

High Impact Weather Event Week

A recipe for Thunderstorms

For EUMETRAIN

Speaker: Helge Tuschy

estofex

<http://www.estofex.org/>





- studied Meteorology / Geophysics at the Leopold-Franzens University of Innsbruck, Austria
- 2002, 2007 internships National Weather Service Amarillo, Texas
- 2004 and 2011 internships Storm Prediction Center (SPC)
- joined Hazardous Weather Testbeds in 2012 and 2013
- talks: ECSS, ICAM, media ...
- forecaster at ESTOFEX / member of ESSL
- Since 2010 forecaster at DWD

The outline

- Ingredients-based forecasting method
- Hodograph
- Excessive rainfall with deep moist convection
- ESTOFEX – a very short overview

Basic ingredients for thunderstorms

Deutscher Wetterdienst
Wetter und Klima aus einer Hand



Lift : - synoptic / mesoscale
- orography

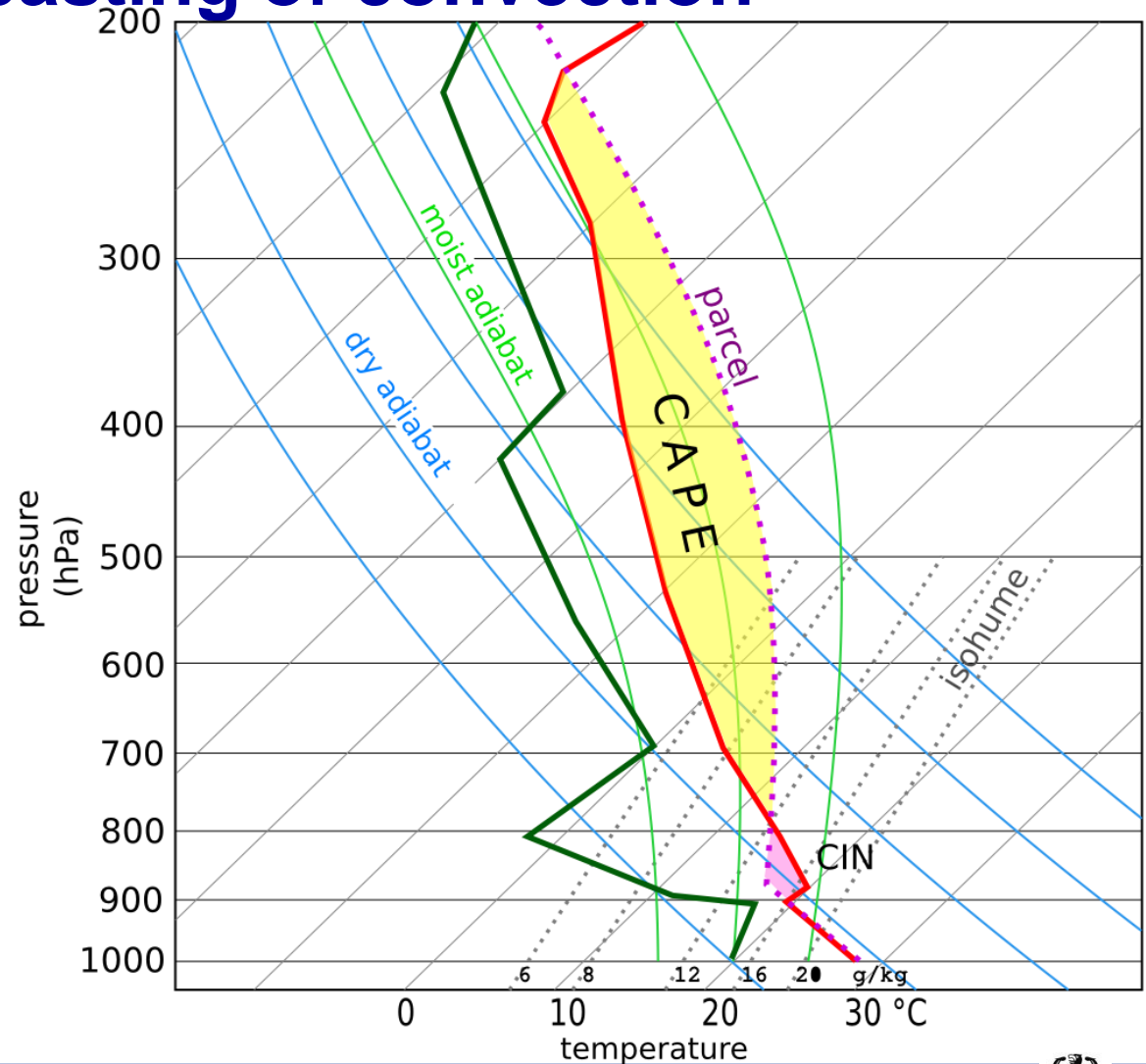
Instability :
- Convective instability

Moisture: - mid-troposphere
- boundary layer



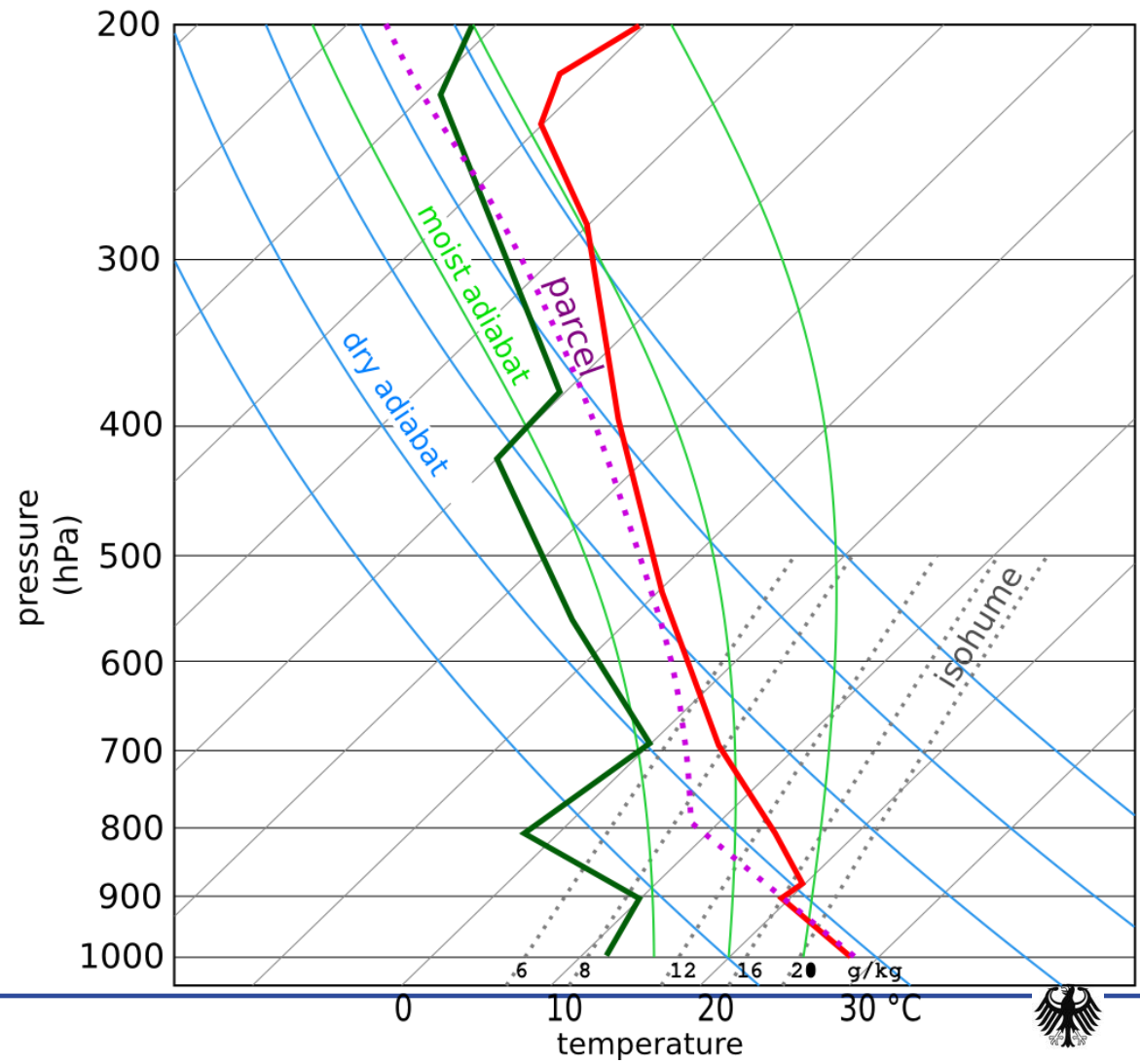
Scientific forecasting of convection

- A buoyant parcel means that the parcel has CAPE
- CAPE requires:
 1. steep lapse rates
 2. low-level moisture



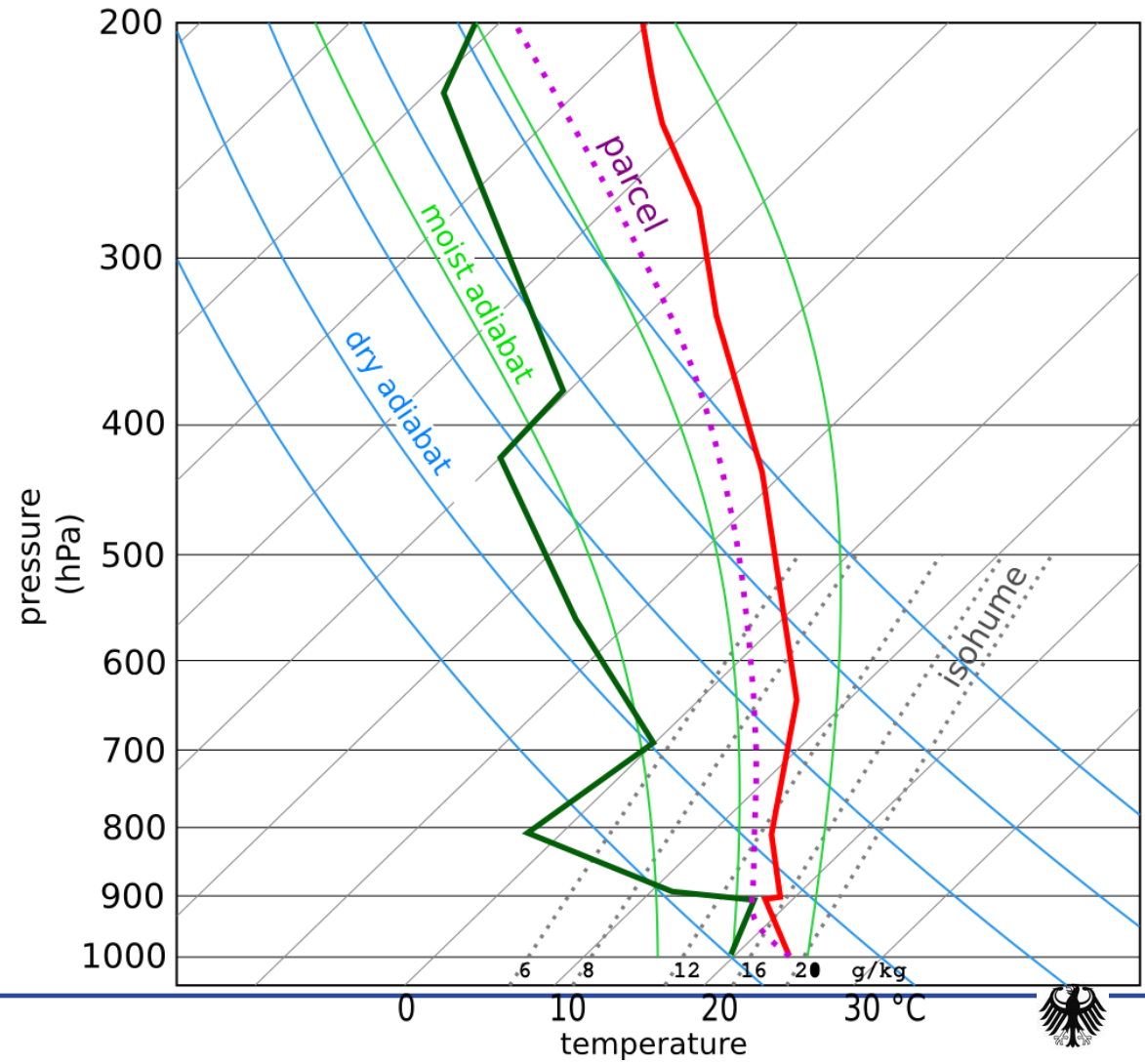
Scientific forecasting of convection

No moisture:
no CAPE



Scientific forecasting of convection

No lapse rates:
no CAPE either



Scientific forecasting of convection

In the atmosphere, low-level moisture and lapse rates evolve relatively independently:

One can track **low-level moisture**

- travels with *low-level winds*
- increases through evaporation
- decreases by condensation or mixing with drier air.

One can track **lapse rates**:

- conserved in geostrophic flow, travelling with *low to mid-level winds*
- modified slowly(!) by large-scale ascent and subsidence
- and rapidly by diabatic processes (heating, radiational cooling)

In contrast: **CAPE**, in model output just suddenly pop up.

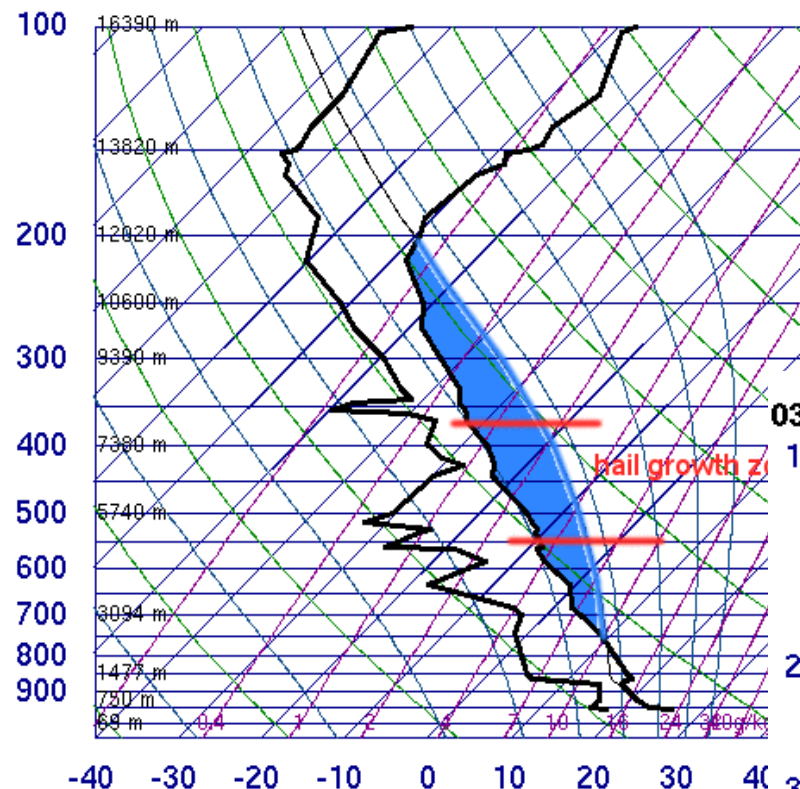
Convective forecast parameters

CAPE



MUCAPE:

Total amount of potential energy available to the most unstable parcel of air found within the lowest 300-mb of the atmosphere.



