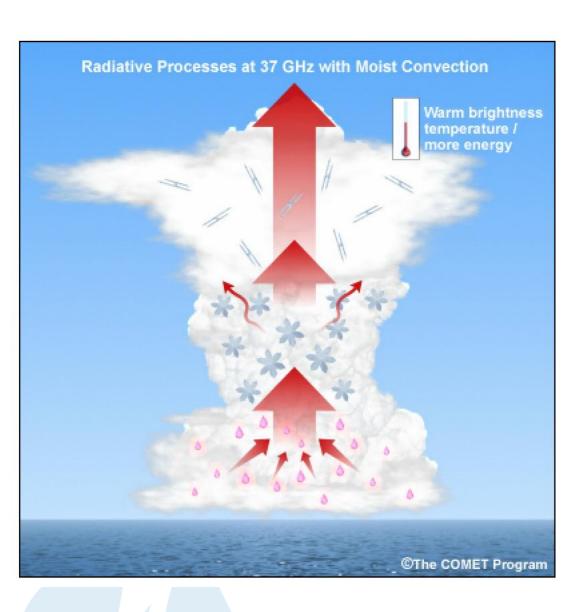


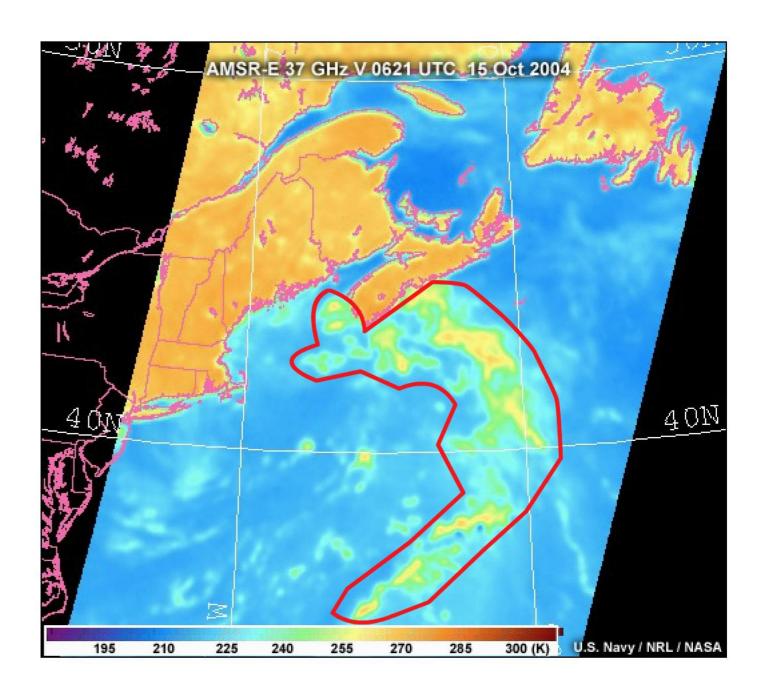






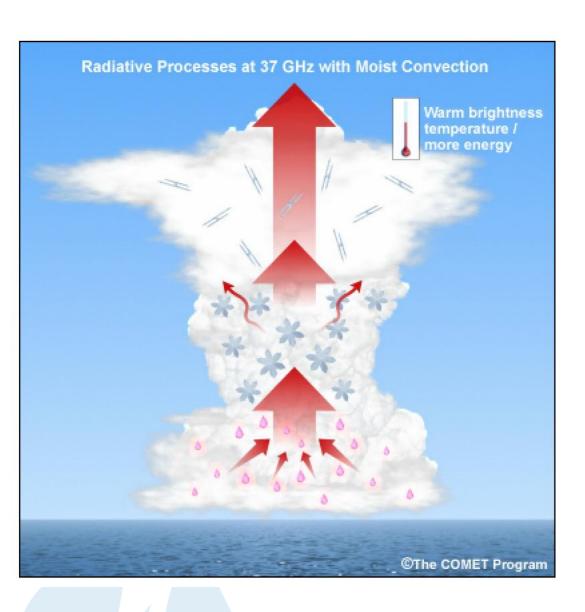
Low-frequency channels Low-frequency channels sense energy from the surface that is then increased by emissions from cloud water and rain drops.

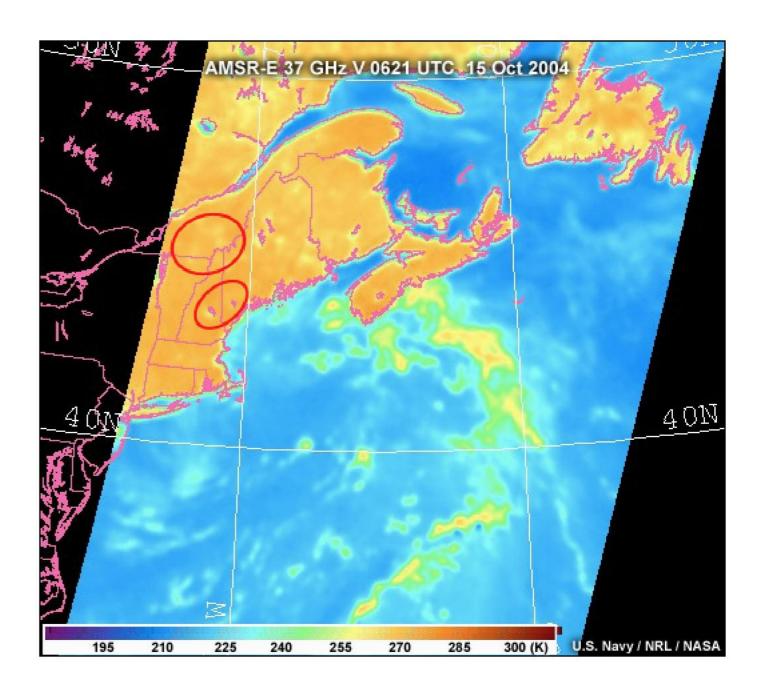




Low-frequency channels

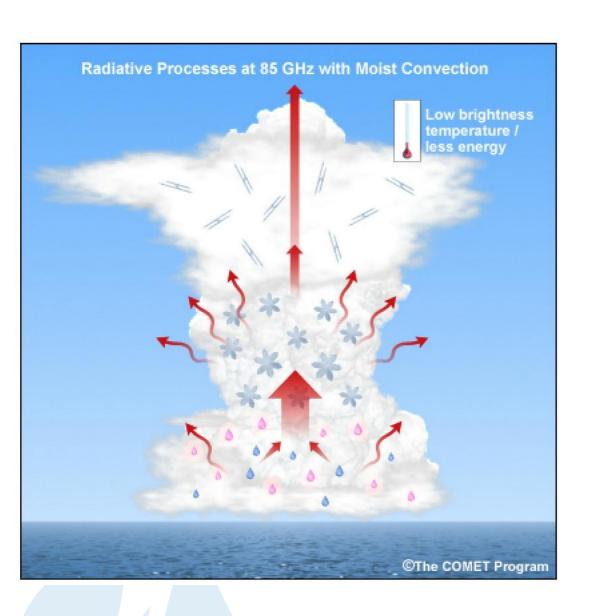
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Low-frequency channels

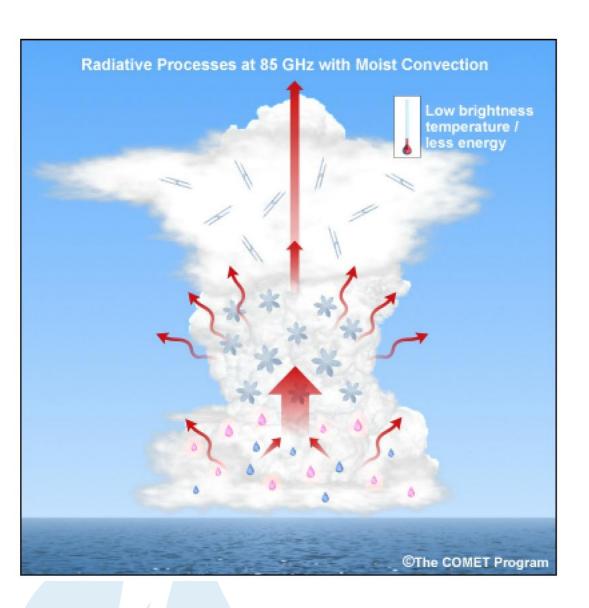
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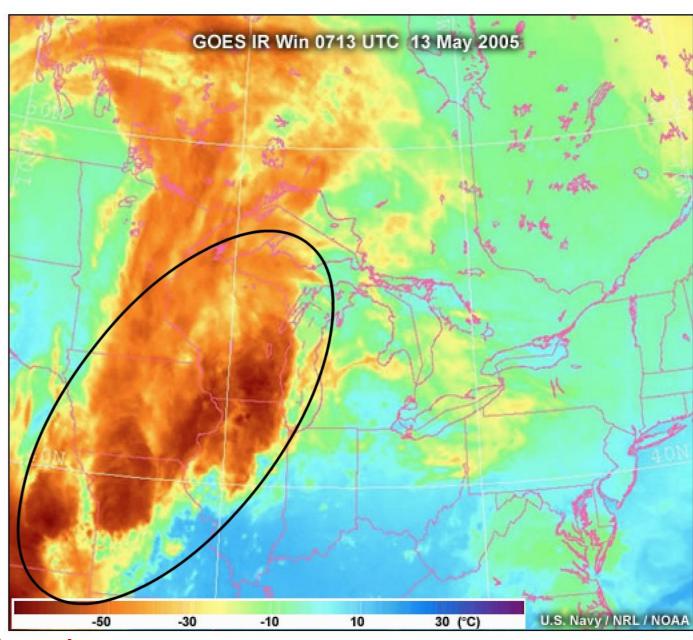


High-frequency channels

Higher-frequency channels are sensitive to Scattering and attenuation of energy by rain and ice hydrometeors, and emission from small raindrops.