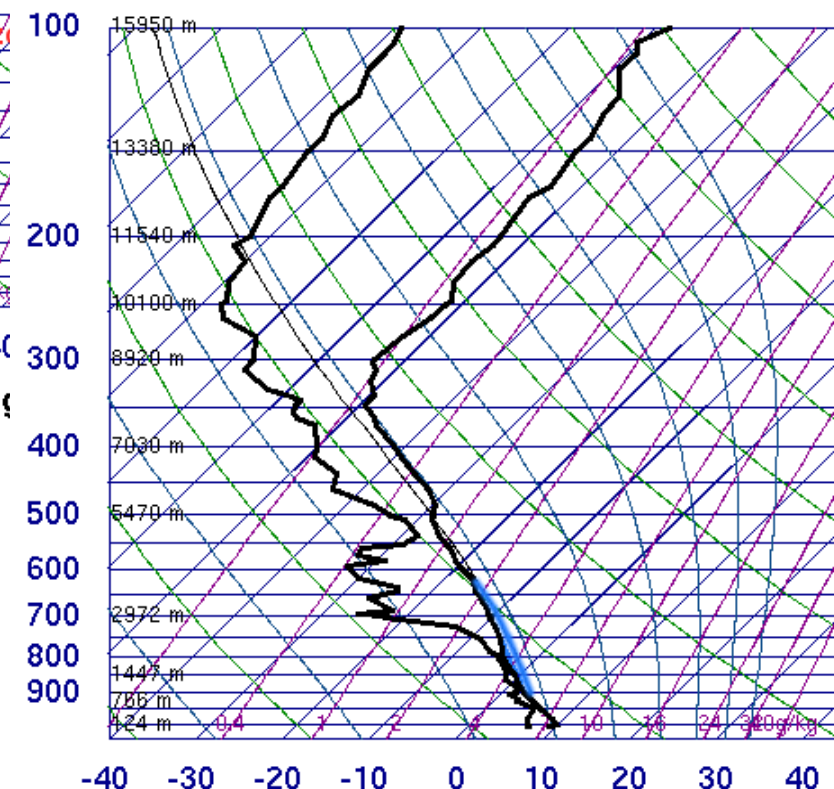
12Z 30 May 2008

University of Wyoming

SLAT 48.25
 SLON 11.55
 SELV 489.0
 SHOW 0.68
 LIFT -6.48
 LFTV -7.07
 SWET 123.6
 KINX 28.50
 CTOT 18.90
 VTOT 30.90
 TOTL 49.80
 CAPE 2177.
 CAPV 2221

LLCAPE vs. mid-level CAPE

03808 Camborne

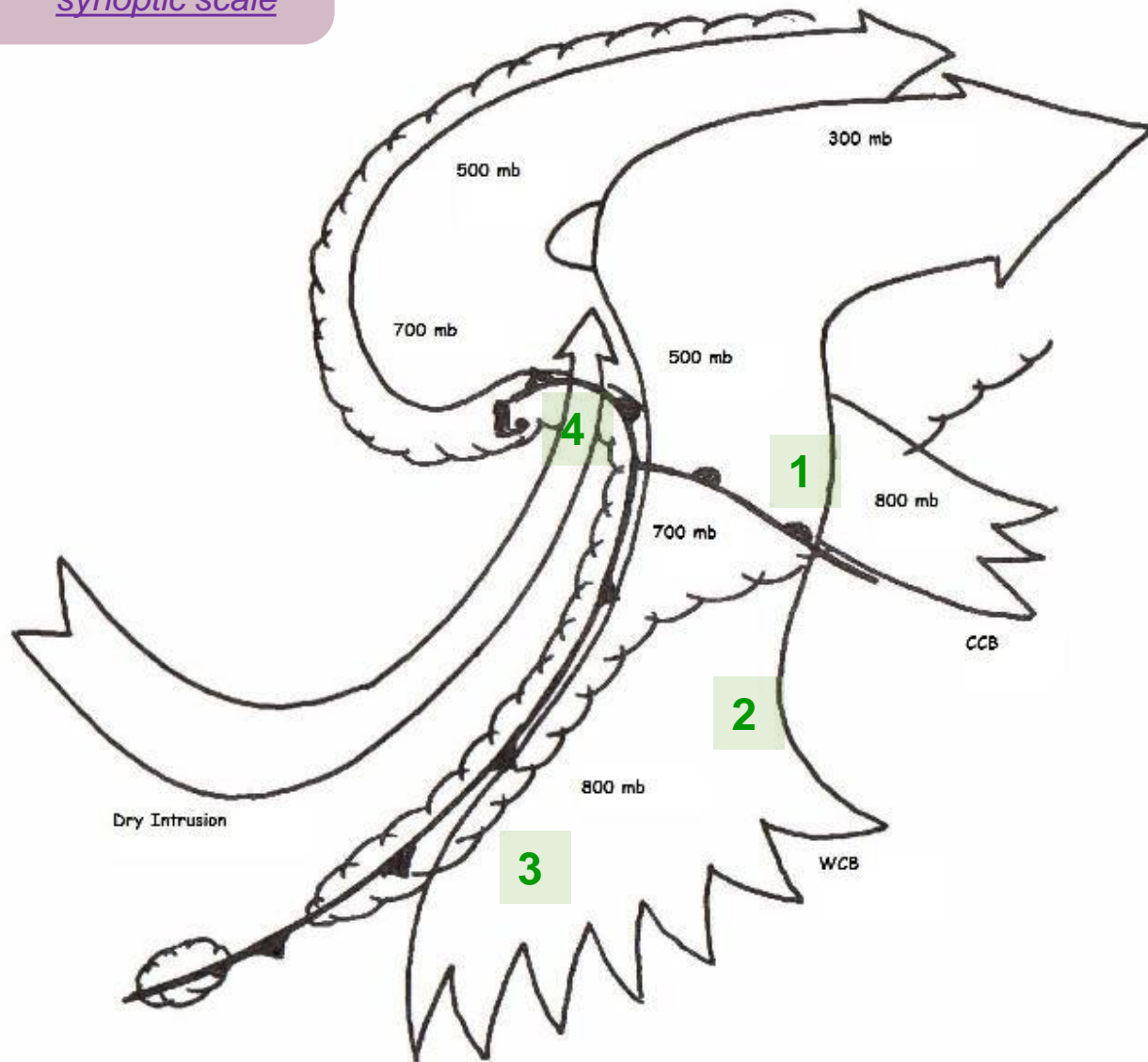


00Z 12 Nov 2008

University of Wyoming

SLAT 50.22
 SLON -5.32
 SELV 88.00
 SHOW 2.30
 LIFT 0.09
 LFTV 0.05
 SWET 189.0
 KINX 14.40
 CTOT 27.10
 VTOT 28.20
 TOTL 55.30
 CAPE 155.1
 CAPV 179.2
 CINS -0.64
 CINV -0.36
 EQLV 502.9
 EQTV 502.5
 LFCT 928.0
 LFCV 933.0
 BRCH 120.1
 BRCV 138.8
 LCLT 278.4
 LCLP 942.7
 MLTH 283.2
 MLMR 5.97
 THCK 5346.
 PWAT 14.03

Lift
synoptic scale



1 Elevated thunderstorms

2 Warm sector convection

3 Cold front

4 Potential instability build-up along cloud head

Carlson (1980)



Elevated instability (e.g. north of warm front)

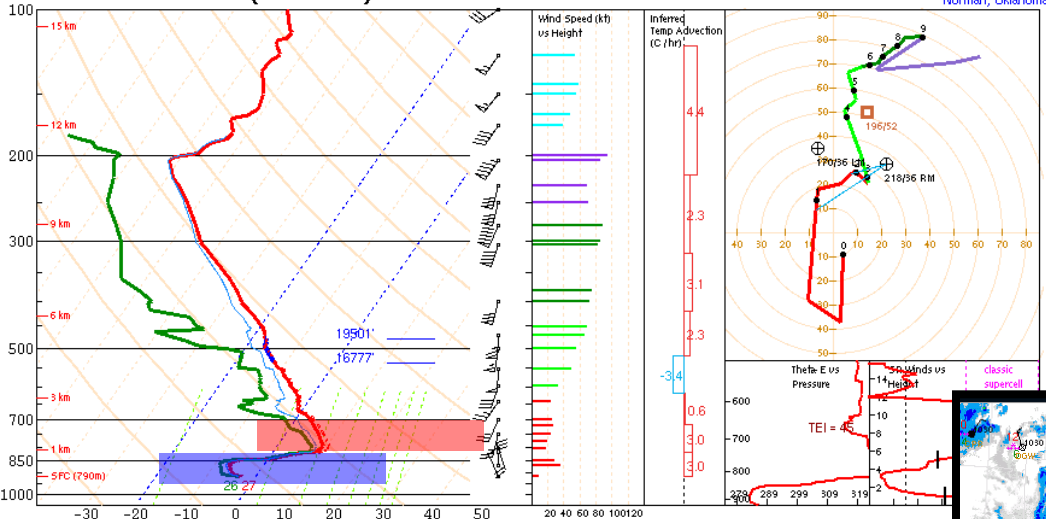
Lift
synoptic scale

Deutscher Wetterdienst
Wetter und Klima aus einer Hand



1

DDC 130409/1800 (Observed)



PARCEL	CAPE	CINH	LCL	LI	LFC	EL	SRH(m2/s2)	Shear(kt)	MnWind	SRW	
SURFACE	14	0	37m	24	37m	1761'	SFC - 1 km	462	25	6/18	27/52
MIXED LAYER	0	0	748m	18	M	2452'	SFC - 3 km	483	33	218/6	38/30
FCST SURFACE	0	0	2574m	7	M	8442'	Eff Inflow Layer	64	21	183/20	68/23
MU (796 mb)	504	-23	1437m	-3	2262m	31781'	SFC - 6 km	79	197/18	56/20	
PW = 0.67 in	3CAPE = 0 J/kg	WBZ = 0'	WINDG = 0.0				SFC - 8 km	89	196/25	75/15	
K = -2	DCAPE = -157 J/kg	FZL = 0'	ESP = 0.0				Lower Half Storm Depth	59	194/31	95/15	
MidRH = 60%	DownT = 53 F	ConvT = M	MMP = 0.98				Cloud Bearing Layer	49	196/46	149/18	
LowRH = 89%	MeanW = 3.5 g/kg	MaxT = 66F					BRN Shear = 229 m/s²				
SigSevere = 0 m3/s3							4-6km SR Wind = 156/33 kt				
Sfc-3km Agl Lapse Rate = -0.2 C/km						Storm Motion Vectors.....				
3-6km Agl Lapse Rate = 7.1 C/km							Bunkers Right = 218/36 kt				
850-500mb Lapse Rate = 2.5 C/km							Bunkers Left = 170/36 kt				
700-500mb Lapse Rate = 7.5 C/km							Corfidi Downshear = 194/94 kt				
							Corfidi Upshear = 193/51 kt				


Supercell = 0.6

Left Supercell = -0.7

Sig Tor (CIN) = 0.0

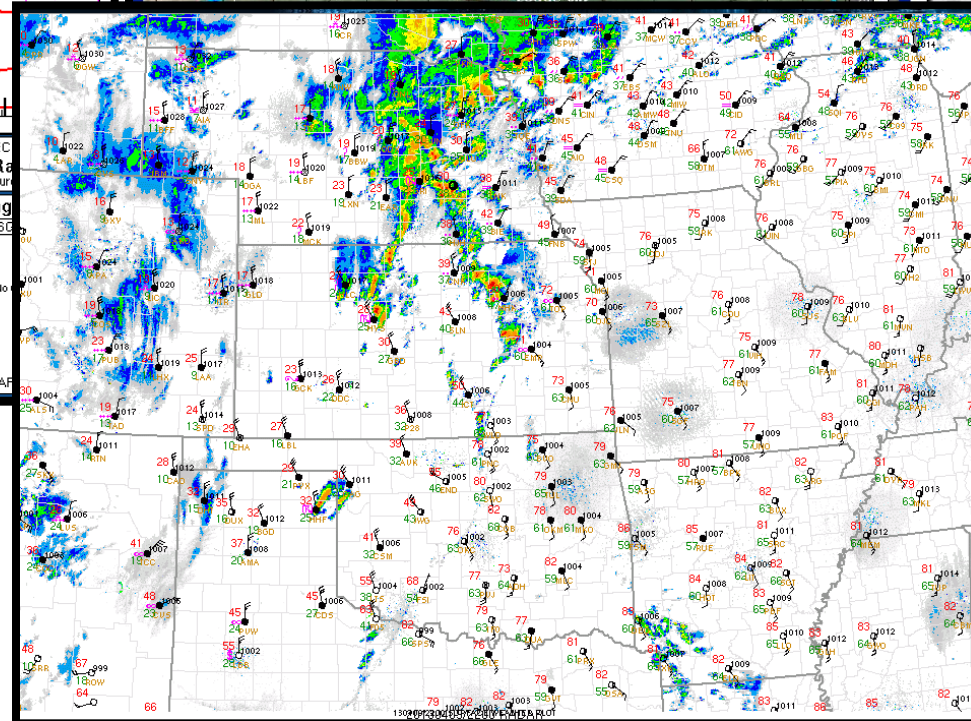
Sig Tor (fixed) = 0.0

Sig Hail = 0.0



1km & 6km AGL Wind Barbs

*** BEST GUESS PRECIP
Based on sfc temperature
Freezing Rain
SARS - Sounding
SUPERCELL
No Quality Matches



Quelle:
<http://www.spc.noaa.gov/>





Thunderstorm and Breezy
28°F
 -2°C

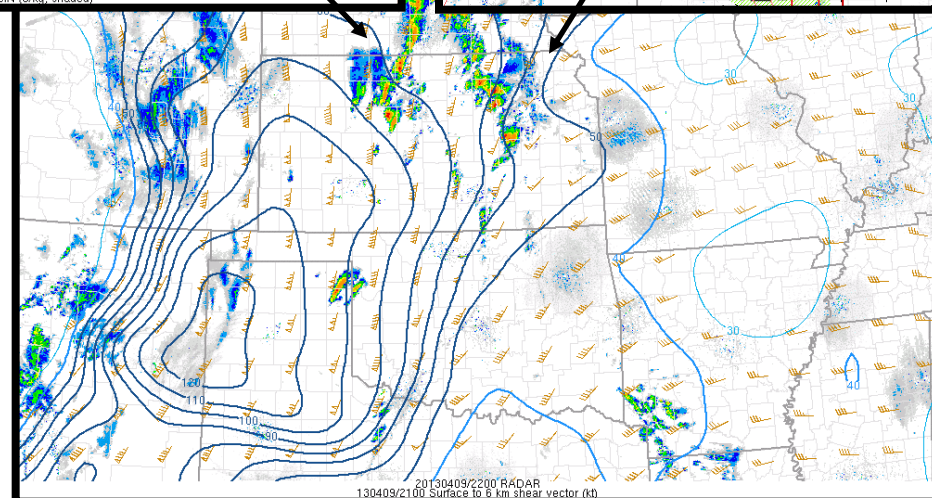
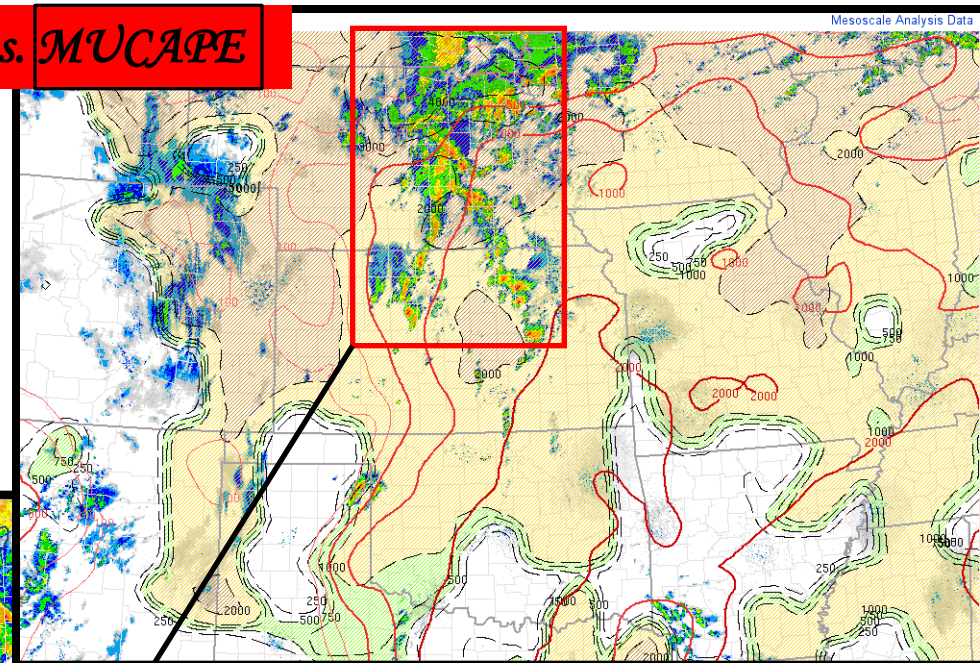
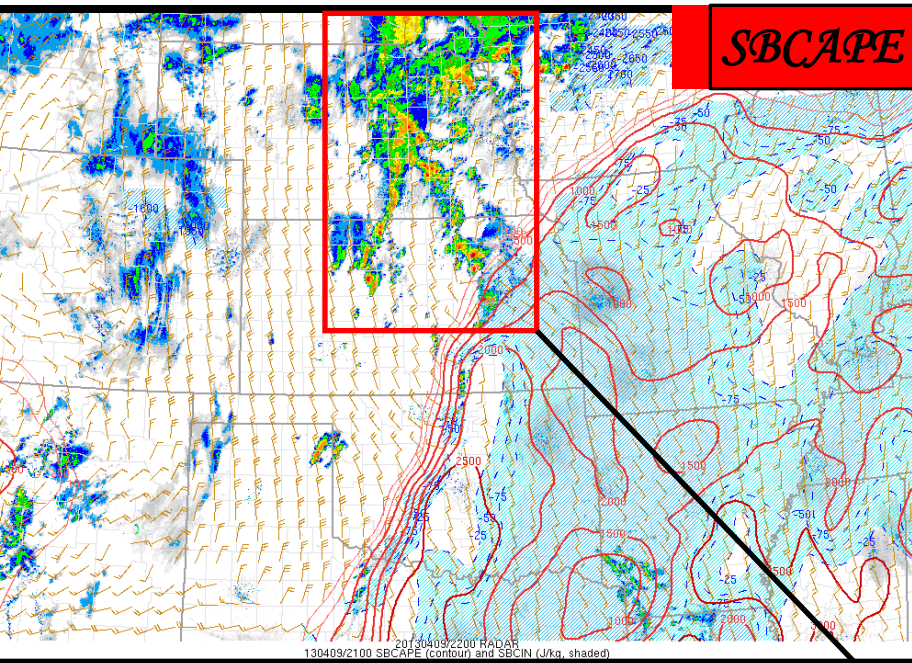
Humidity 86%
 Wind Speed NW 25 G 40 mph
 Barometer 29.88 in
 Dewpoint 25°F (-4°C)
 Visibility 3.00 mi
 Wind Chill 13°F (-11°C)
 Last Update on 9 Apr 4:55 pm CDT