Precipitation-size ice particles act to cool the observed Brightness temperatures.

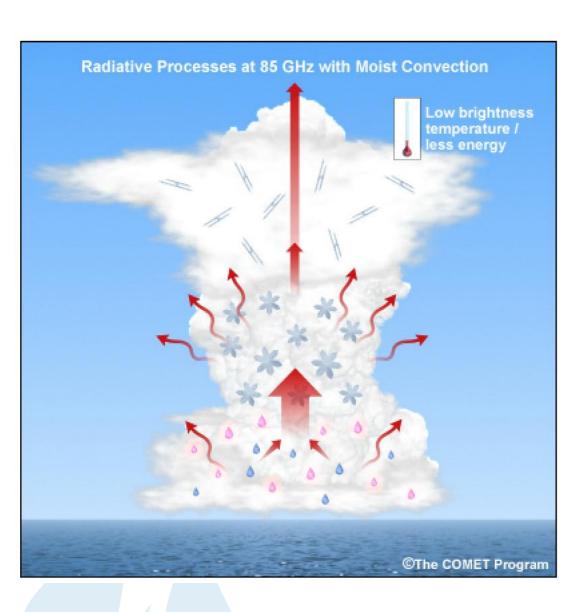


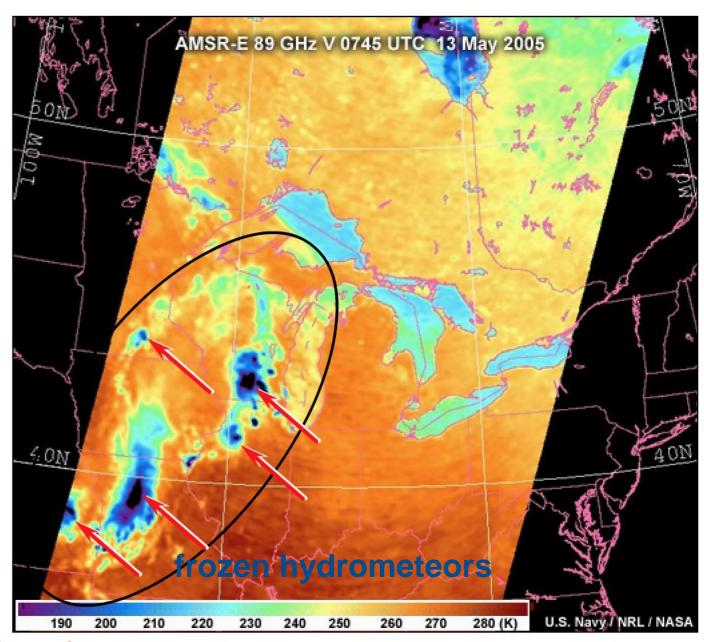


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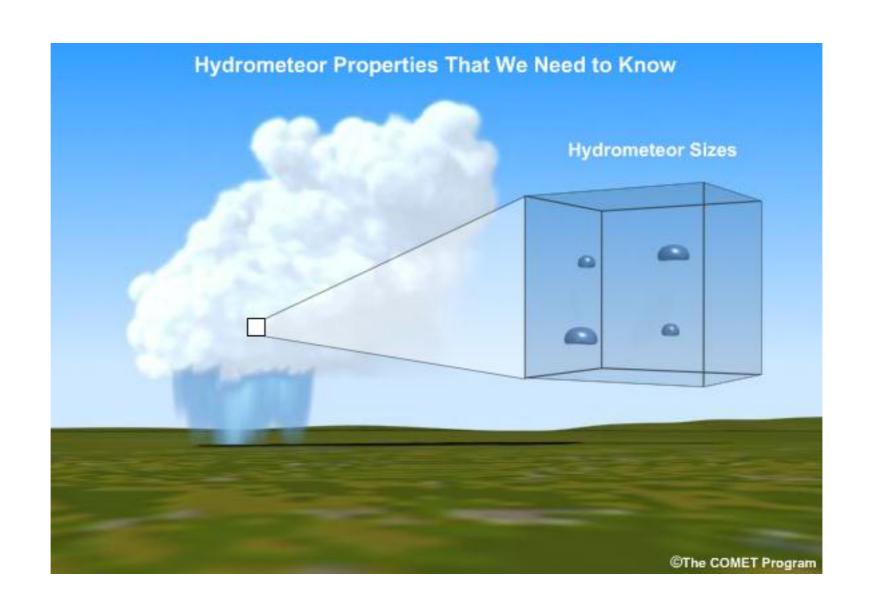
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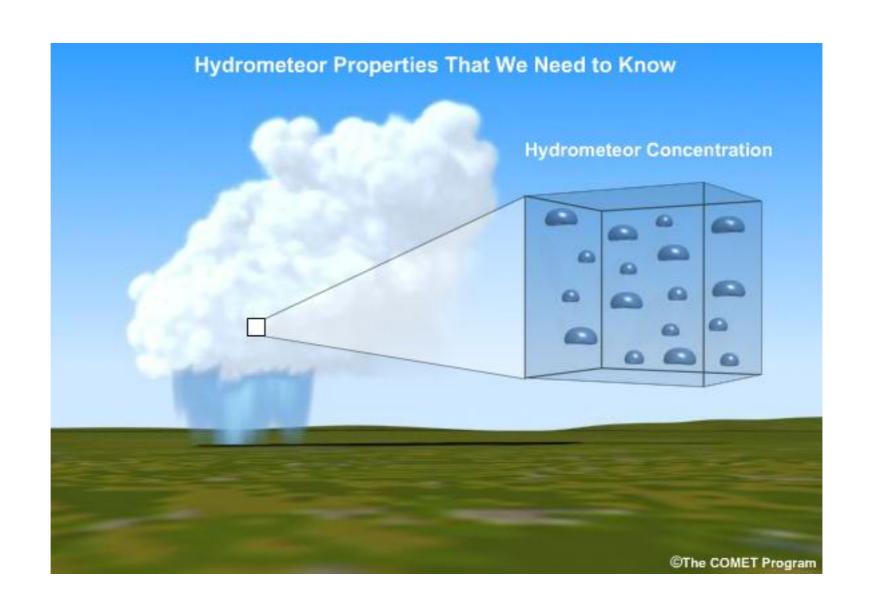
Precipitation-size ice particles act to cool the observed Brightness temperatures.

What do we need to know?

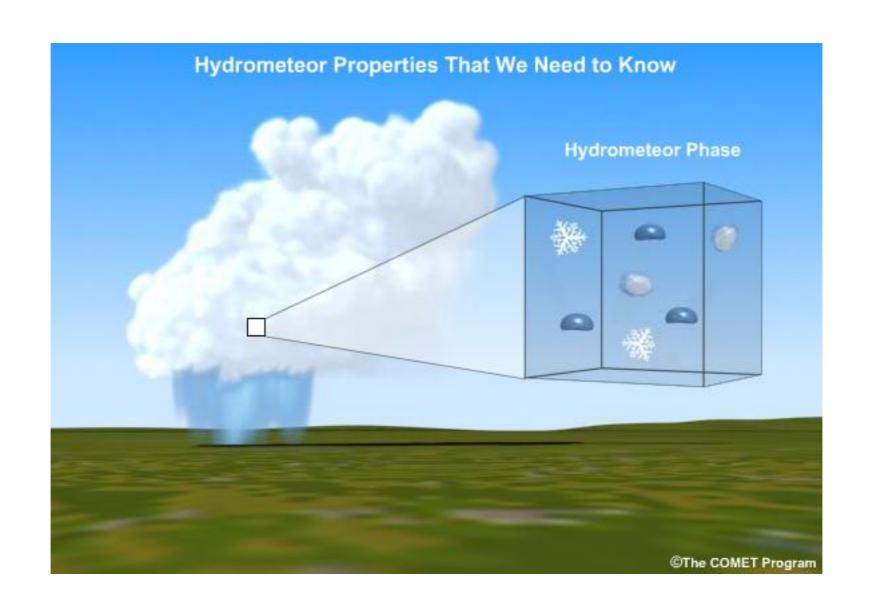




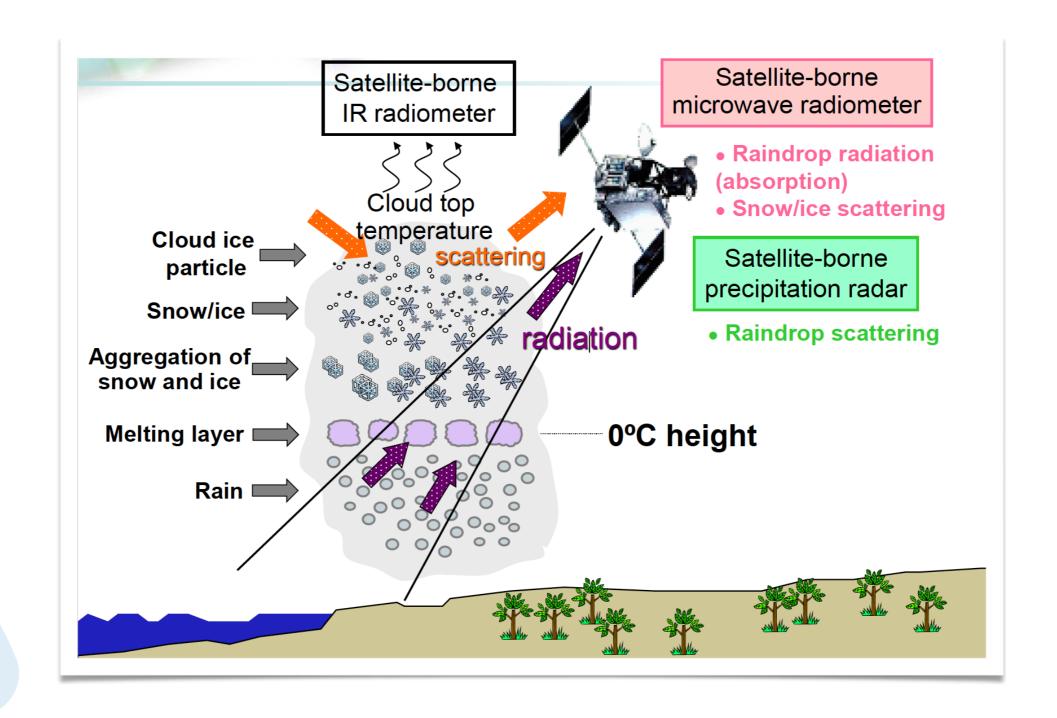


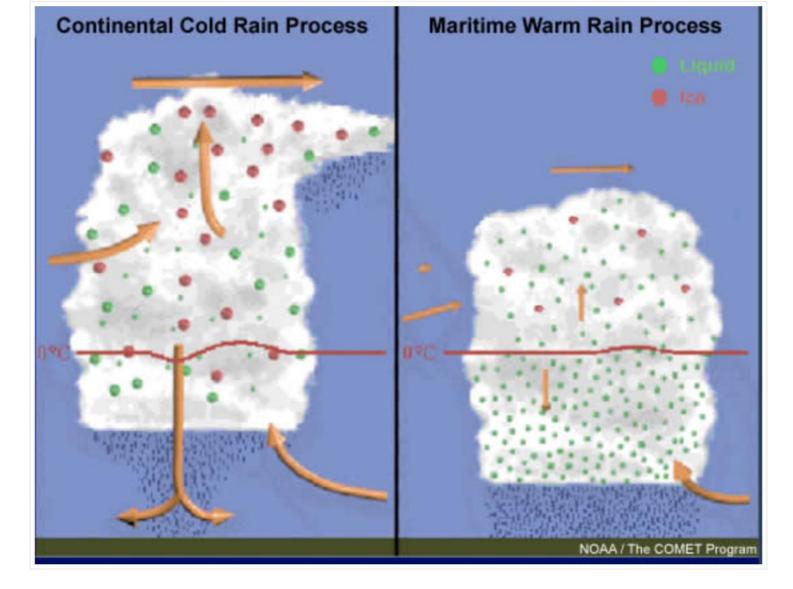










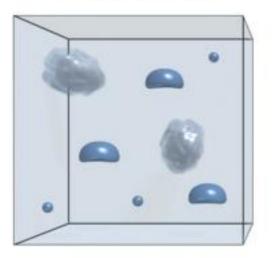


In some situations, especially in maritime tropical environments, precipitation can be produced predominately via the warm rain process. Here, precipitation particles mainly grow in the liquid phase at altitudes where the temperature is greater than 0°C . By contrast, the very common cold rain process describes a situation in which precipitation particles grow mainly in the ice and snow phase and then melt on the way to the surface.

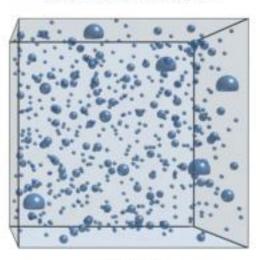
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Hydrometeor Distributions with Equivalent Reflectivity but Different Rainfall Rates

Cold Rain Process



Warm Rain Process

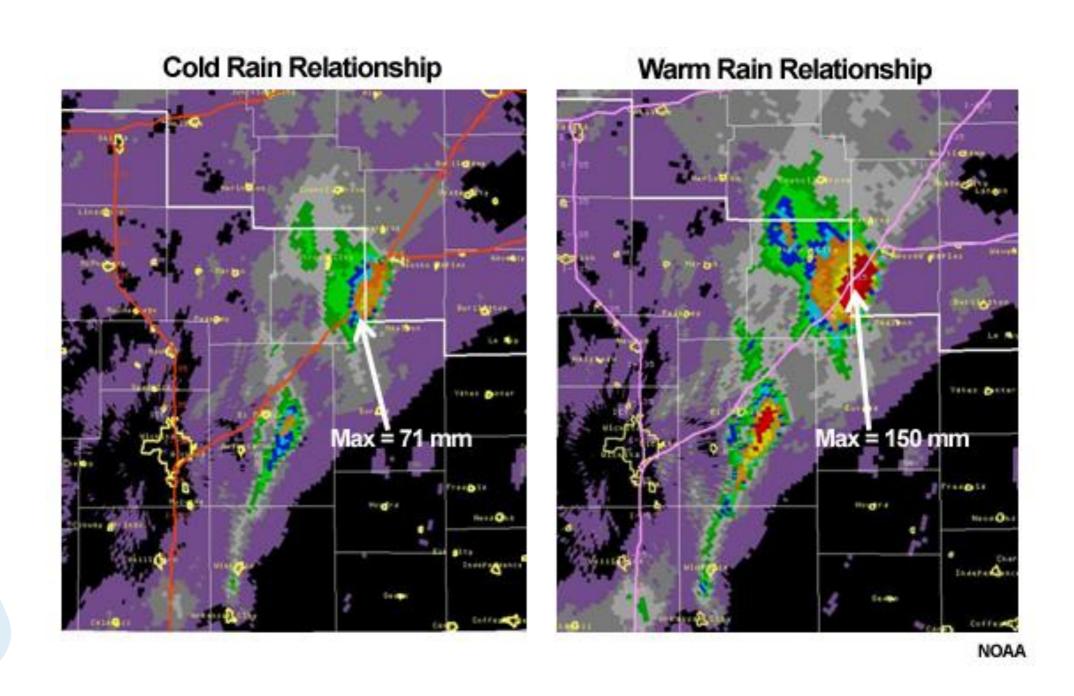


Greater rainfall rate

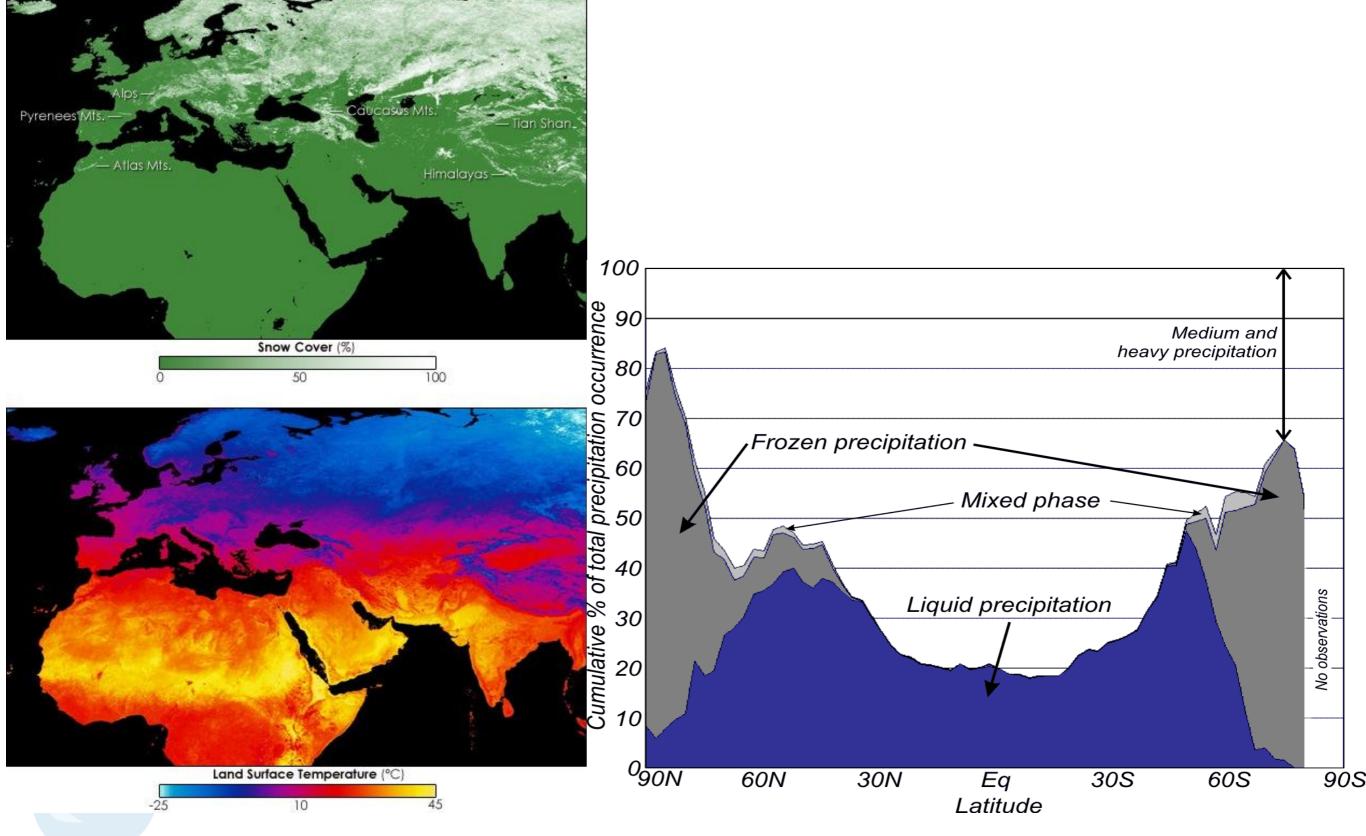
Reflectivity values are dependent on the number of precipitation particles present and the 6th power of the particle diameter. Rain rates are calculated from the reflectivity using a simple relationship that is based on empirically-derived estimates of the populations of each size of raindrop. During the warm rain processes, there are usually high concentrations of medium-to-small raindrops. By contrast, during cold rain processes, there tend to be lower concentrations of hydrometeors, and those hydrometeors can vary from small to very large.