Lift
synoptic scale

Deutscher Wetterdienst
Wetter und Klima aus einer Hand



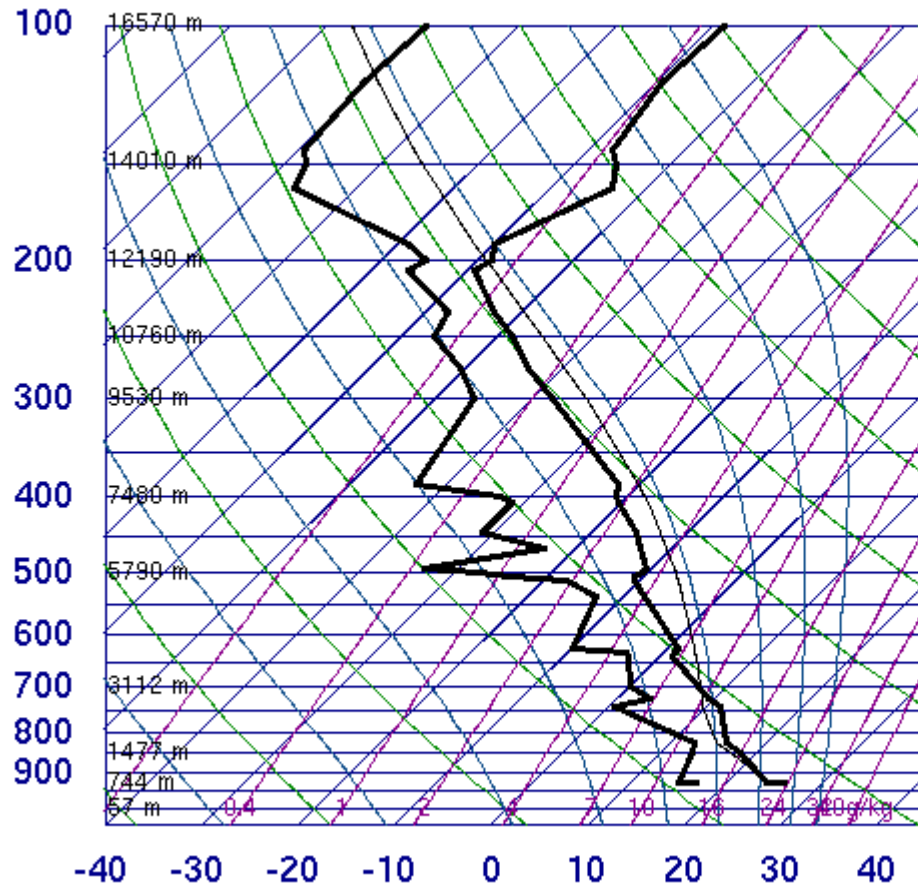
SBCAPE vs. MUCAPE



Convection in warm sector

Tornado outbreak in Poland, 15th August 2008

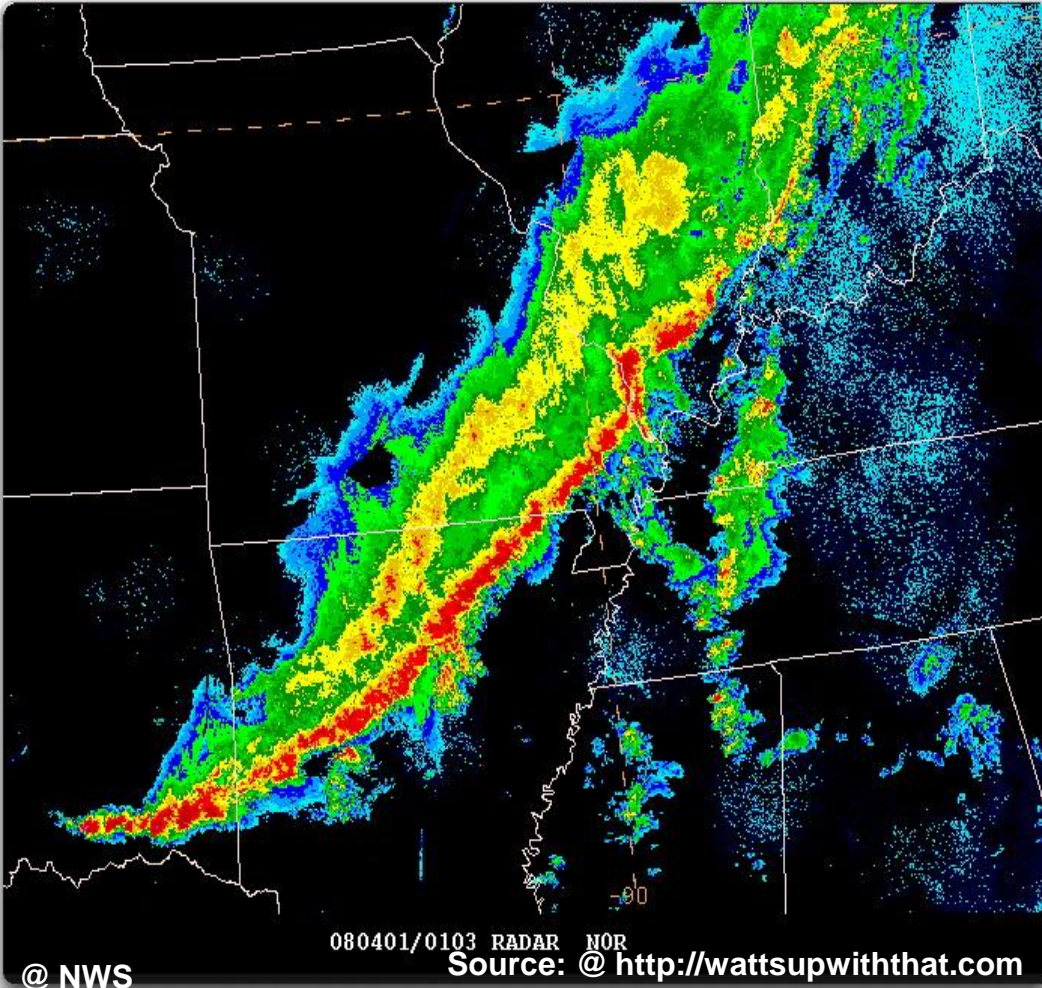
11952 Poprad-Ganovce



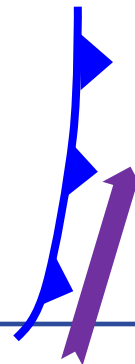
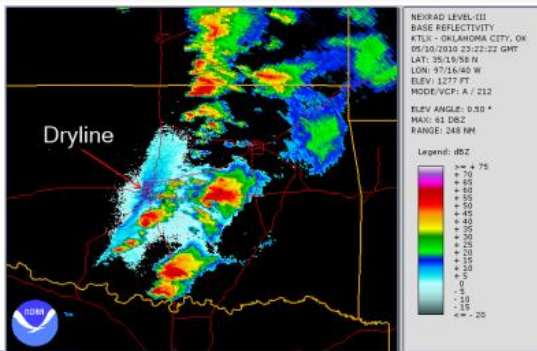
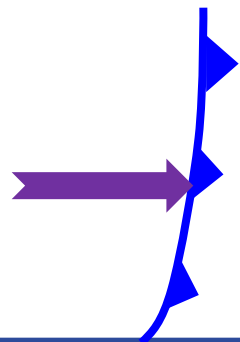
SLAT	49.03
SLOH	20.31
SELV	706.0
SHOW	-3.65
LIFT	-3.48
LFTV	-4.09
SWET	310.6
KINX	35.30
CTOT	23.90
VTOT	28.90
TOTL	52.80
CAPE	1030.
CAPV	1132.
CINS	-82.5
CINV	-58.2
EQLV	201.3
EQTV	201.3
LFCT	706.0
LFCV	719.7
BRCH	5.42
BRCV	5.95
LCLT	286.0
LCLP	803.7
MLTH	304.4
MLMR	11.74
THCK	5733.
PWAT	30.11

Break the cap:

- Mesoscale forcing** (e.g. convergent boundary layer flow)
- Increase of **low-level moisture** (advection, evapotranspiration)
- Diabatic heating** (cloud cover in warm sector)
- Synoptic lift** (short wave?)



- Strong cross-frontal circulation often results in widespread initiation
- Closer to best forcing of mid/upper-level trough
- Orientation of the shear vector is important for degree of line-up of thunderstorms:



e.g. 0-6 km shear vector

60-70 kt SW-erly 0-6 km shear vector

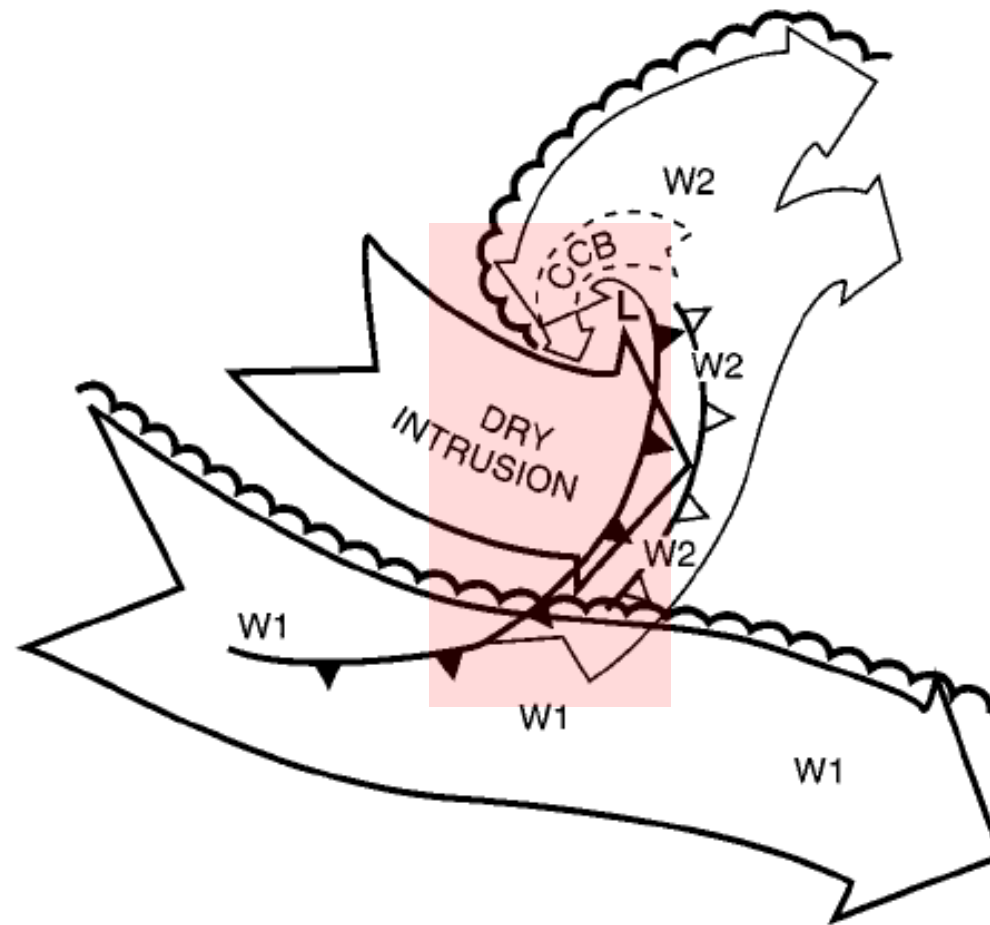
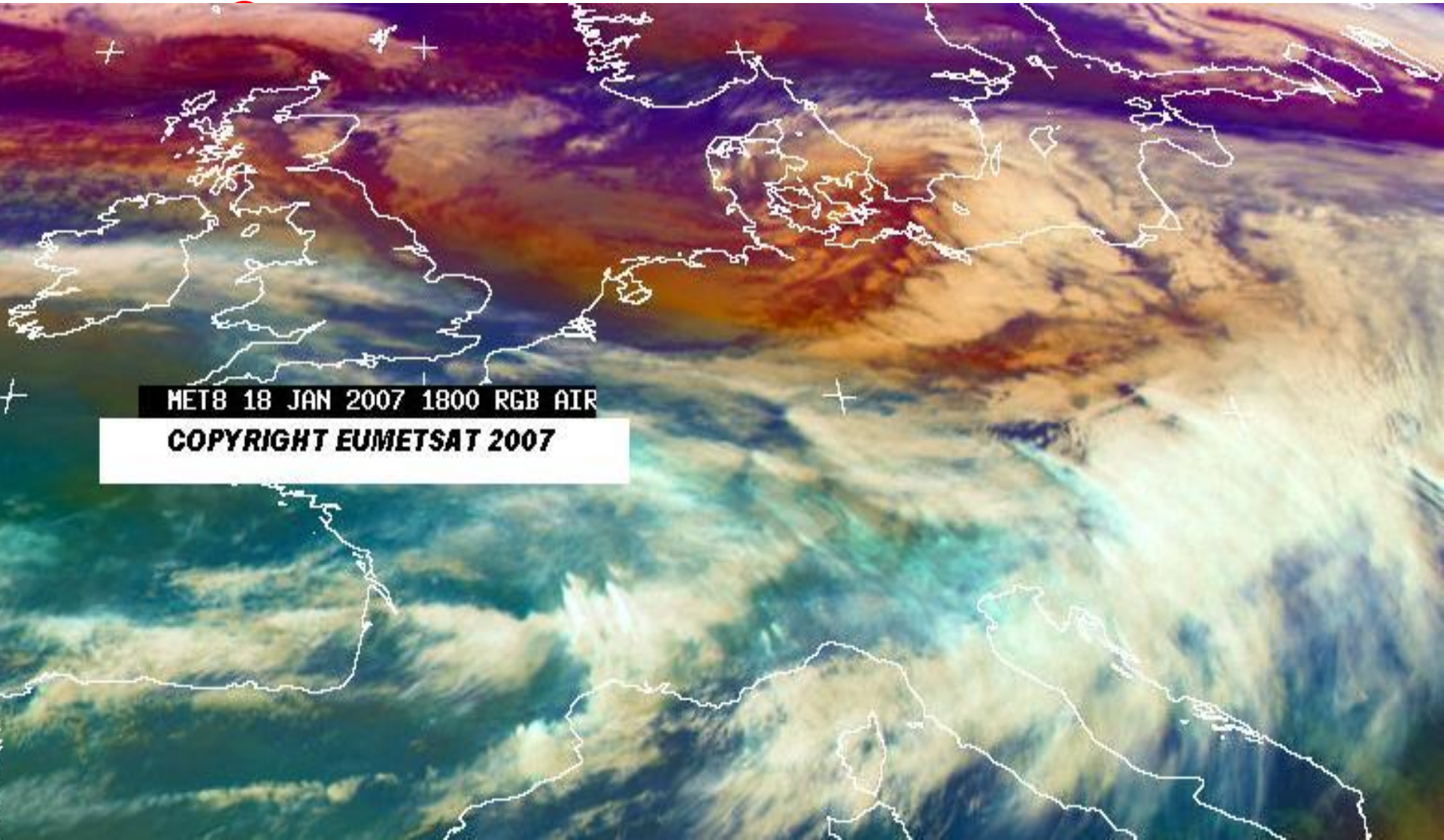


Figure 5. Conceptual model showing system-relative airflow associated with the diffluent-flow type of cyclogenesis. The arrows labelled W1 and W2 are the primary and secondary warm conveyor belts. The dashed arrow labelled CCB is a cold conveyor belt. The dry intrusion is seen to overrun W2 over a broad region to produce an upper cold front at its leading edge. (After Young, 1994.)

Kyrill 18/19.01.2007



MET8 18 JAN 2007 1800 RGB AIR

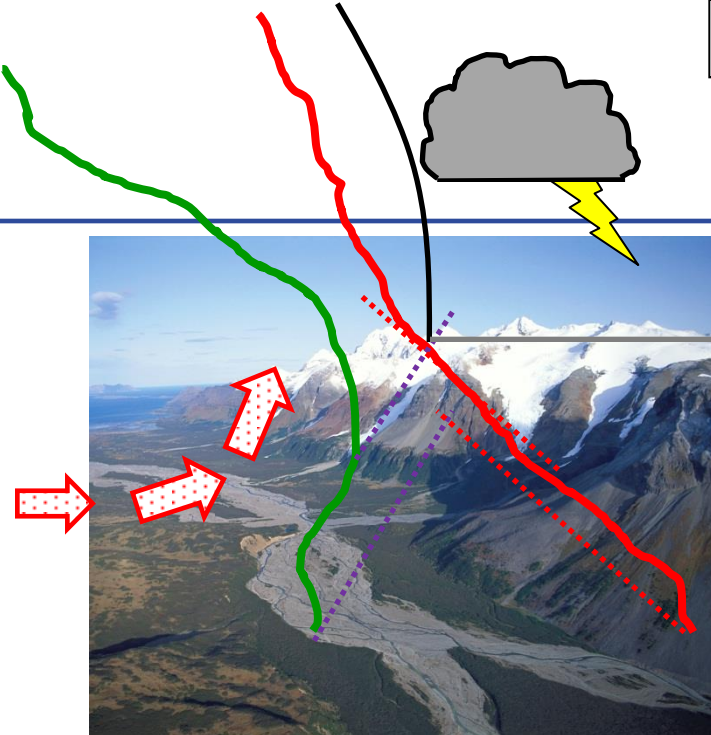
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Orographic lift

Lift

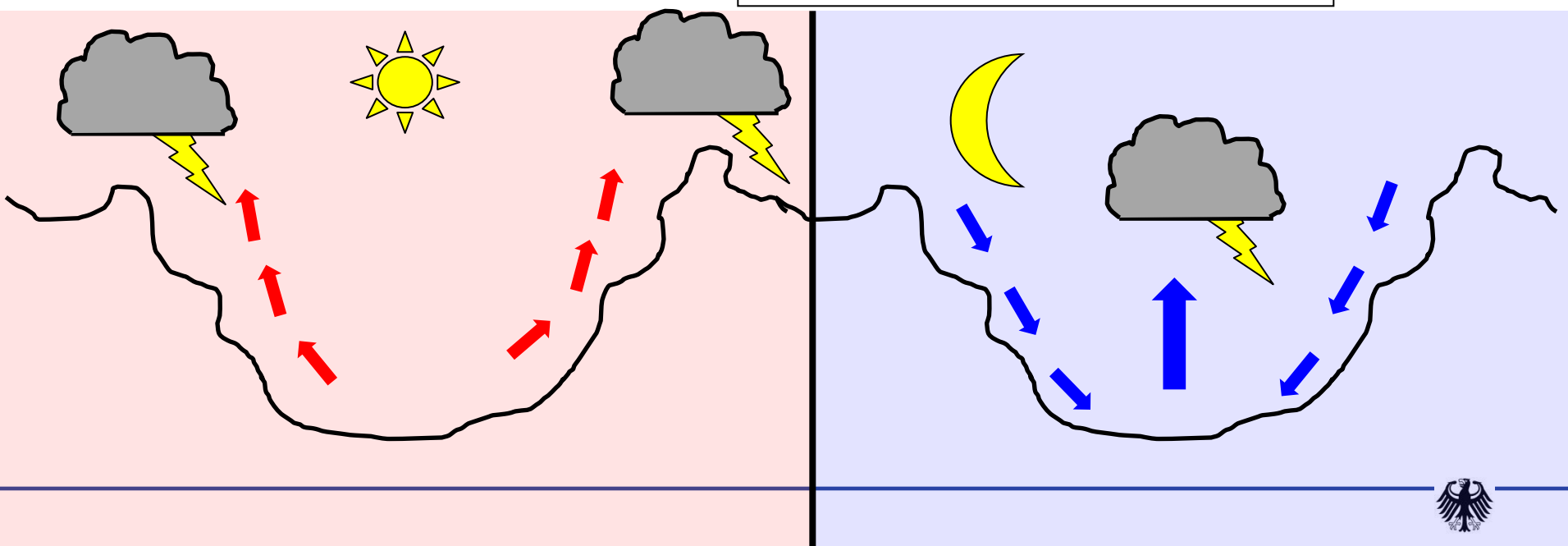
mesoscale

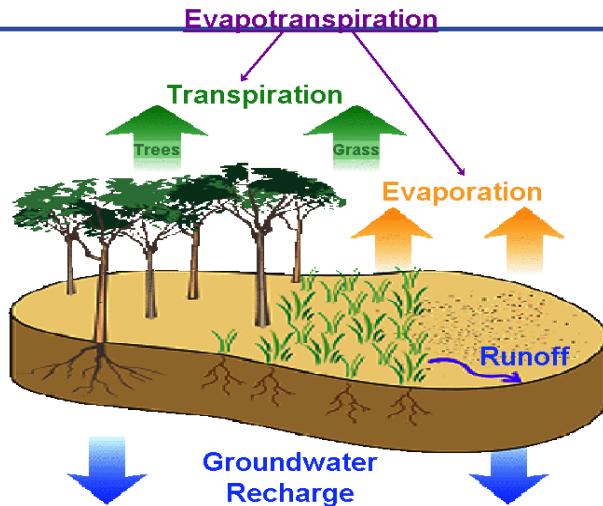
Deutscher Wetterdienst
Wetter und Klima aus einer Hand



Lifted Condensation Level (LCL)

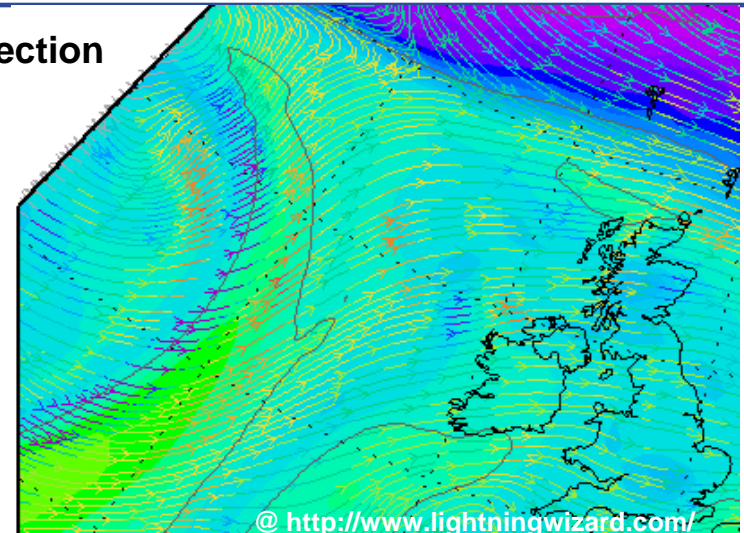
Mountain – Valley circulation





@ <http://weather.vouhead.gr/vouhead-images/evapotranspiration002-en.gif>

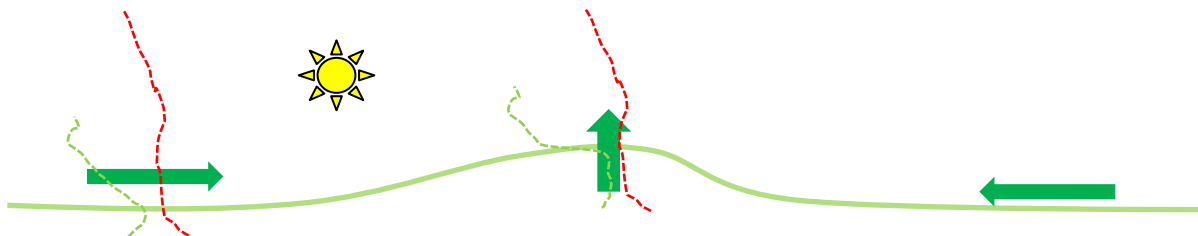
Advection



@ <http://www.lightningwizard.com/>

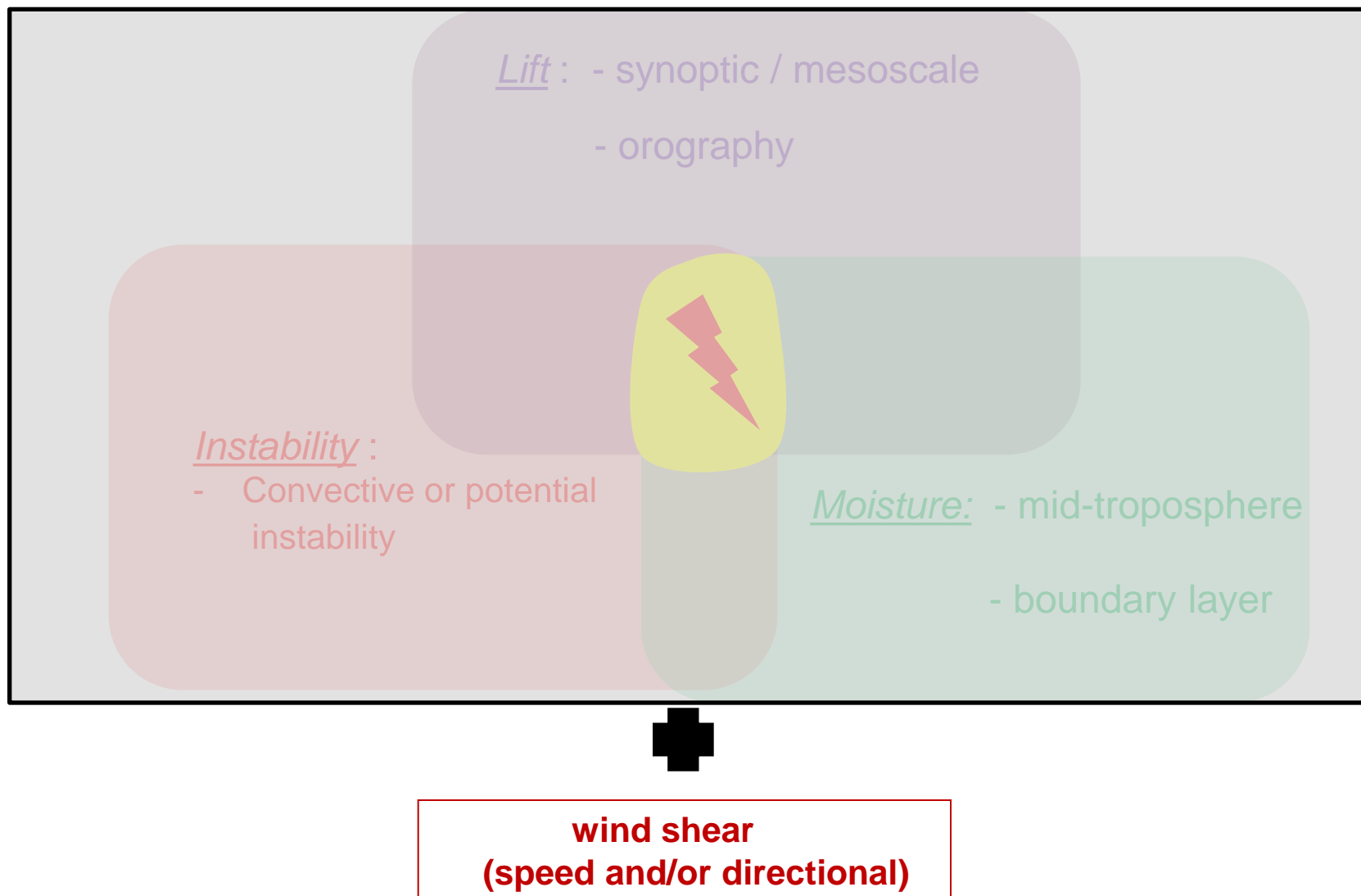
Moisture pooling along convergent „thermal“ boundary

- Moisture advection cannot create maxima, only translation of a maximum is possible
- Moisture convergence ~ velocity convergence ~ upward motion
- ➡ deepening of the boundary layer moisture ➡ less diurnal mixing