Reflectivity Factor - Rain Rate Relationships		
Reflectivity	Continental, Cold Rain Z = 300 * R ^{1.4}	Maritime, Warm Rain Z = 250 * R ^{1.2}
45 dBZ	27.9 mm hr ⁻¹	56.5 mm hr ⁻¹
50 dBZ	63.4 mm hr ⁻¹	147.2 mm hr ⁻¹
55 dBZ	144.2 mm hr ⁻¹	384.6 mm hr ⁻¹ © The COMET Program

Is it the same precipitation episode?



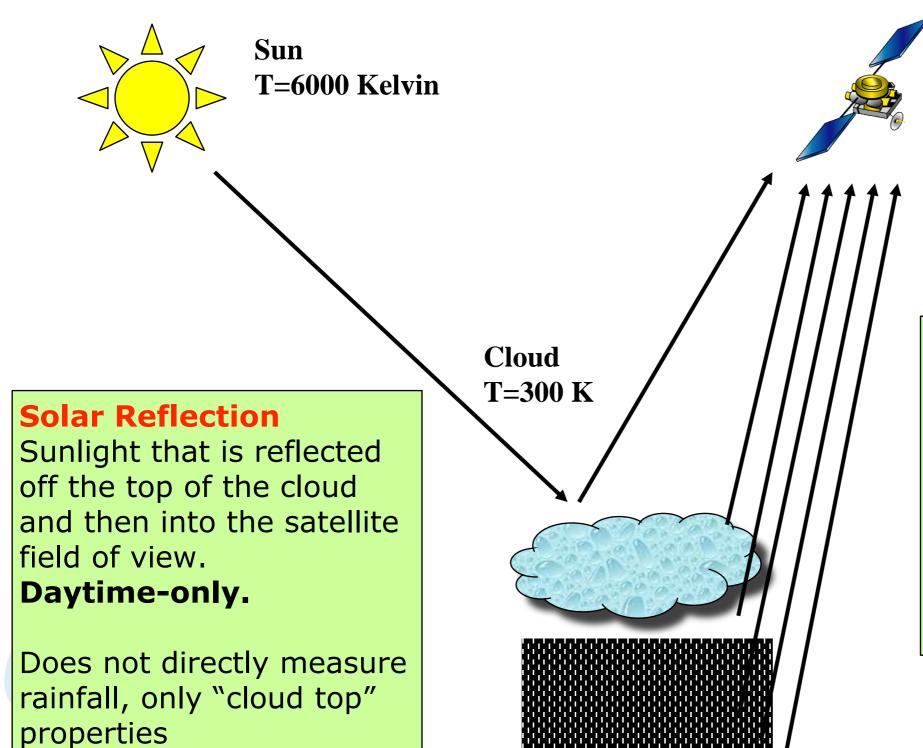
How important is it to estimate snowfall?



What are we observing and how?



Two Types of Satellite Measurements



Earth's surface (70% ocean, 30% land)

Thermal Emission

Satellite

Thermal emission that originates near the top of the cloud (infrared to near infrared) or within the cloud (microwave) and into the satellite field of view.

Day and night.

(not drawn to scale)

The Satellite "Beamfilling" Problem

We don't know the spatial pattern of the underlying rainfall at the time that the satellite flies over

Therefore, when one interprets the satellite signal (radiances), there will be a systematic underestimate of rainfall (e.g., 10 mm h⁻¹)

But it's only raining in this fraction of the sensor's field of view (e.g., 25 mm h⁻¹)

Satellite sensor receives a signal for all Earth scenes that fall within this cone ("field of view")

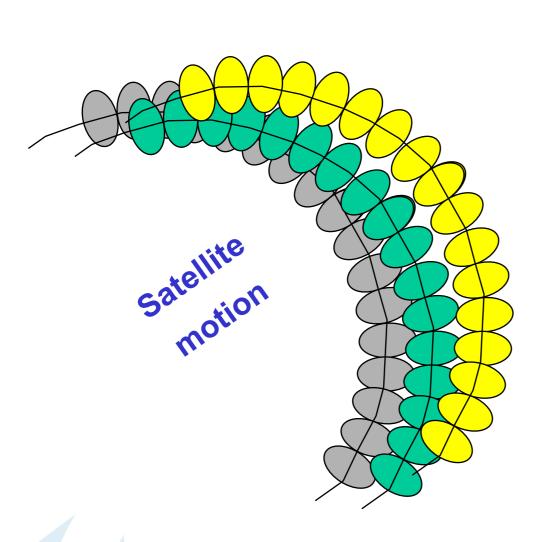
Satellite

movement

Earth's surface

(not drawn to scale)

Characteristics of the Satellite Sensor Scanning



Cross-track scanning (AMSU-B)

Pixels grow as viewing angle grows away from nadir. Pixels are spaced further apart, but oversampling in the along-track direction occurs at the edges.

Conical scanning (SSMIS, TMI, GMI)

Scan lines are segments of a cycloidal pattern. The along-track separation is the same everywhere, but the curvature causes over-sampling at the edges.

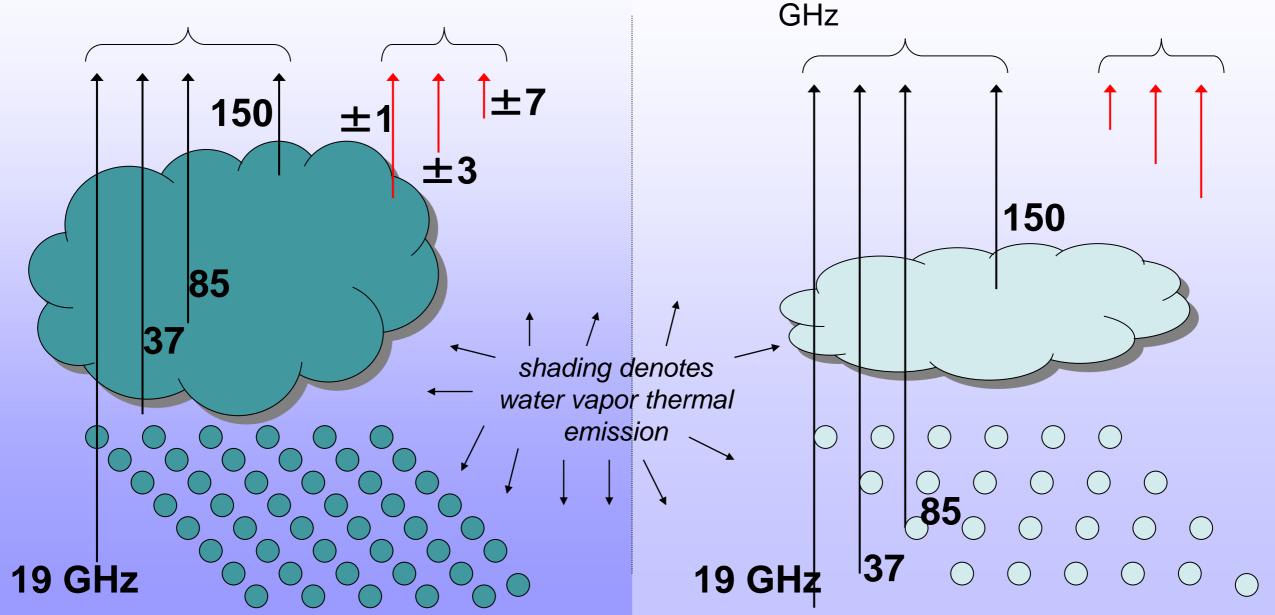
rectangular map grid

Using Imaging and Sounding Channels on SSMIS

19-85 GHz imaging channels are the most useful 183 GHz sounding channels
less useful – water vapor
and temperature variations
less significant

Significant surface contribution between 19-85

Surface is usually opaque at 150 and 183 GHz – signals due to snow and drizzle



Tropical Rainfall

convective clouds with ice region above rain, warm ocean background

Higher Latitude Drizzle

Lower altitude clouds, mainly non-convective, cold ocean background, 24 marzo 2015