Basic ingredients for organized thunderstorms

Deutscher Wetterdienst
Wetter und Klima aus einer Hand

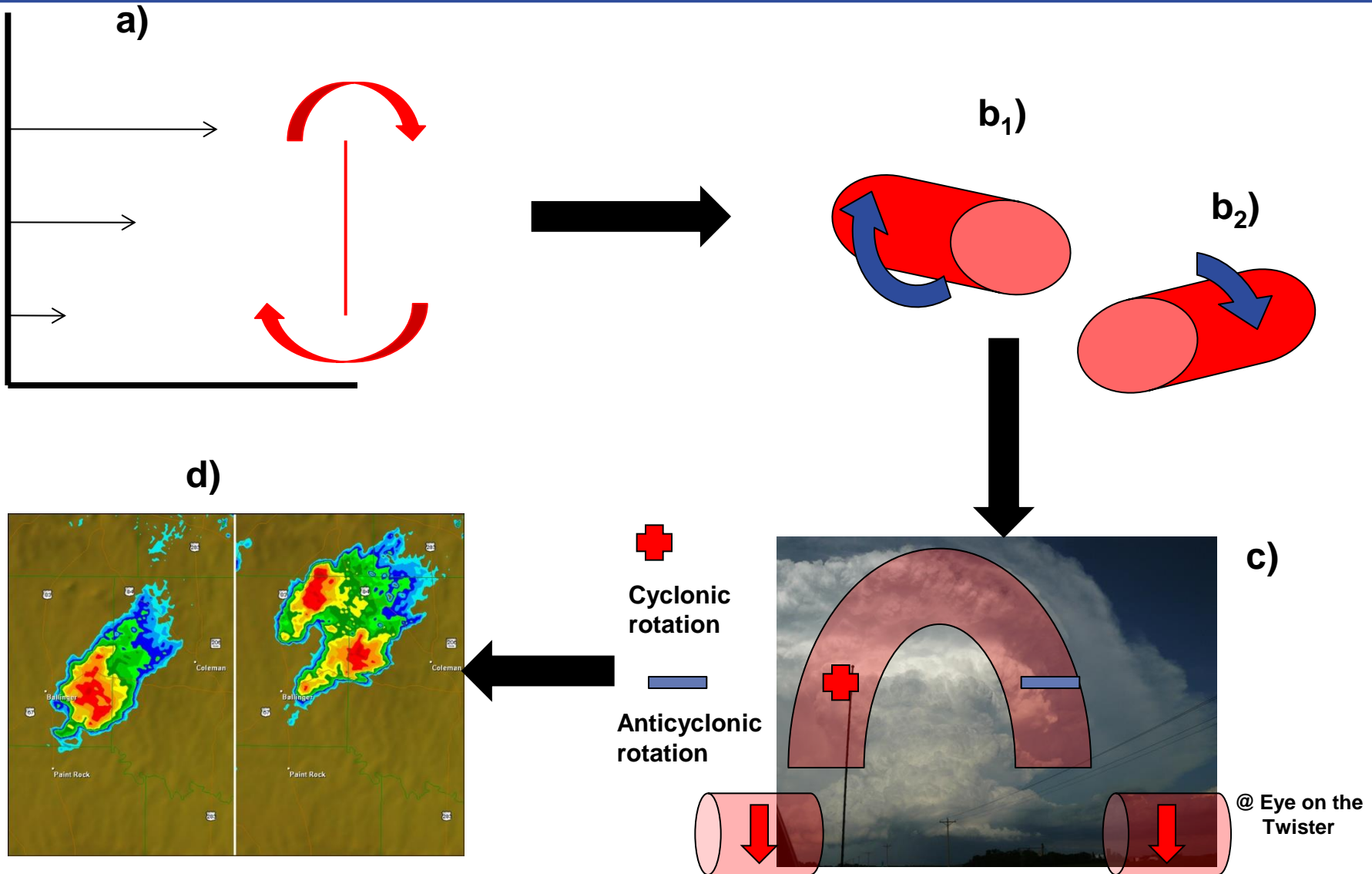


Convective forecast parameters: speed shear



Speed shear

Deutscher Wetterdienst
Wetter und Klima aus einer Hand

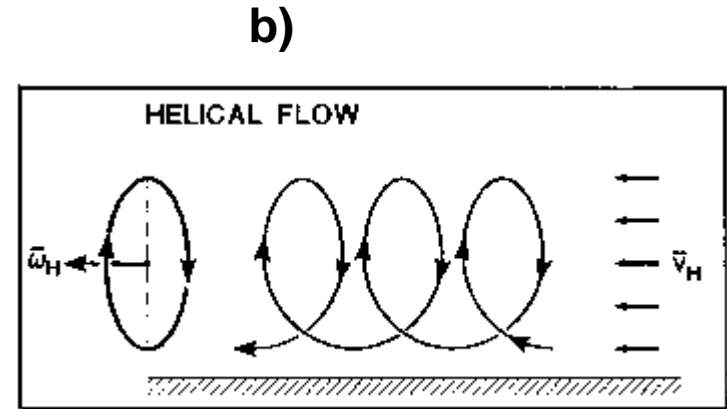
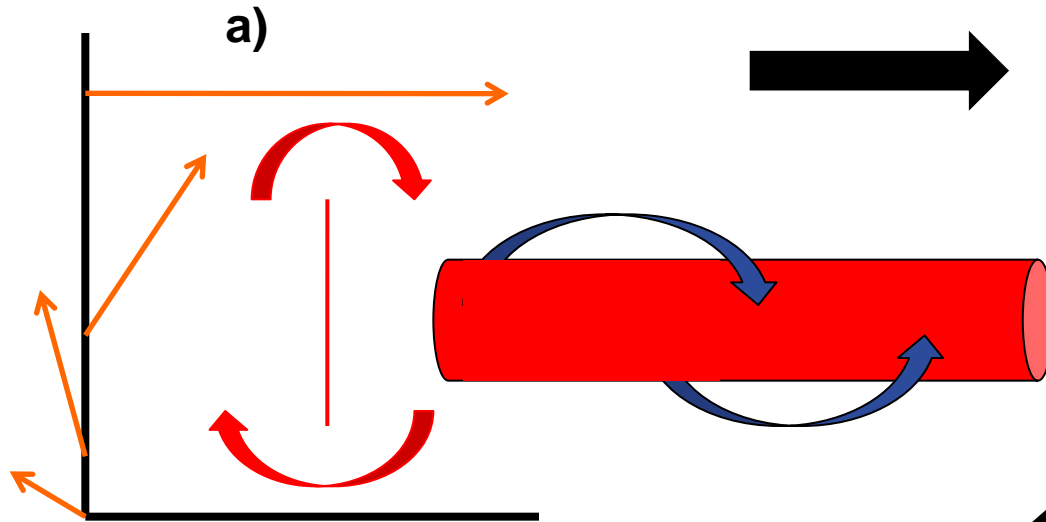


Convective forecast parameters: directional shear

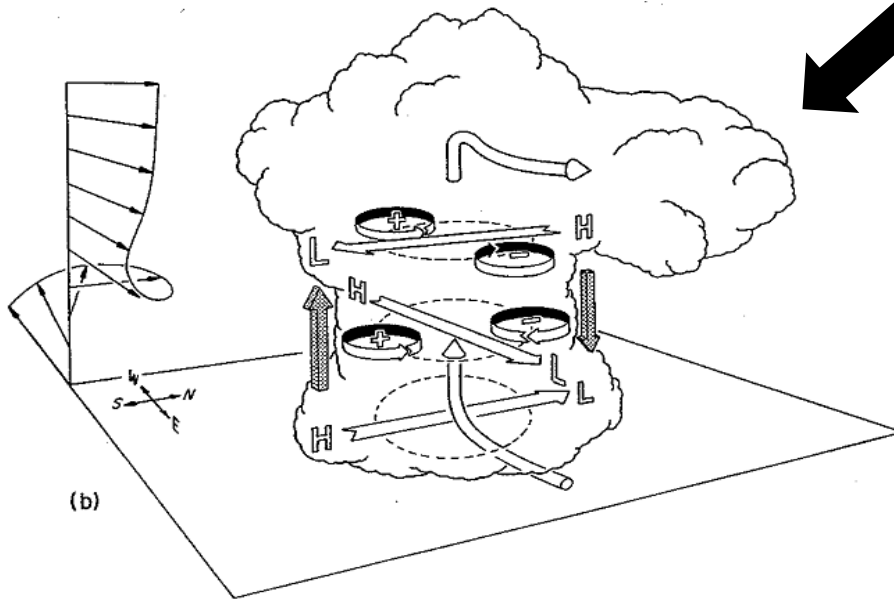


Speed and directional shear (helicity)

Deutscher Wetterdienst
Wetter und Klima aus einer Hand



@ A REVIEW FOR FORECASTERS ON THE
APPLICATION OF HODOGRAPHS TO
FORECASTING SEVERE THUNDERSTORMS
by Charles A. Doswell III



@ Dynamics of tornadic thunderstorms
by Joseph B. Klemp



What is an hodograph ?

180 °

Deutscher Wetterdienst
Wetter und Klima aus einer Hand



90 °

270 °

0 °

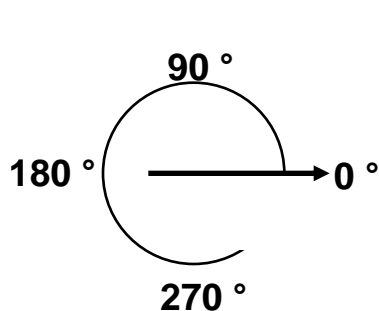
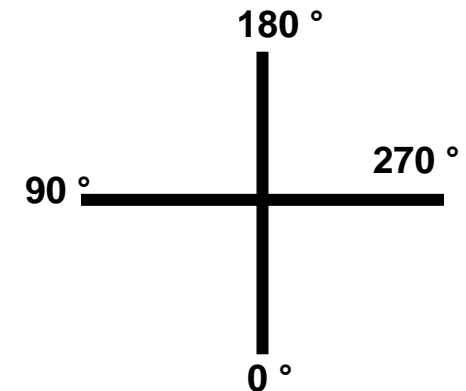
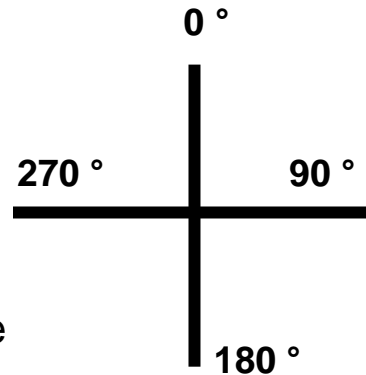
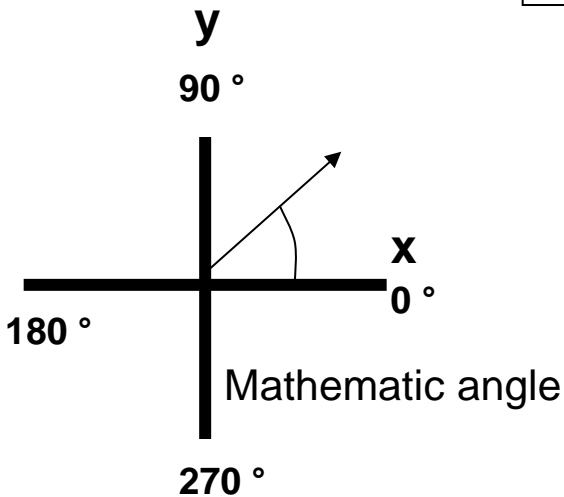




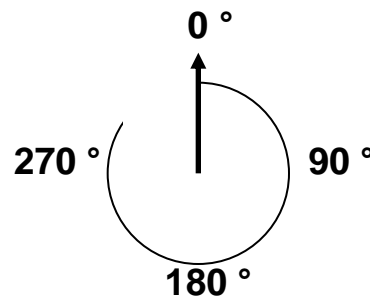
wind direction is defined to the y-axis



- **180°**: in meteorology, we look, where the wind comes from



Mathematical



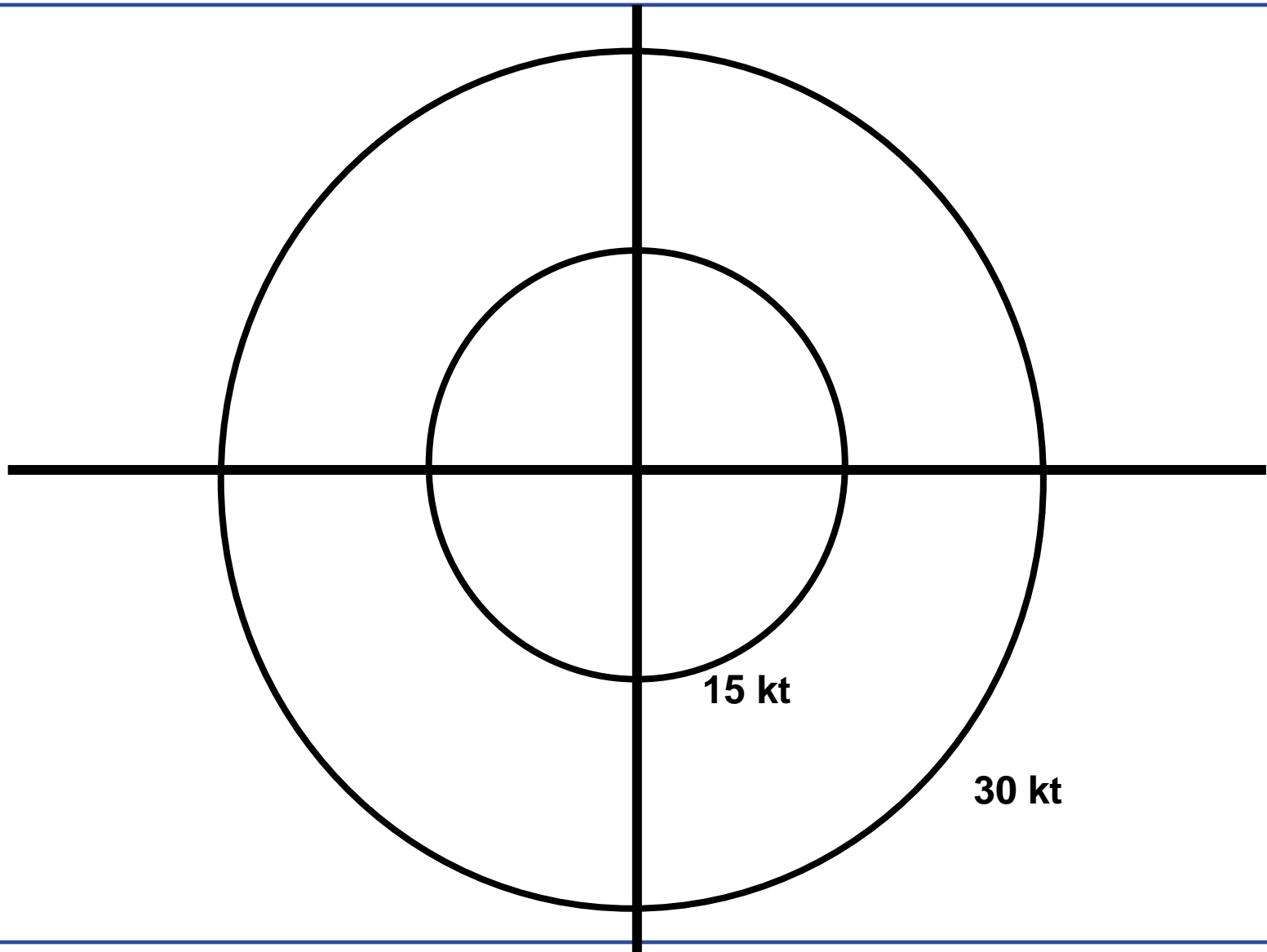
Meteorological

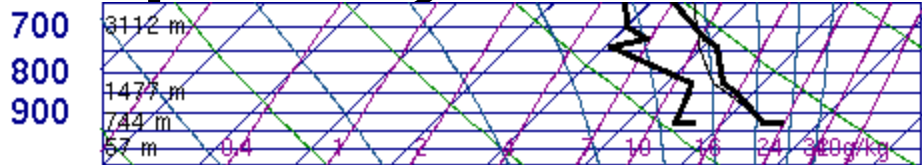
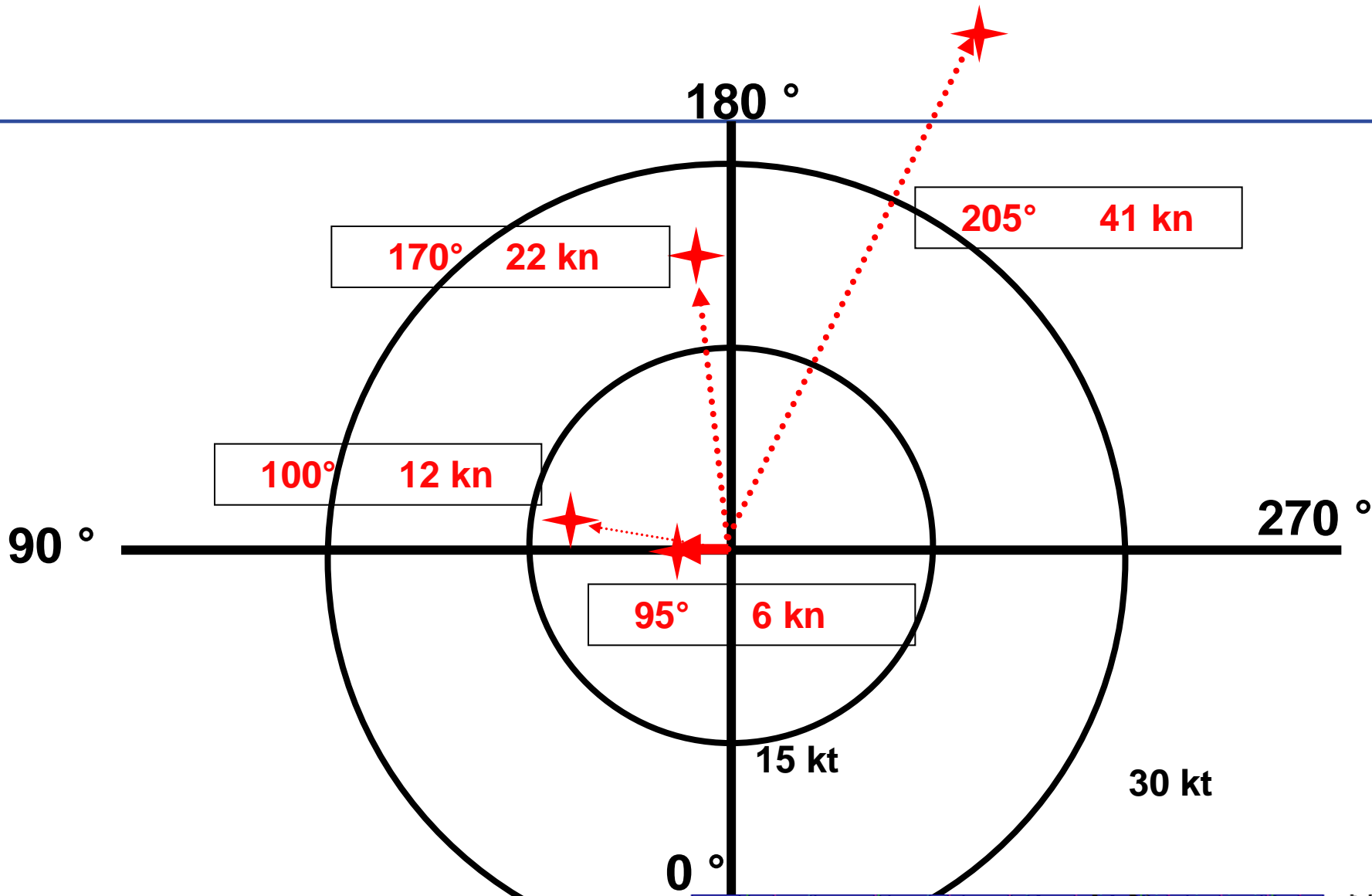
- ✓ north wind = from 0° or 360°
- ✓ east wind = from 90°
- ✓ south wind = from 180°
- ✓ west wind = from 270°