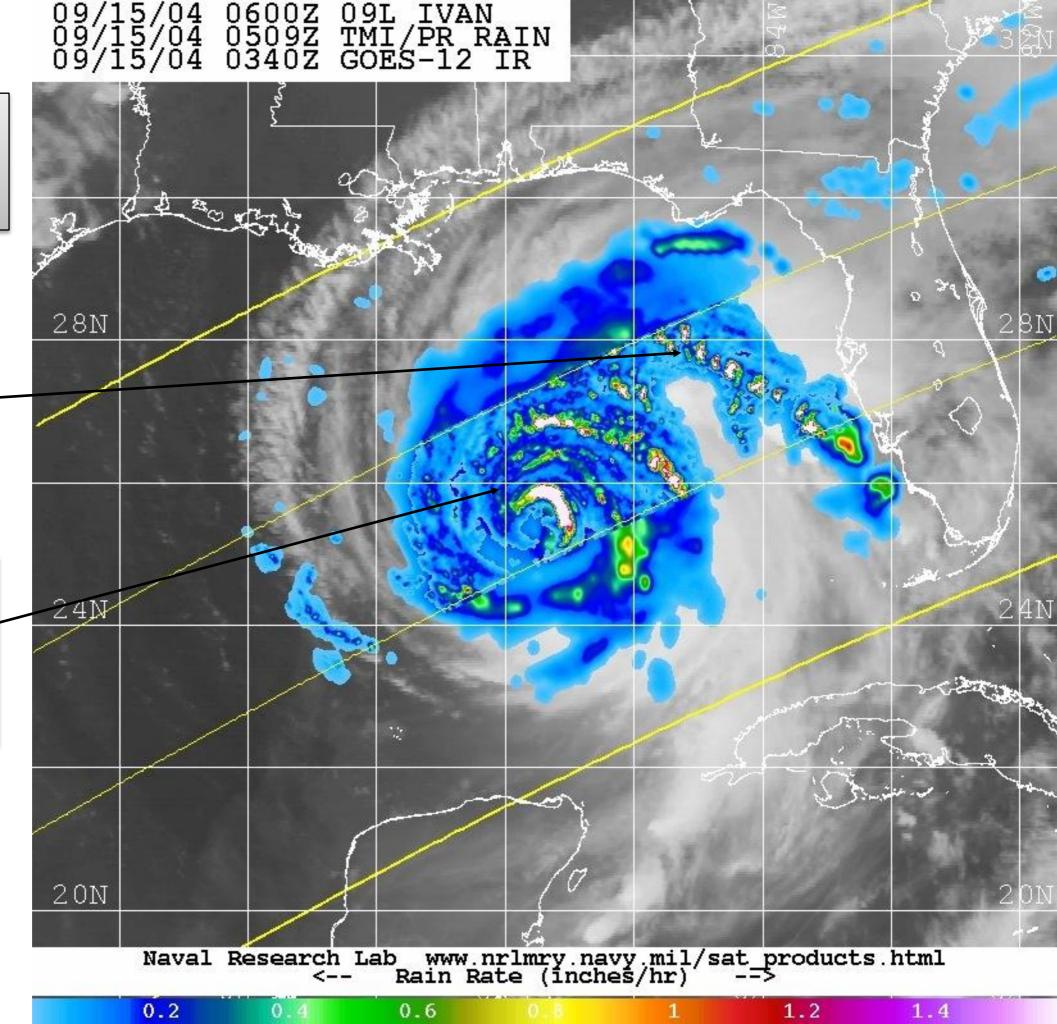
TRMM Satellite Sensors TRMM altitude was 350-km at launch, boosted **Flight Direction** to 402-km in 2001 to extend lifetime Lower altitude= finer spatial resolution **TRMM Microwave Imager (TMI)** 10-85 GHz **Precipitation Radar** (PR) 13.8 GHz 5-40 km **Microwave radiometer** Radar swath width=220km swath width=850km 5km

TRMM TMI/PR 15 Sep 2004 0509 UTC Over-Ocean

TMI can't delineate fine-scale structure

PR-estimated precip is displaced from the TMI-estimated precip due to parallax

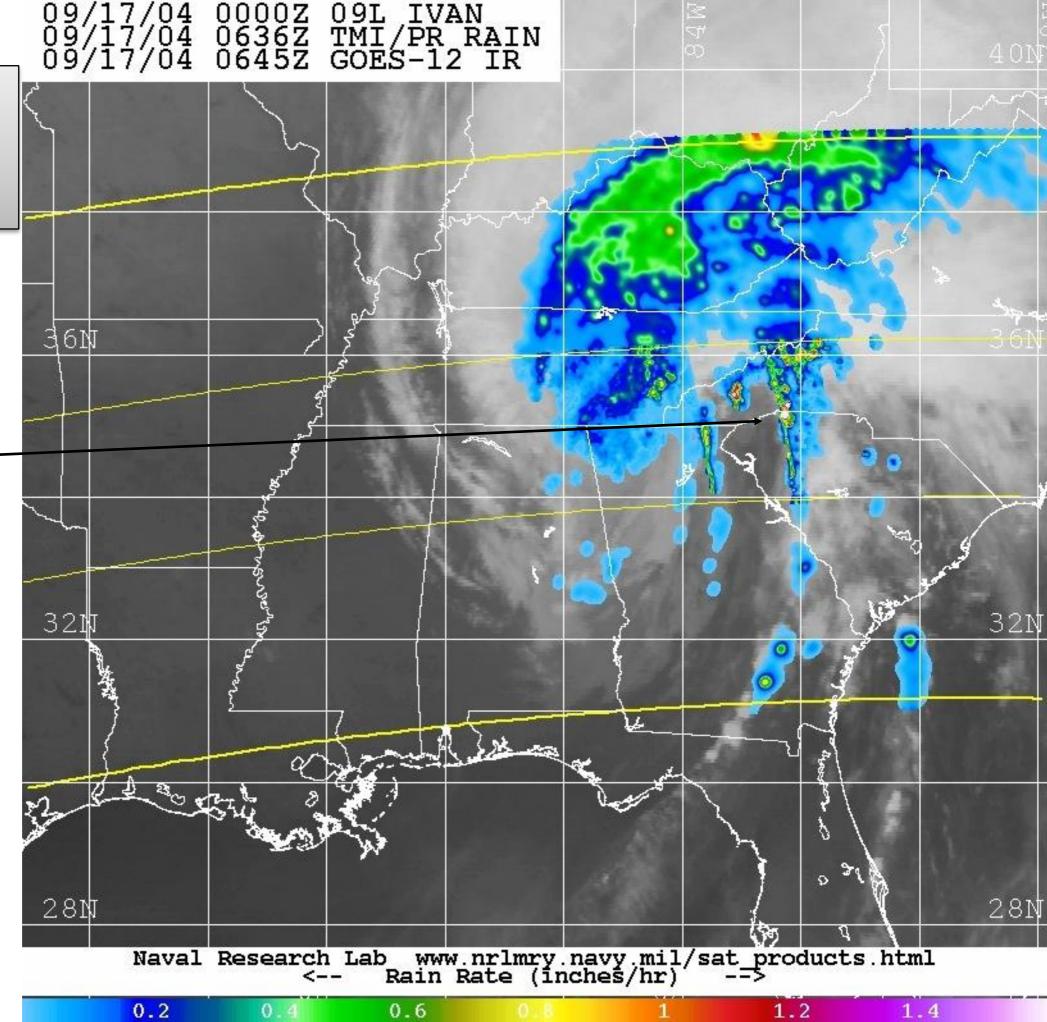
satellite motion

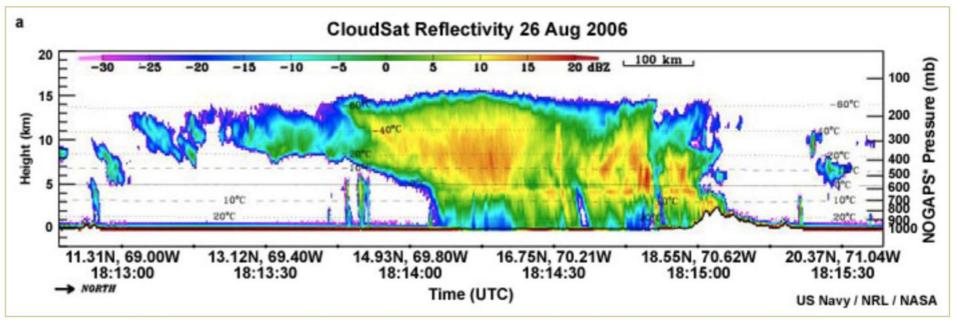


TRMM TMI/PR 17 Sep 2004 0636 UTC Over-Land

TMI can't capture
the heavy isolated
rain events (the tail
of the rainfall
histogram, i.e. the
"few big events")
that the PR picks
up

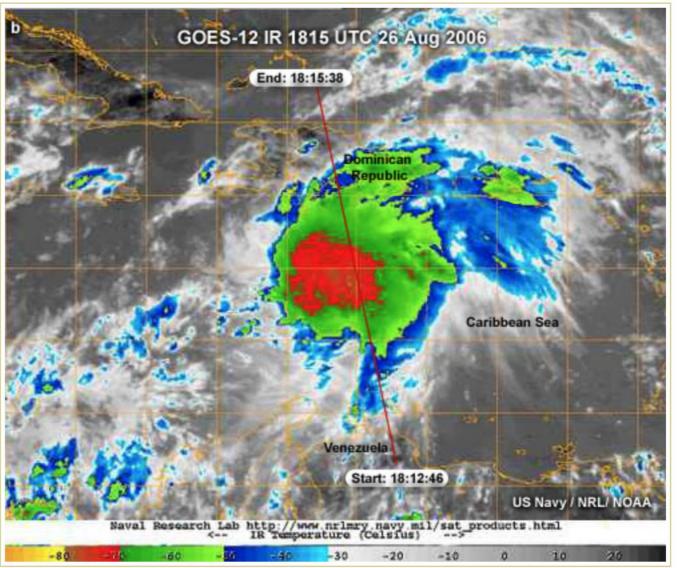
satellite motion





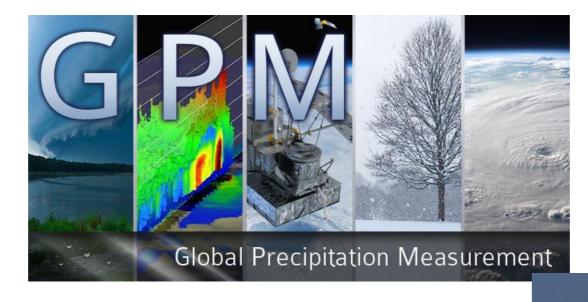
CloudSat profile through Tropical Storm Ernesto.

A broad area of high reflectivity extends south of the mountains of the Dominican Republic. Red and orange areas indicate the presence of large amounts of cloud water and/or ice while blue areas above indicate cloud ice. Wavy blue lines along the bottom of the cloud mass indicate intense rainfall. The top-down satellite-IR view misses two small thunderstorms beneath the cirrus anvil. The role of orographic lift in producing large amounts of cloud water is indicated by high reflectivity along the mountain peaks. At the same altitude over the ocean, reflectivity values are mostly lower.



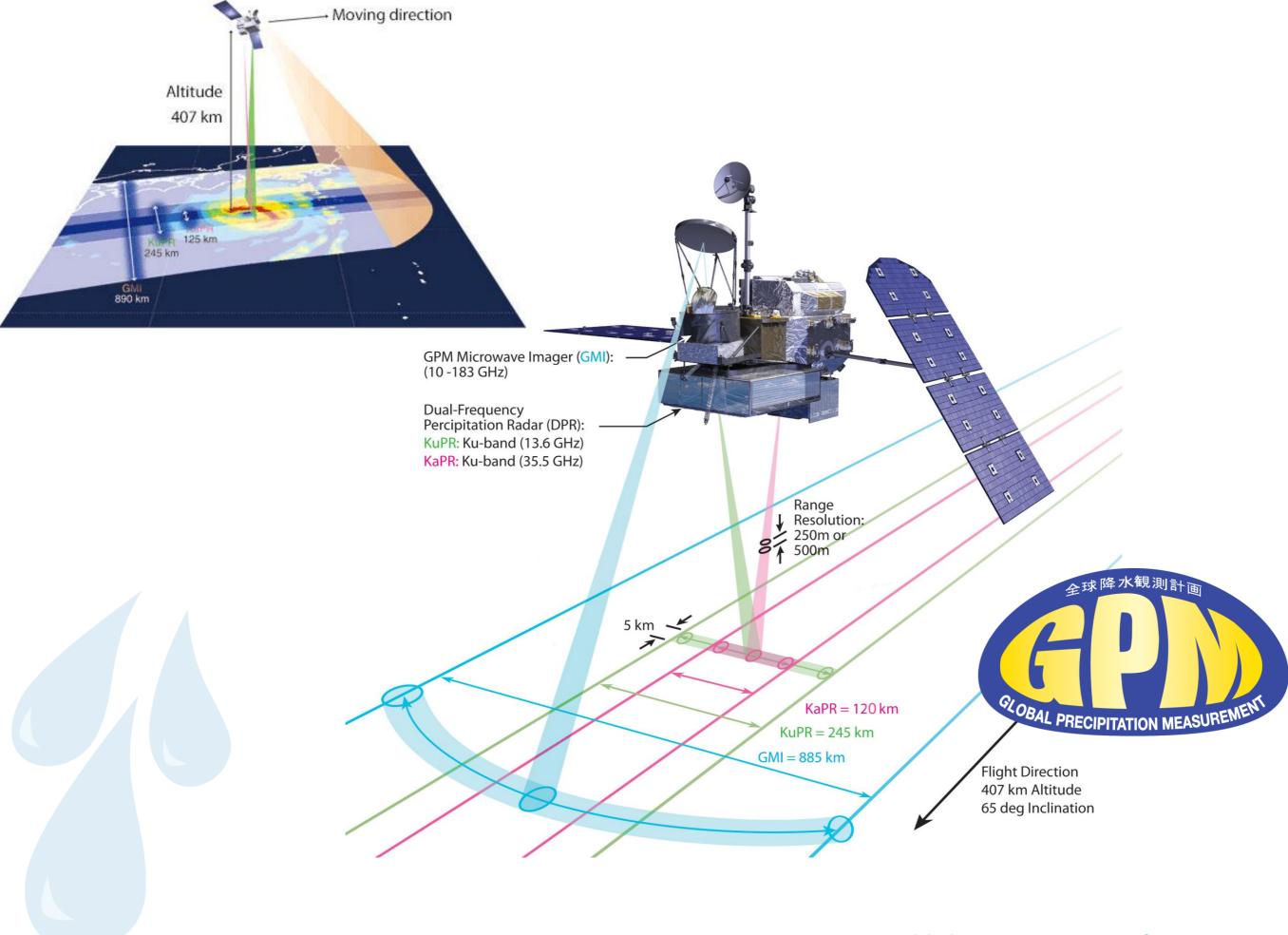
The Global Precipitation Measurement

GPM mission

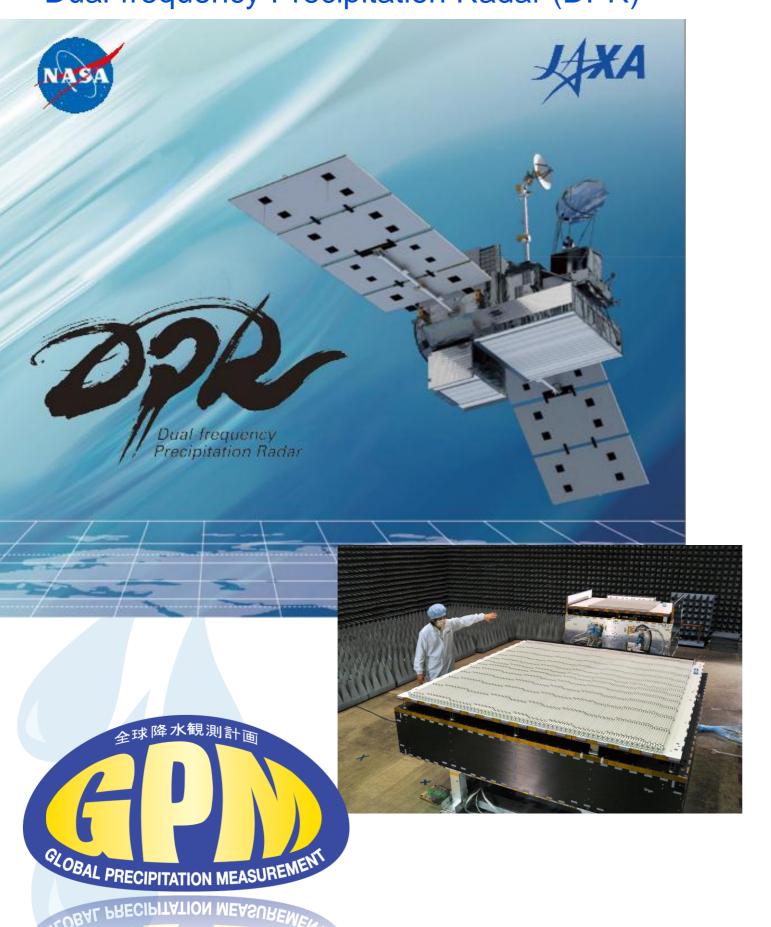




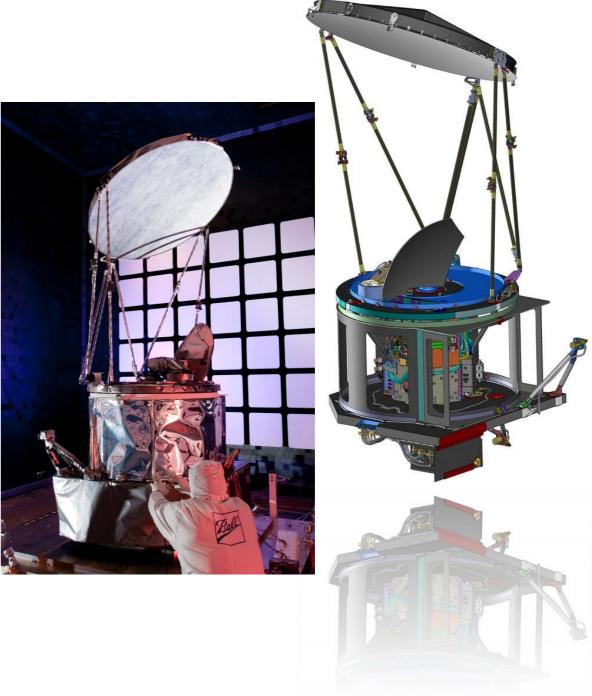


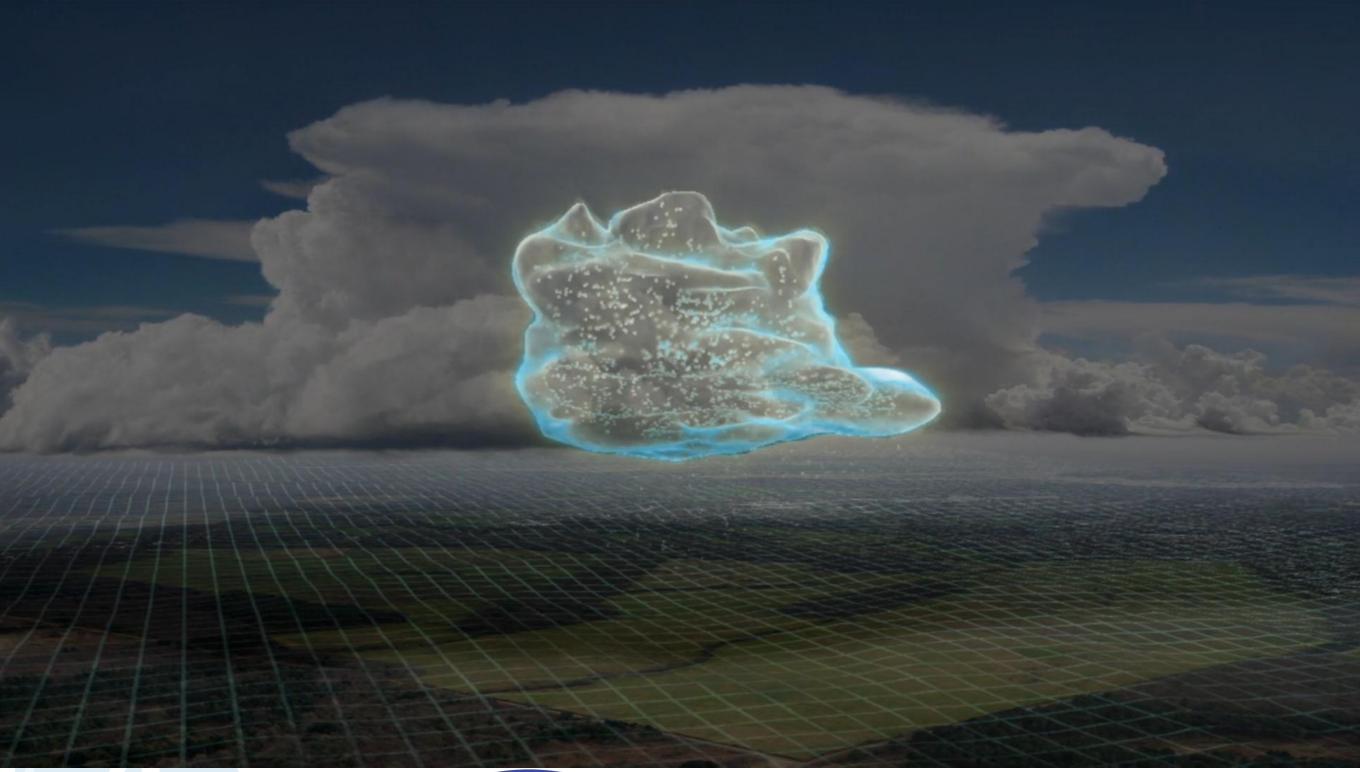


Dual-frequency Precipitation Radar (DPR)



GPM Microwave Imager (GMI)