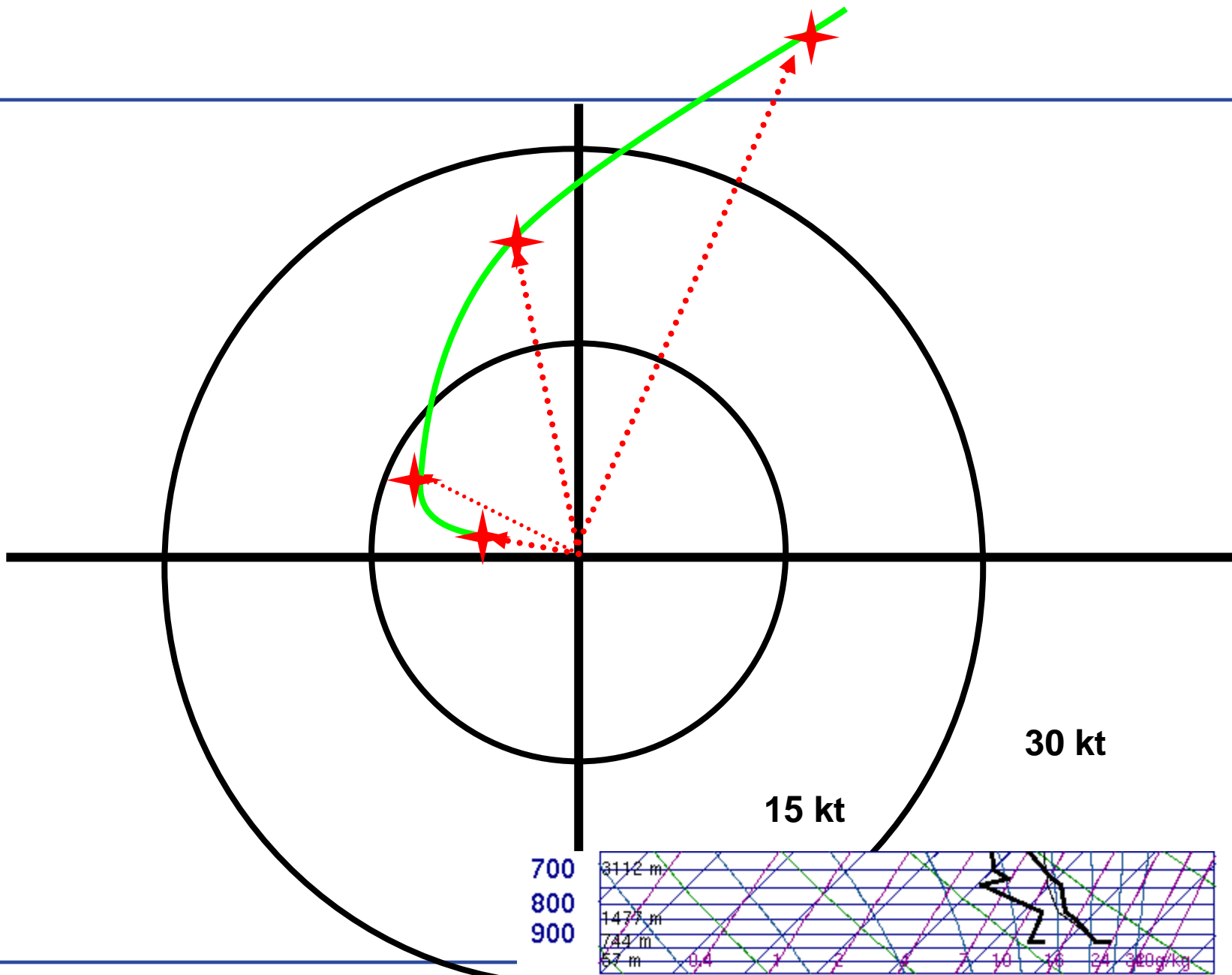
Hodograph

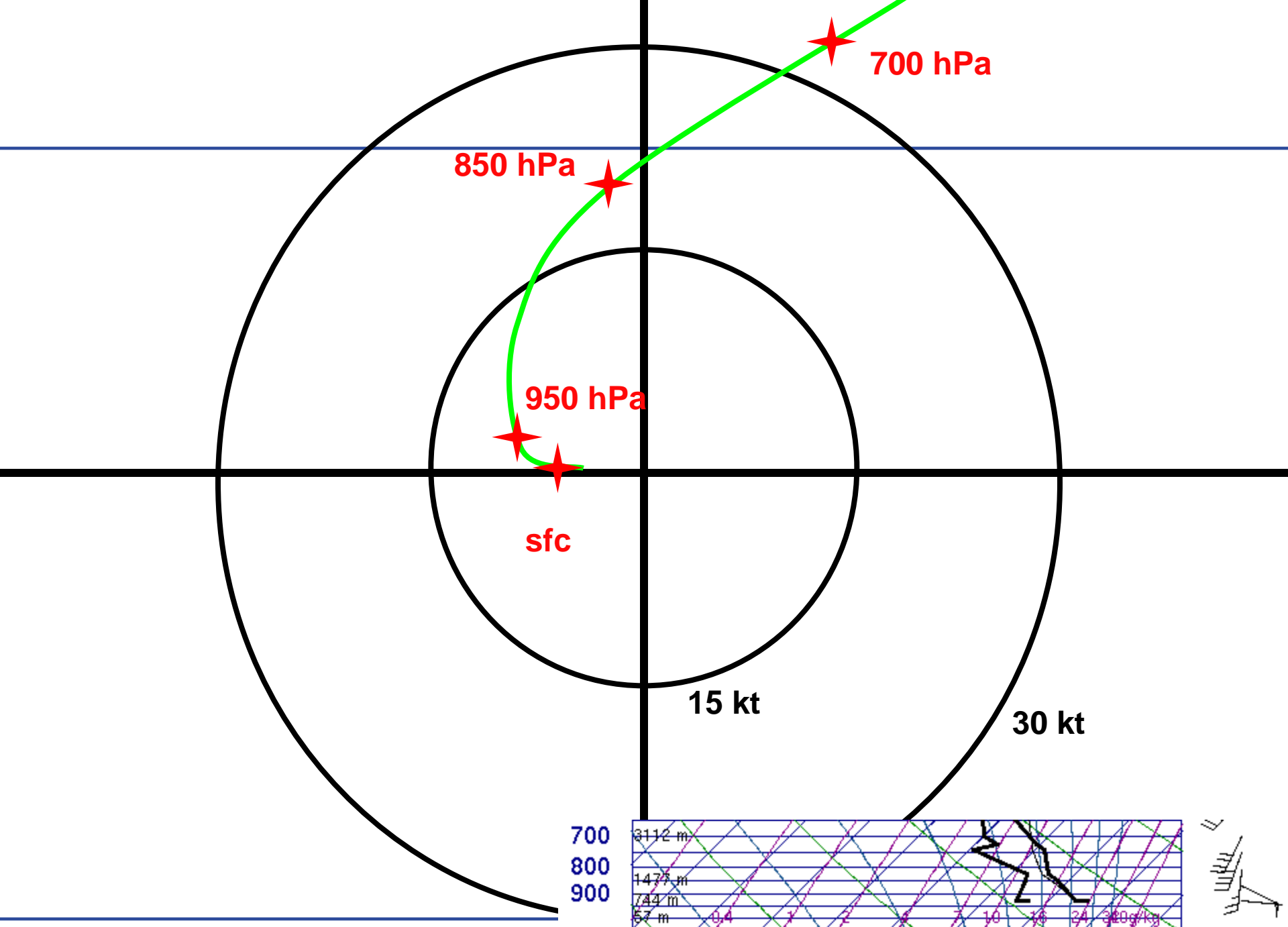
- ✓ north wind = pointing to the N
- ✓ east wind = pointing to the E
- ✓ south wind = pointing to the S
- ✓ west wind = pointing to the W











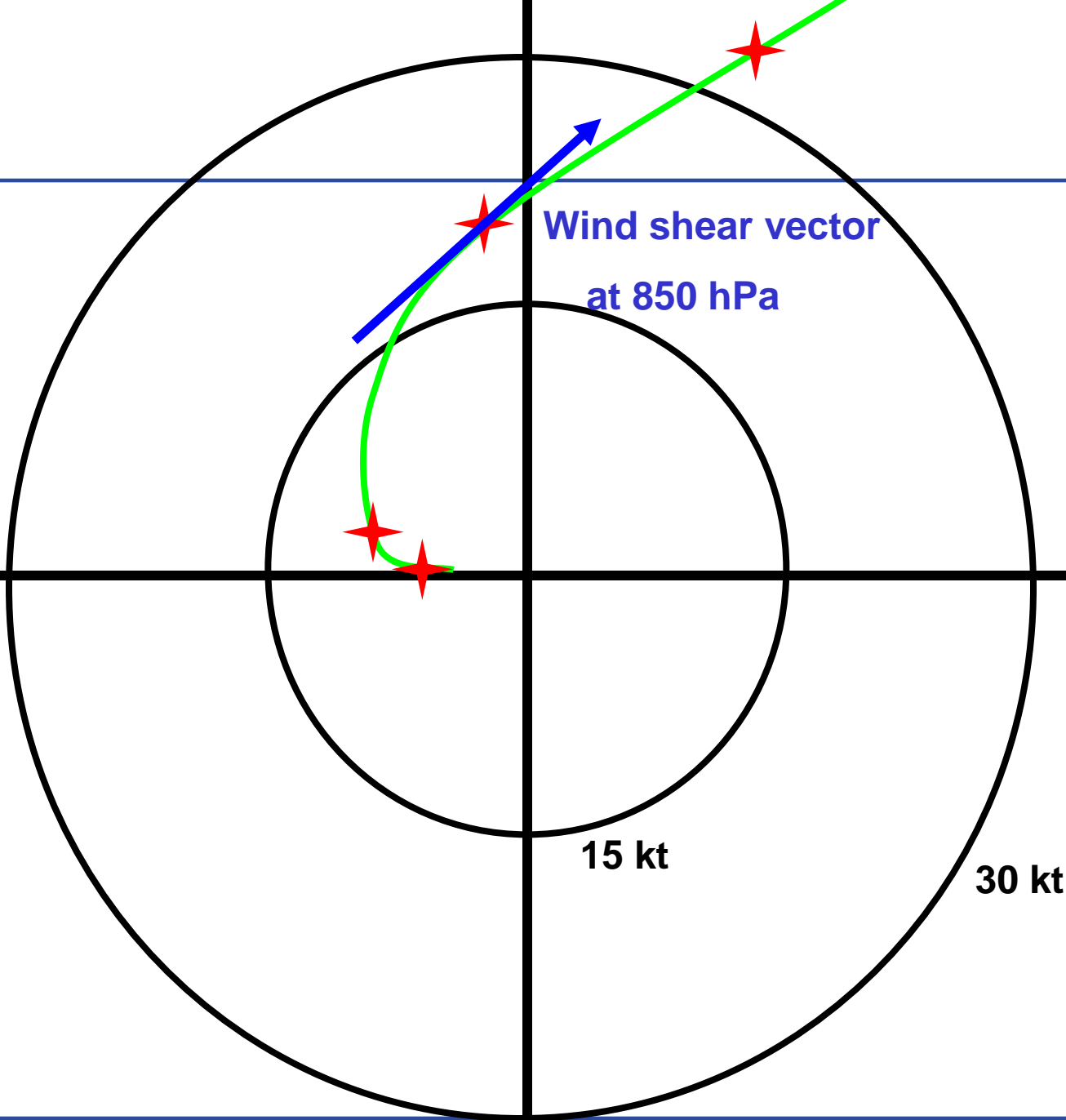


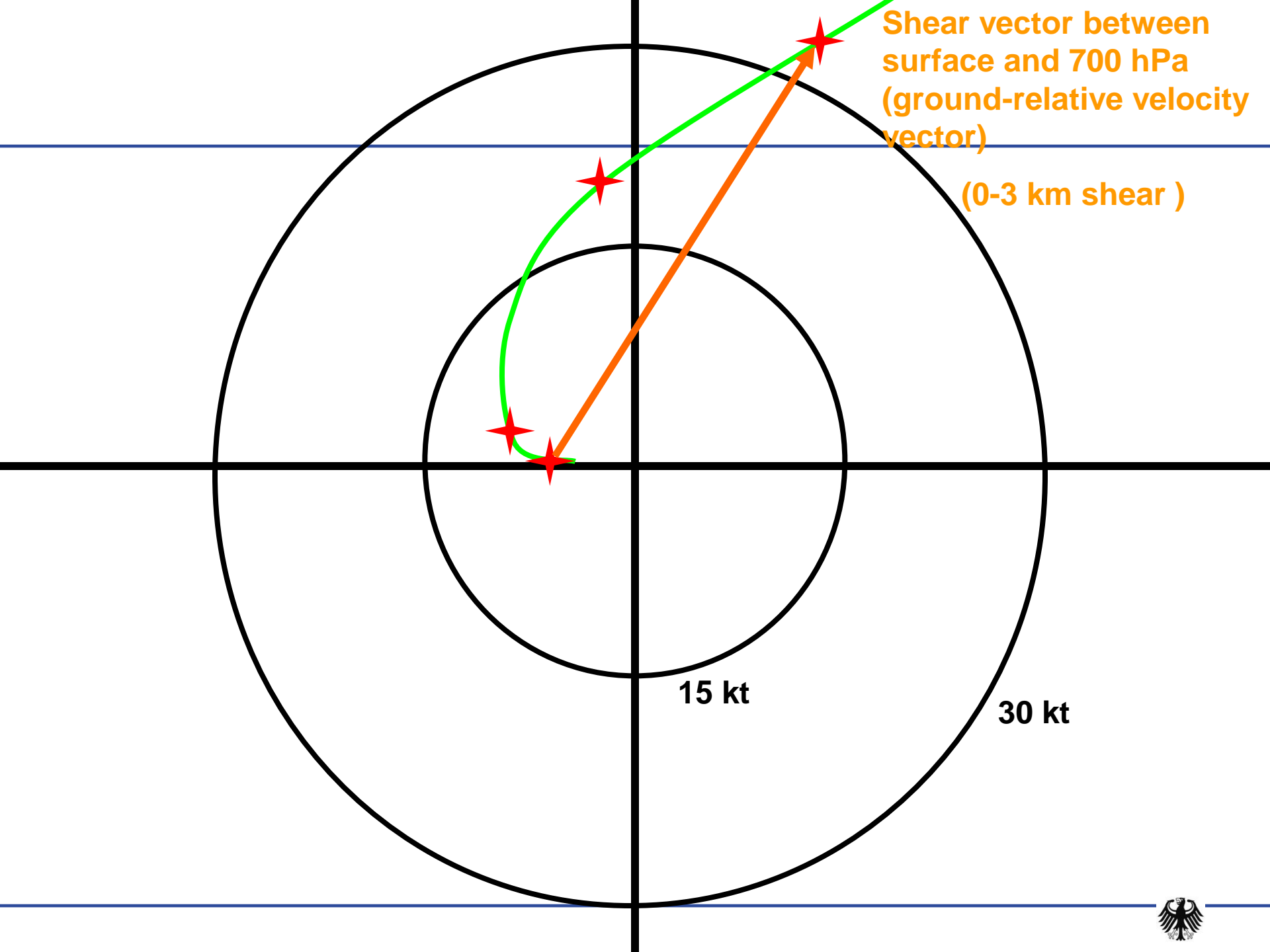
- Wind speed increases (significantly) with height
- Wind changes direction with height

Wind shear vector

- tangent to the hodograph, pointing towards increasing altitudes
- Shear between 2 levels is vector between both levels (thermal wind vector)

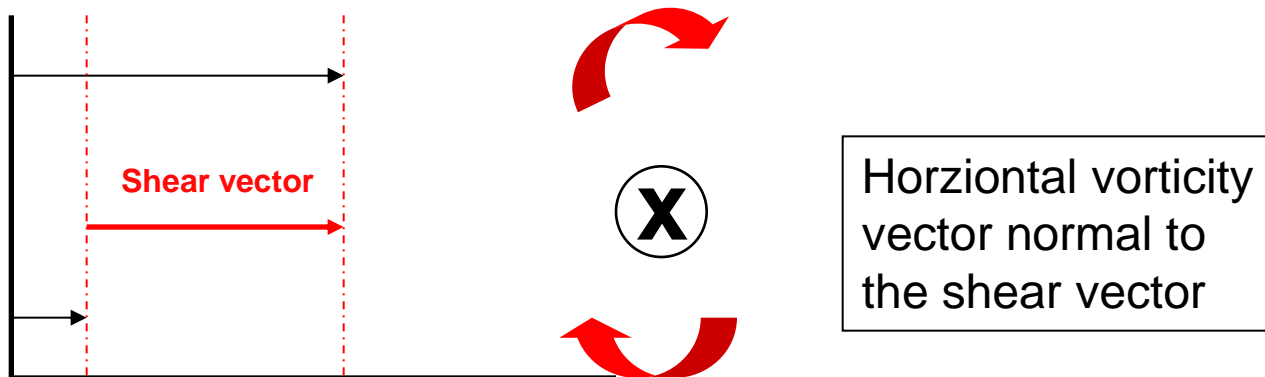






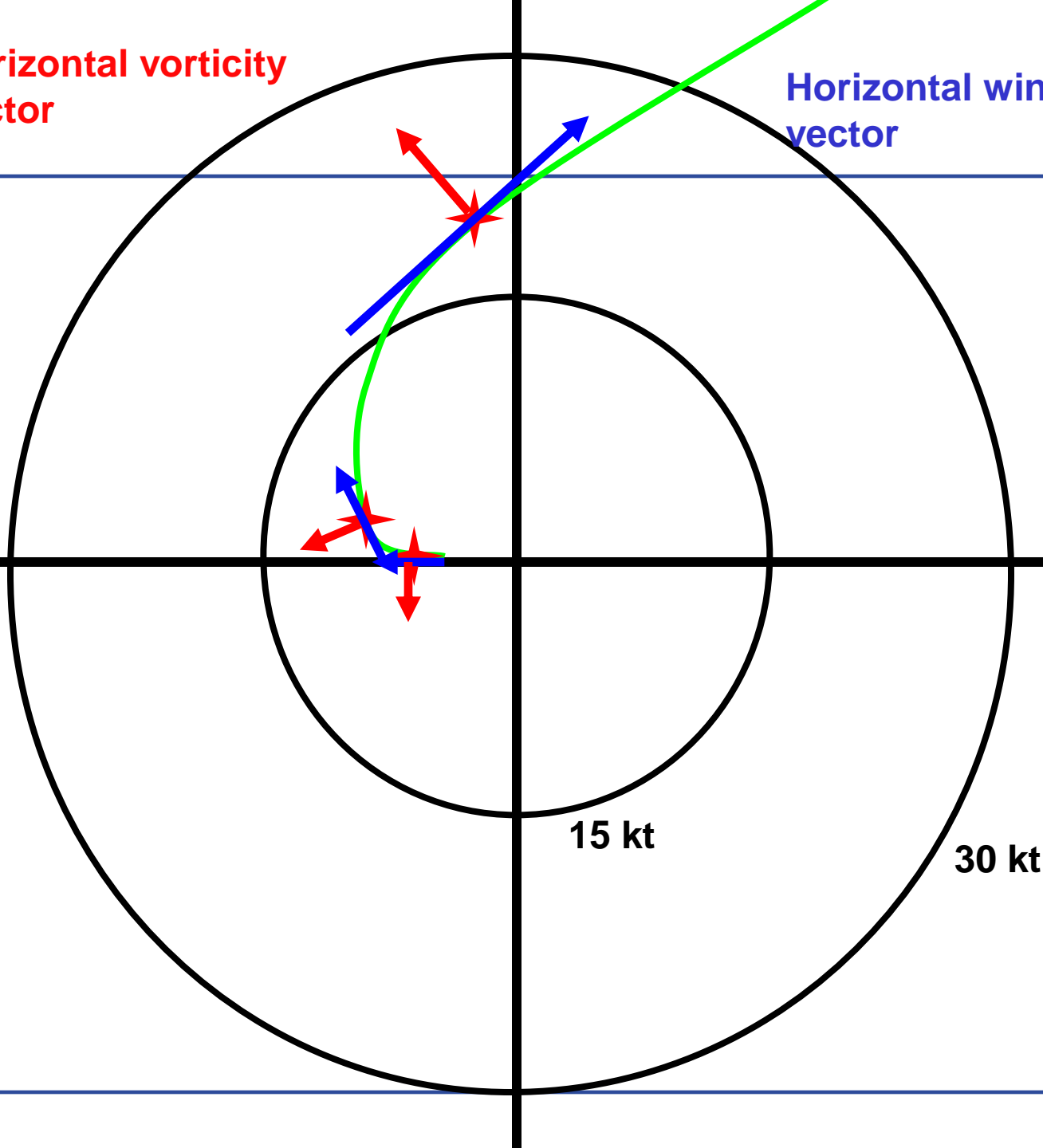
Horizontal vorticity vector

→ Achieved either by buoyancy gradient or shear



Horizontal vorticity
vector

Horizontal wind shear
vector

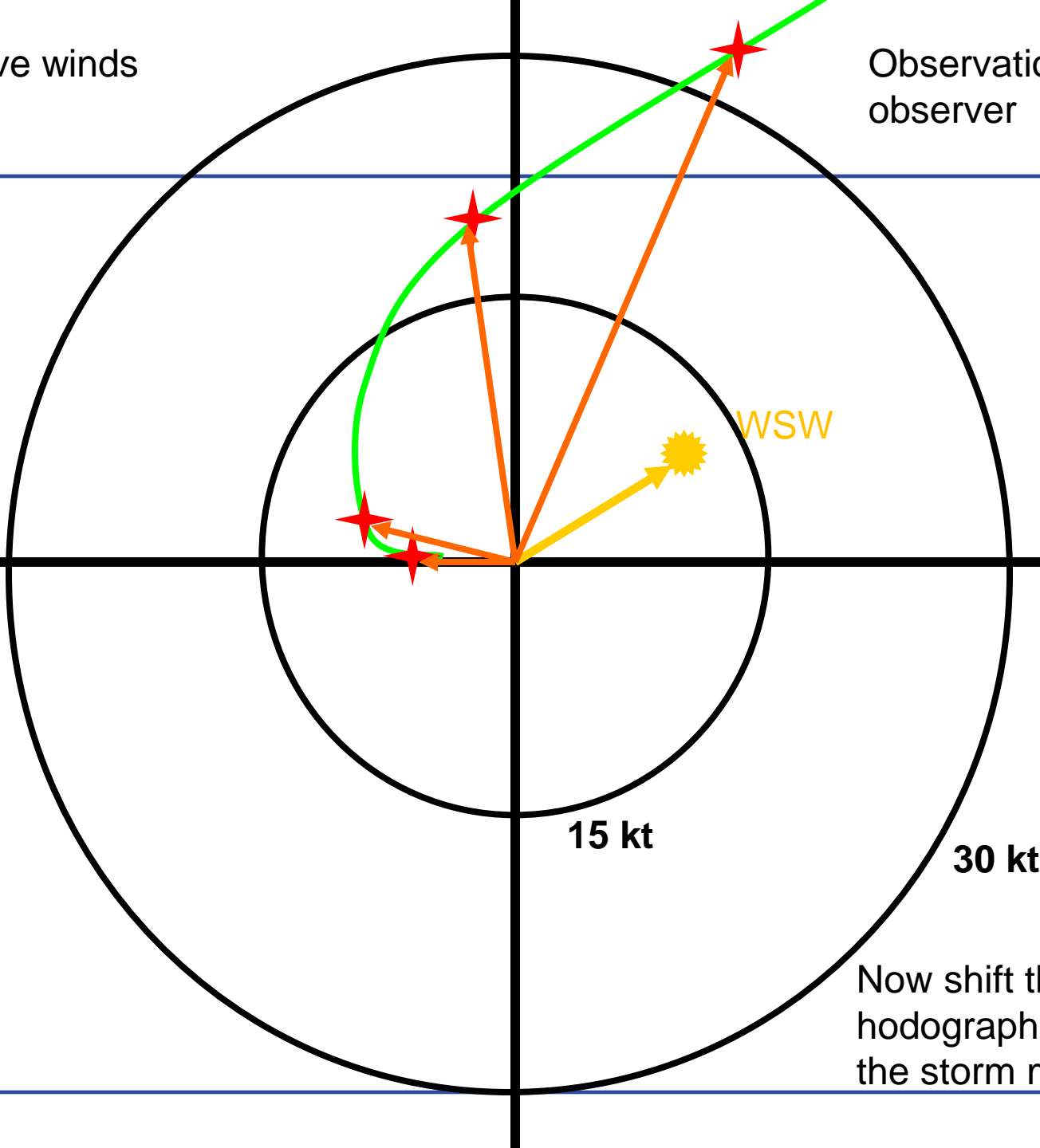


Storm relative wind

- Storm motion needed (e.g. Bunkers 30R75)
- Important wind to determine expected storm dynamics
- To see rotating updrafts, the storm-relative winds have to be aligned to the horizontal vorticity vector (streamwise helicity)

Storm relative winds

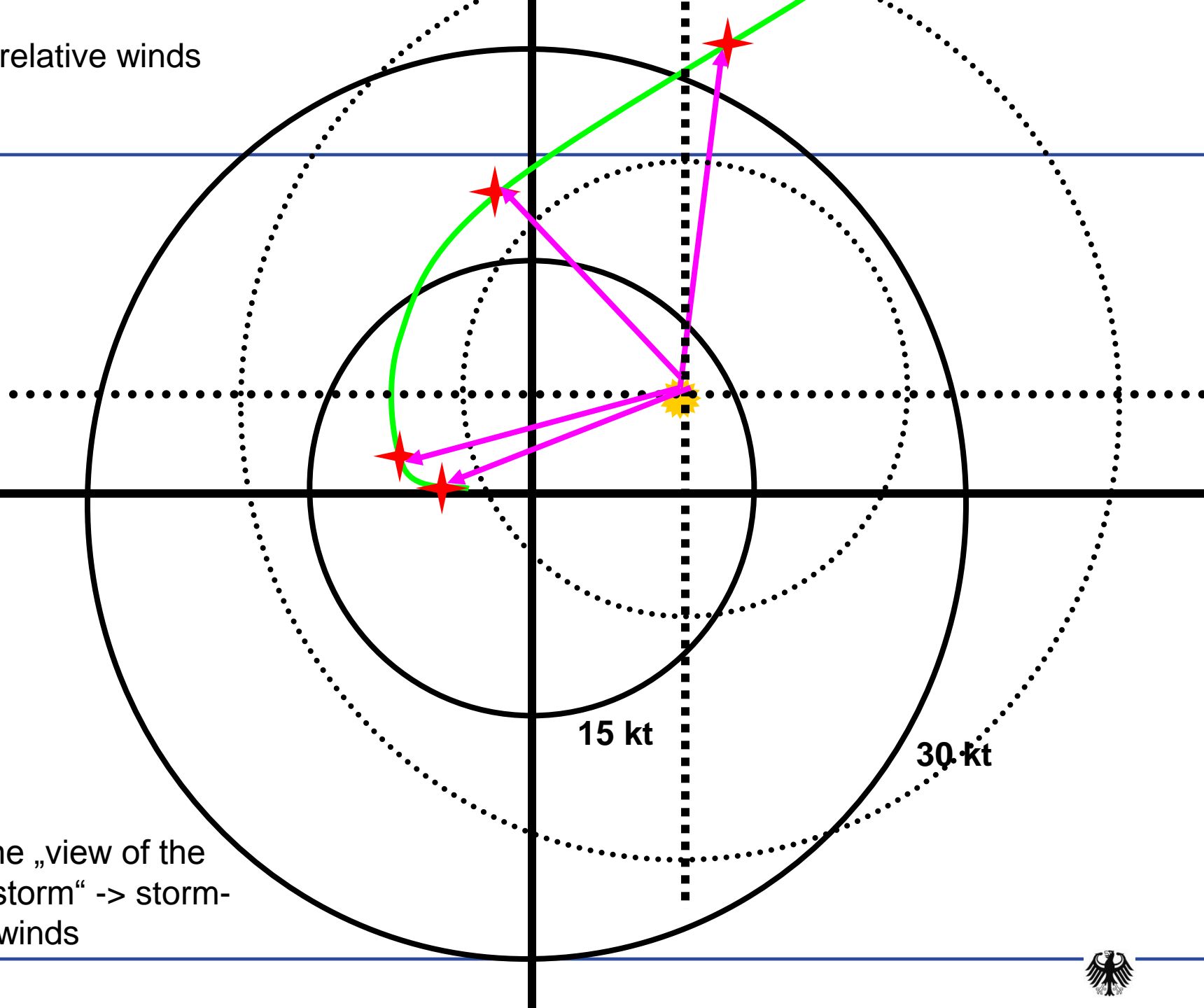
Observation for an observer



Now shift the
hodograph towards
the storm motion



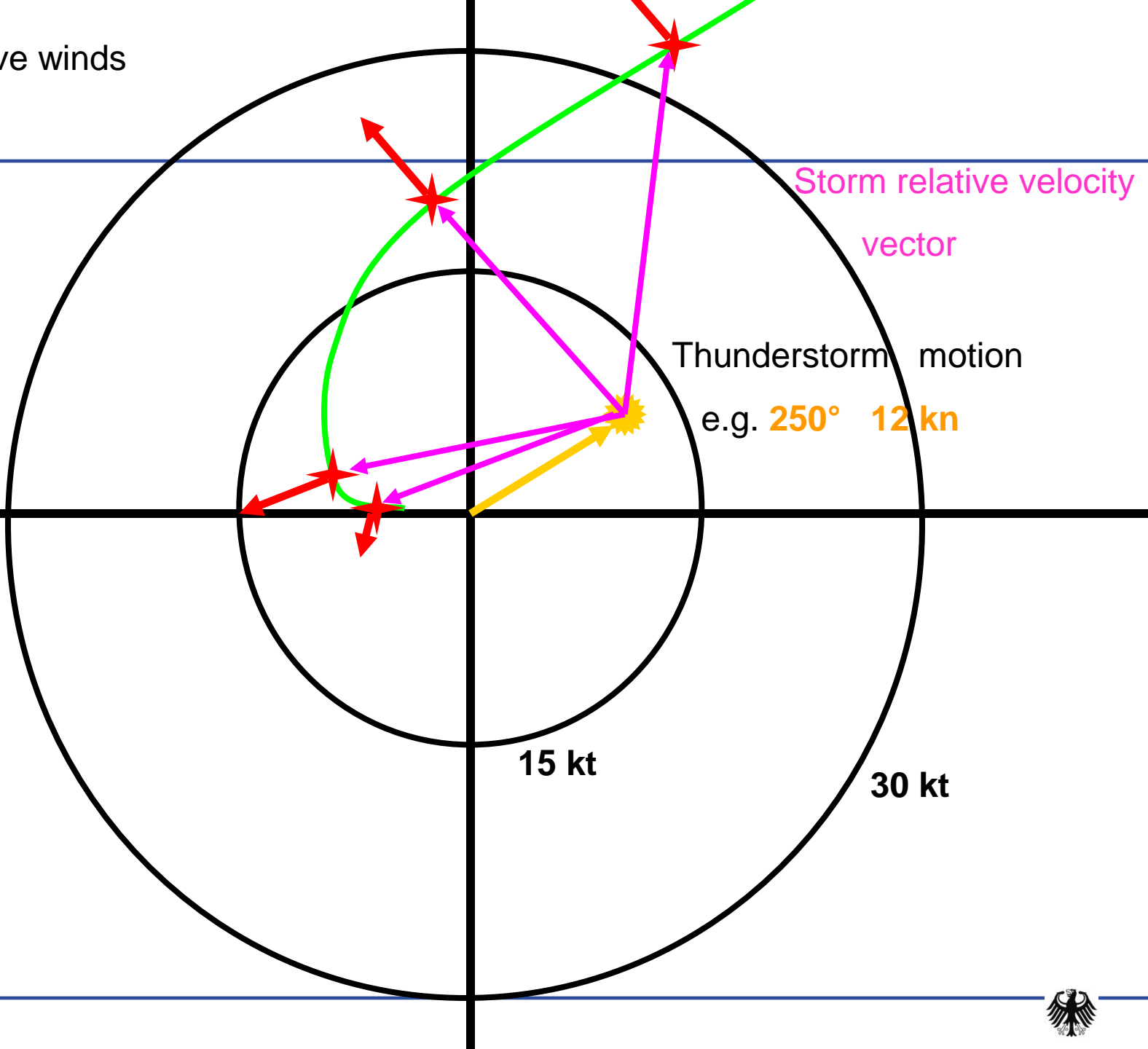
Storm relative winds



That's the „view of the
thunderstorm“ -> storm-
relative winds

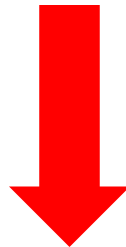


Storm relative winds



Favorable environment for organized updrafts

- Strong storm-relative winds within the low-levels (at least 10 m/s)
- Storm-relative vectors and horizontal vorticity vectors have to align nicely
- The wind shear vector has to reveal strong turning (veering or backing) with height