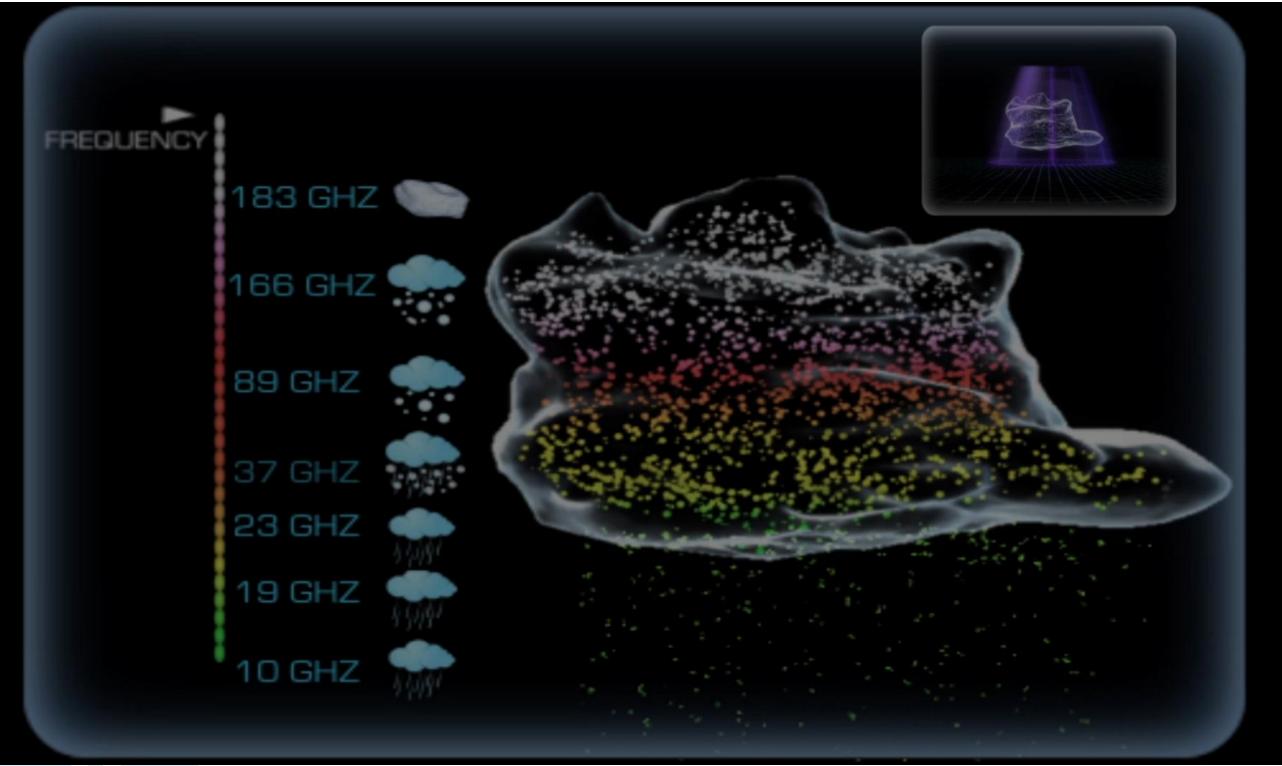








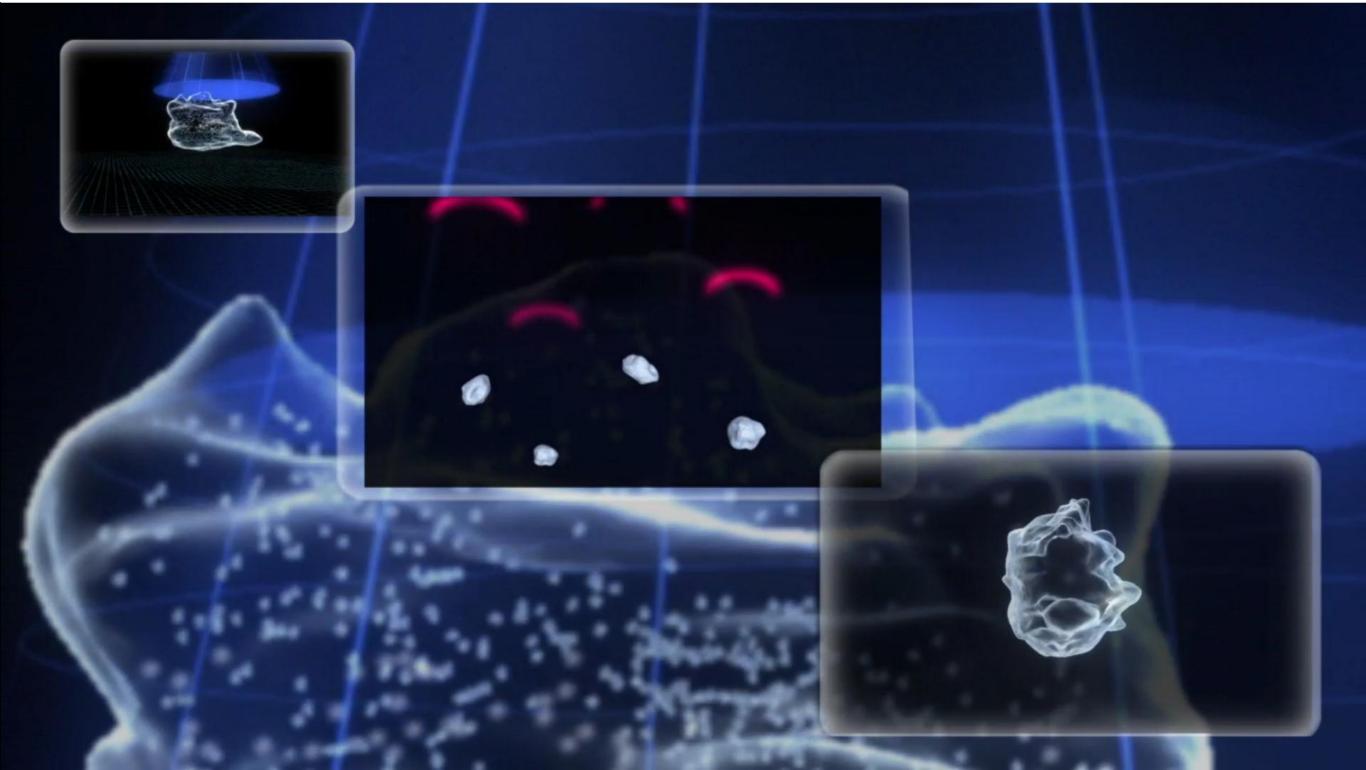












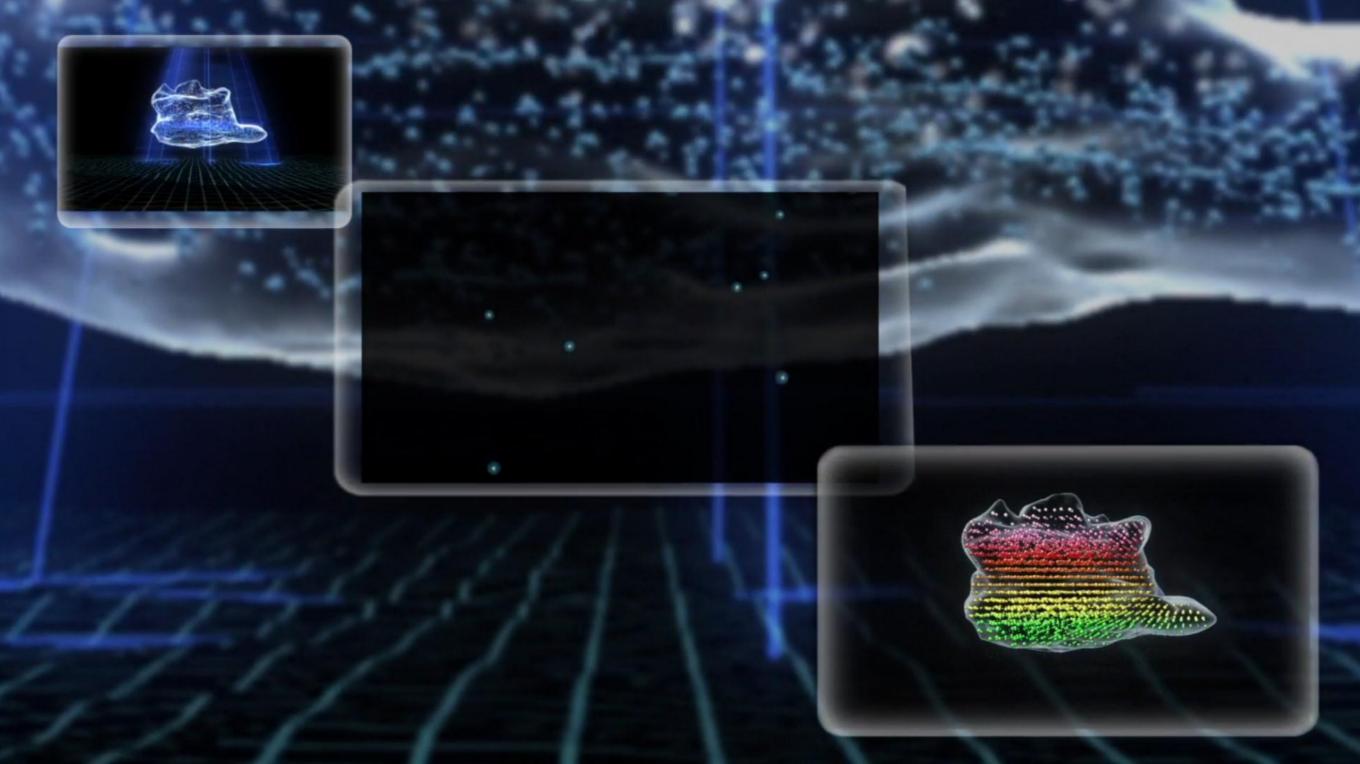








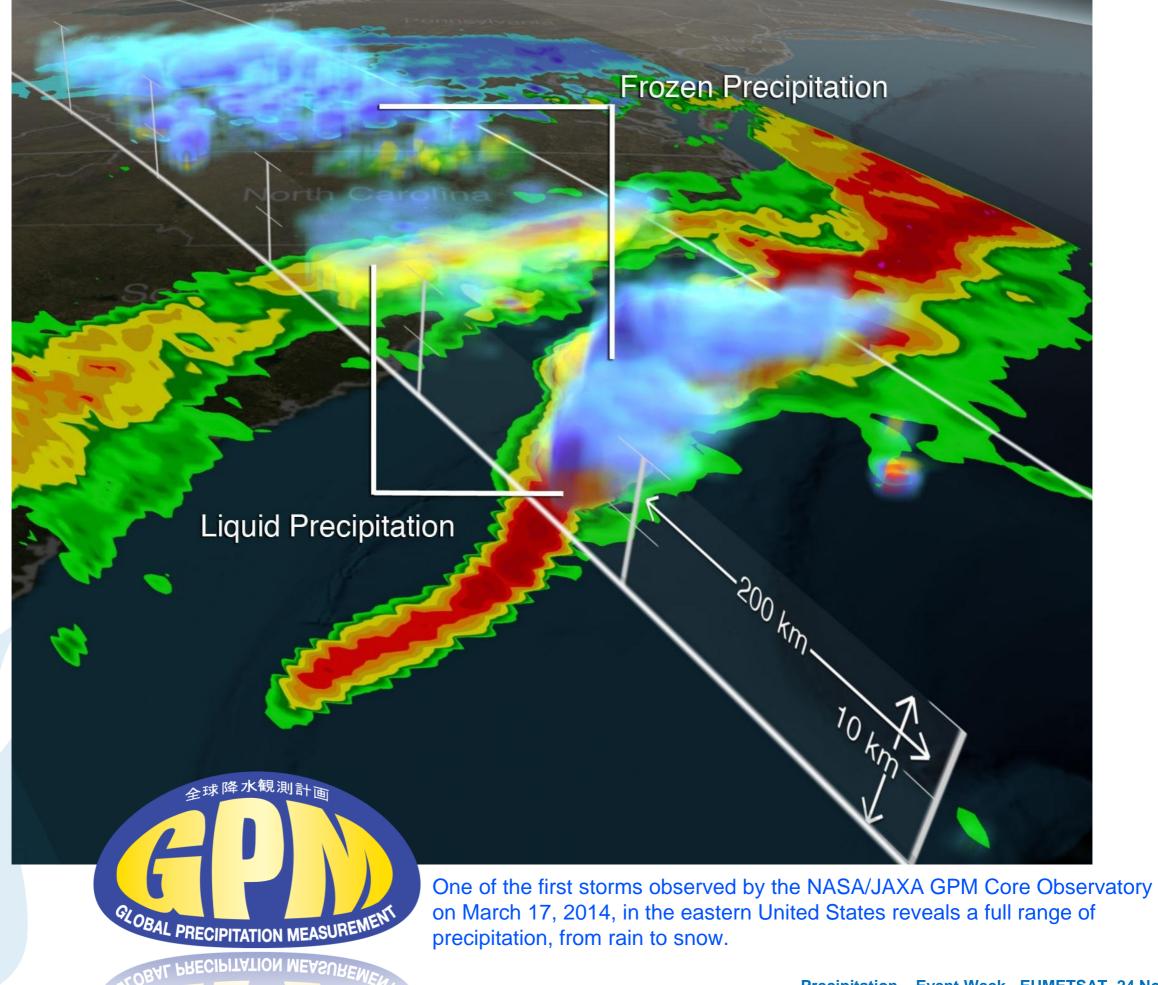


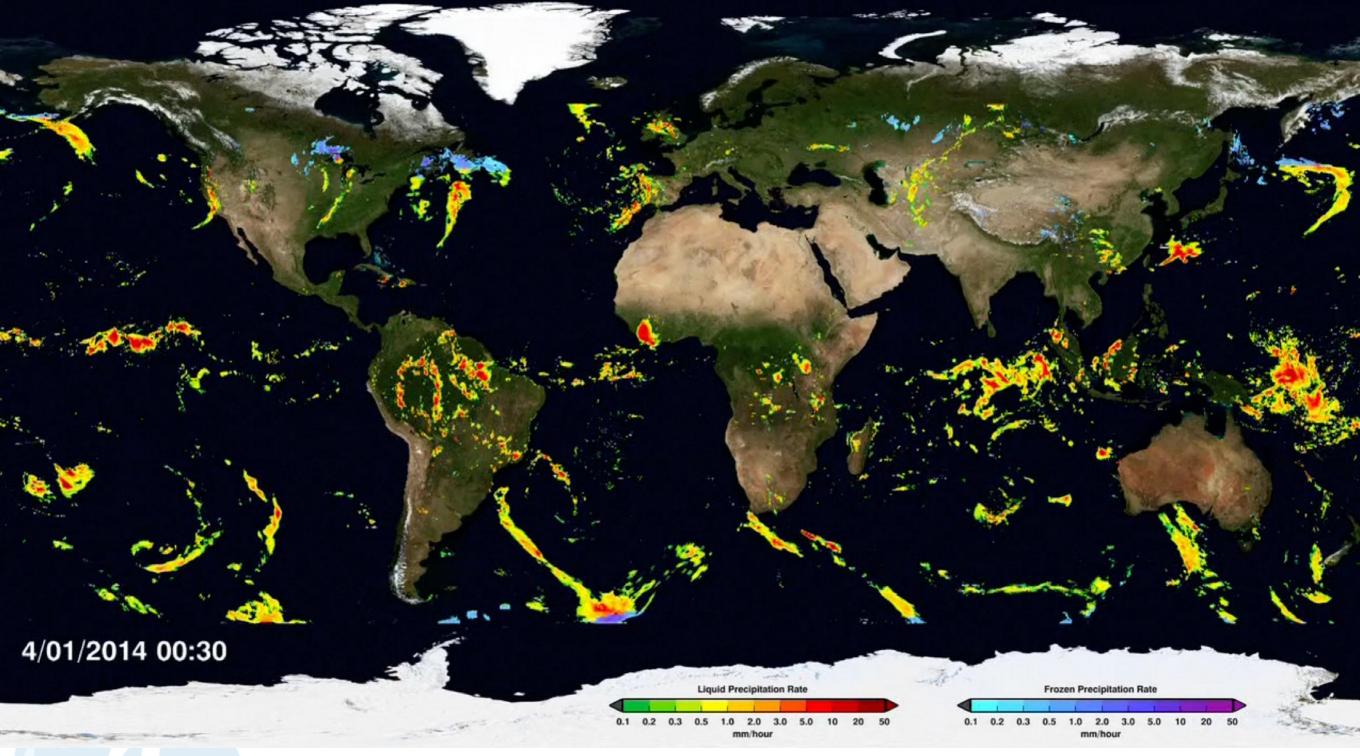




GPM constellation





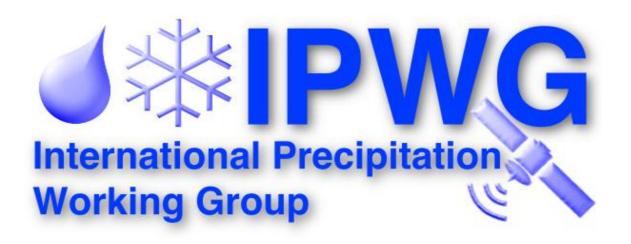


Global rainfall (yellow-red) and snowfall (blue) map from GPM, April-September 2014





A one stop shop



http://ipwg.isac.cnr.it/



That's all folks!



Thank you