Storm relative helicity (SRH or SREH for „environmental“)

Storm relative winds

Storm relative helicity

Storm relative winds
(arrow) and storm relative
helicity (colored area);
here: 1-3 km SRH

Horiz. Vort.
vector

Thunderstorm motion

e.g. **320° 12 kn**

270 °

180 °

90 °

15 kt

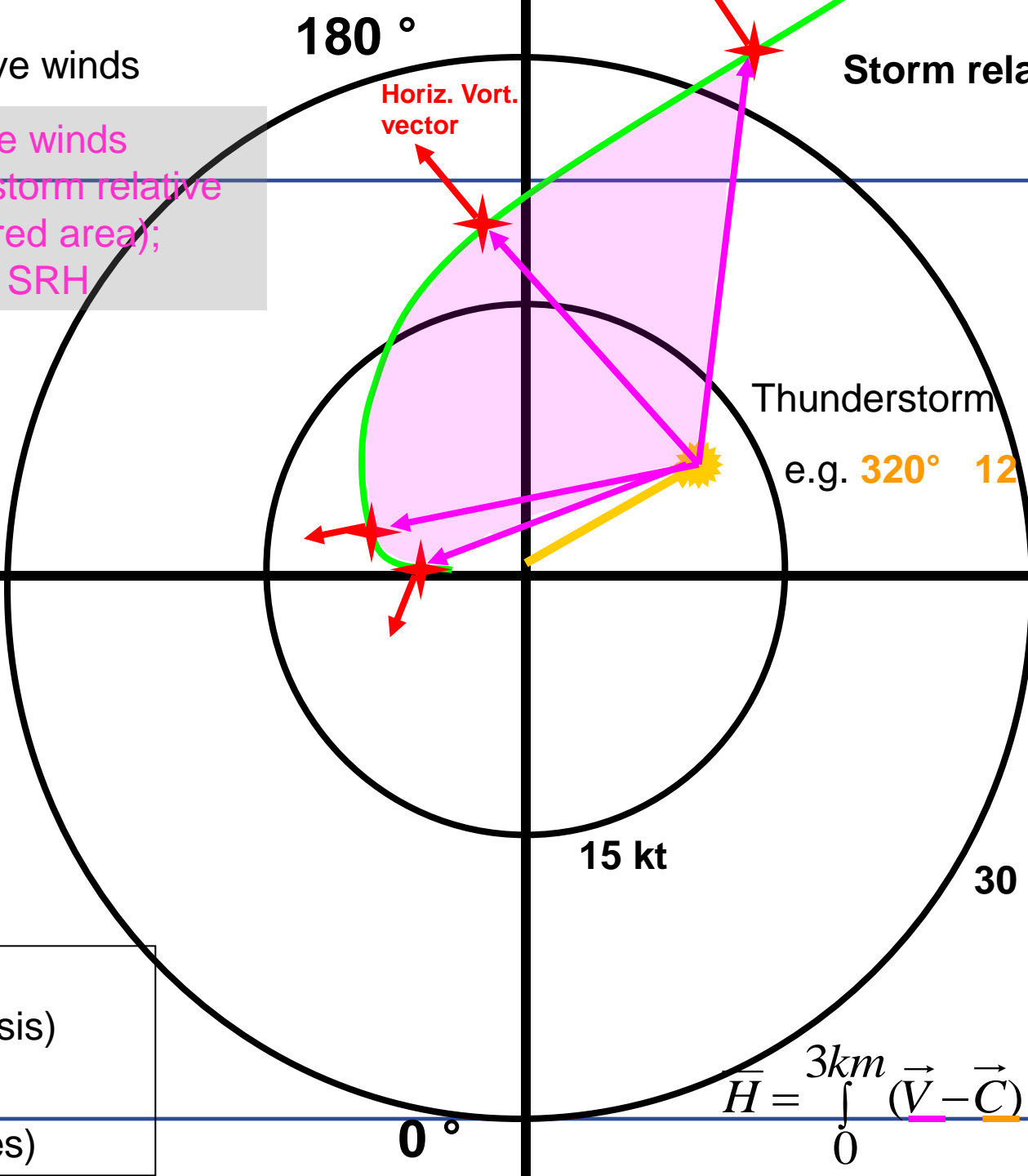
30 kt

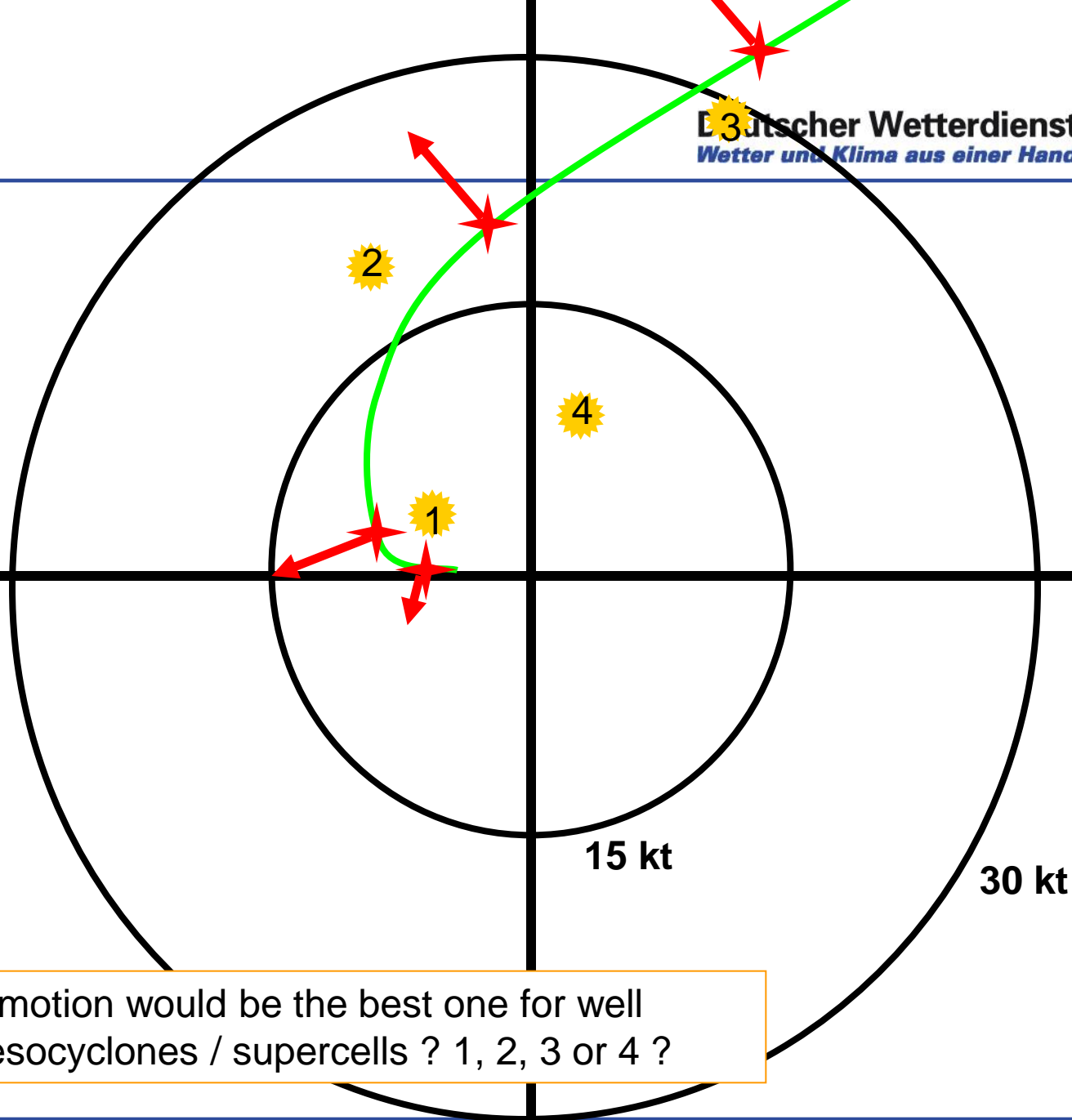
0 °

Sfc. To 1 km
(tornadogenesis)

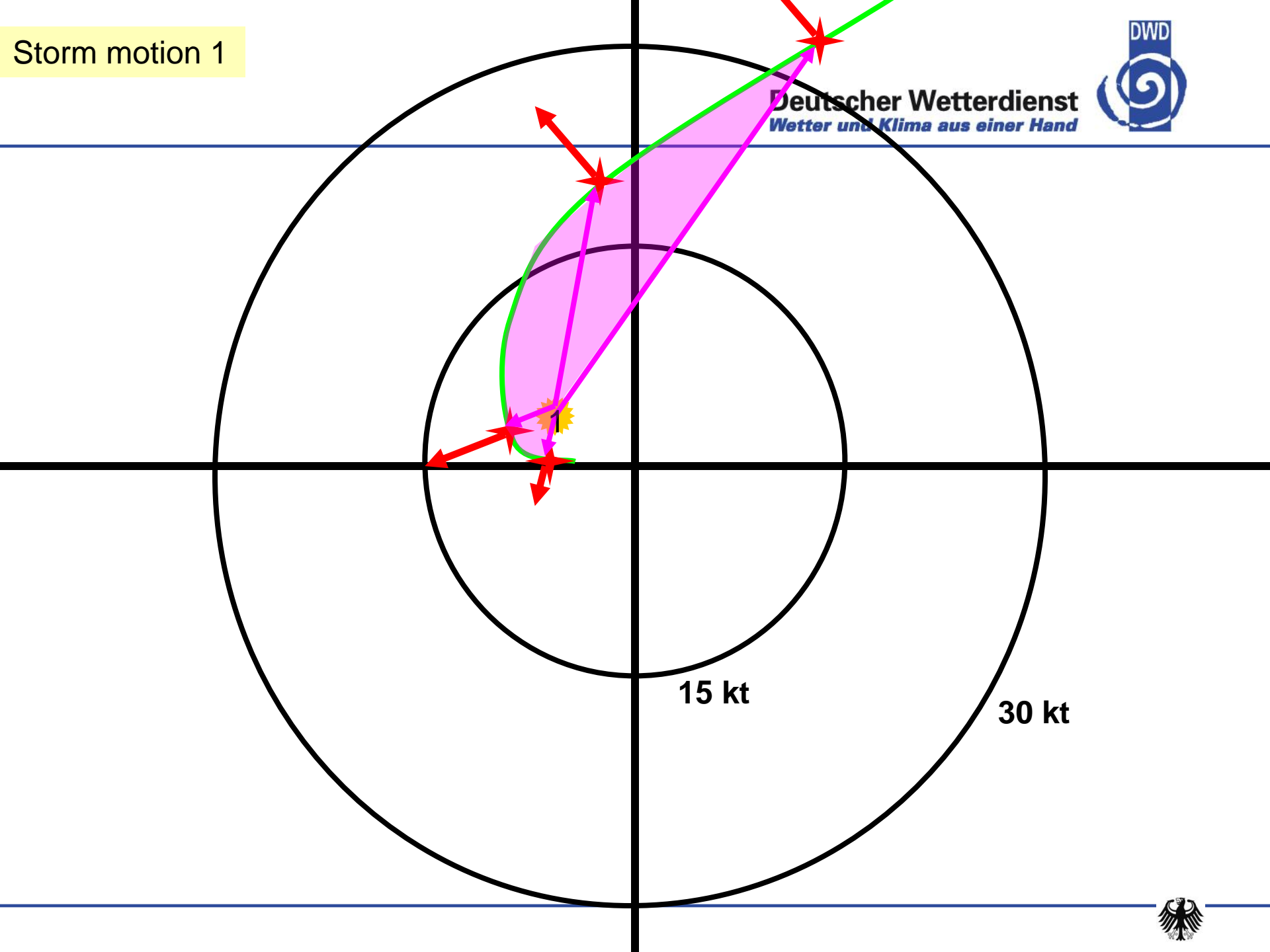
Sfc. To 3 km
(mesocyclones)

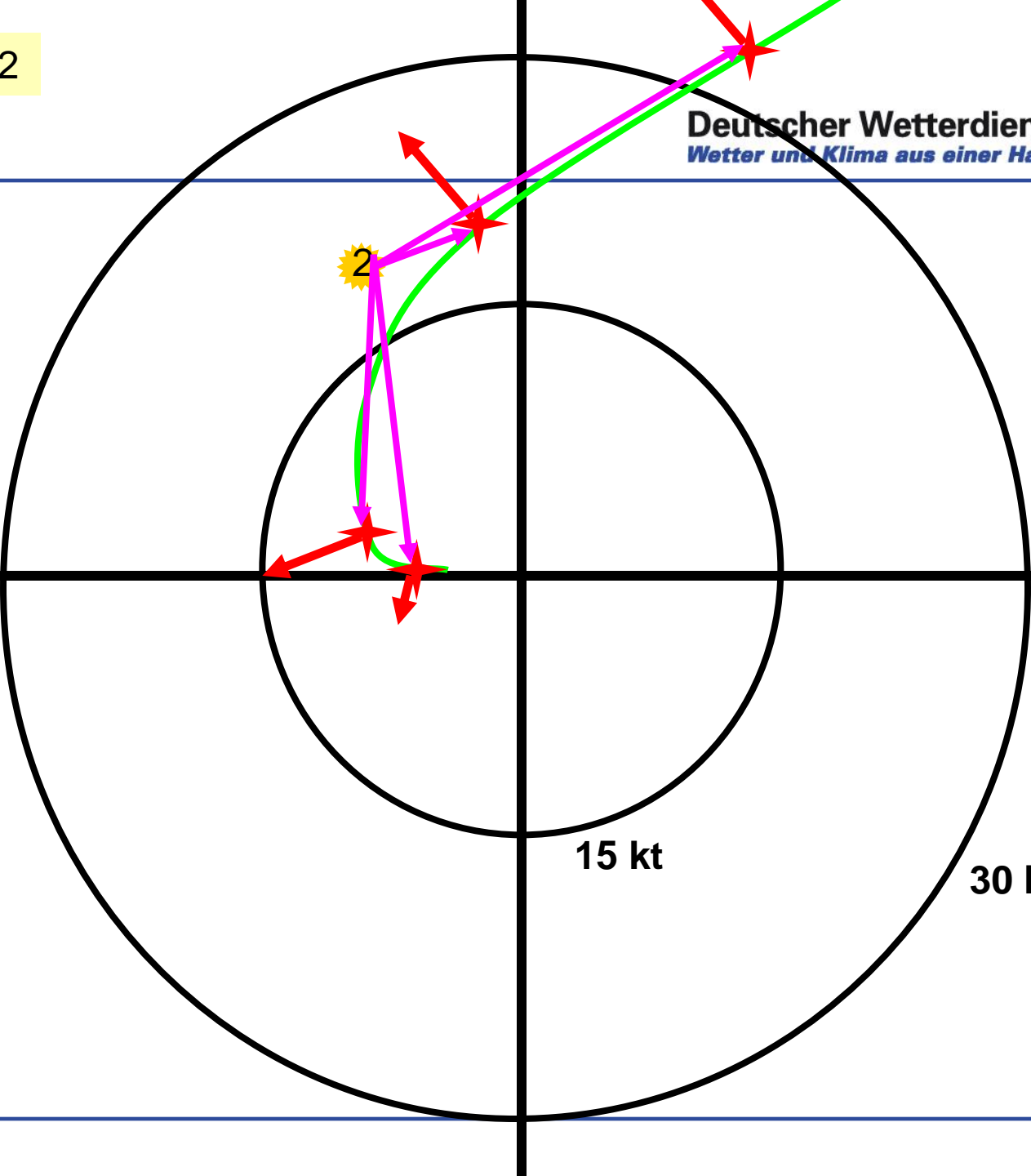
$$\bar{H} = \int_0^{3km} (\vec{V} - \vec{C}) * (\hat{k} \times \frac{\partial \vec{V}}{\partial z}) dz$$

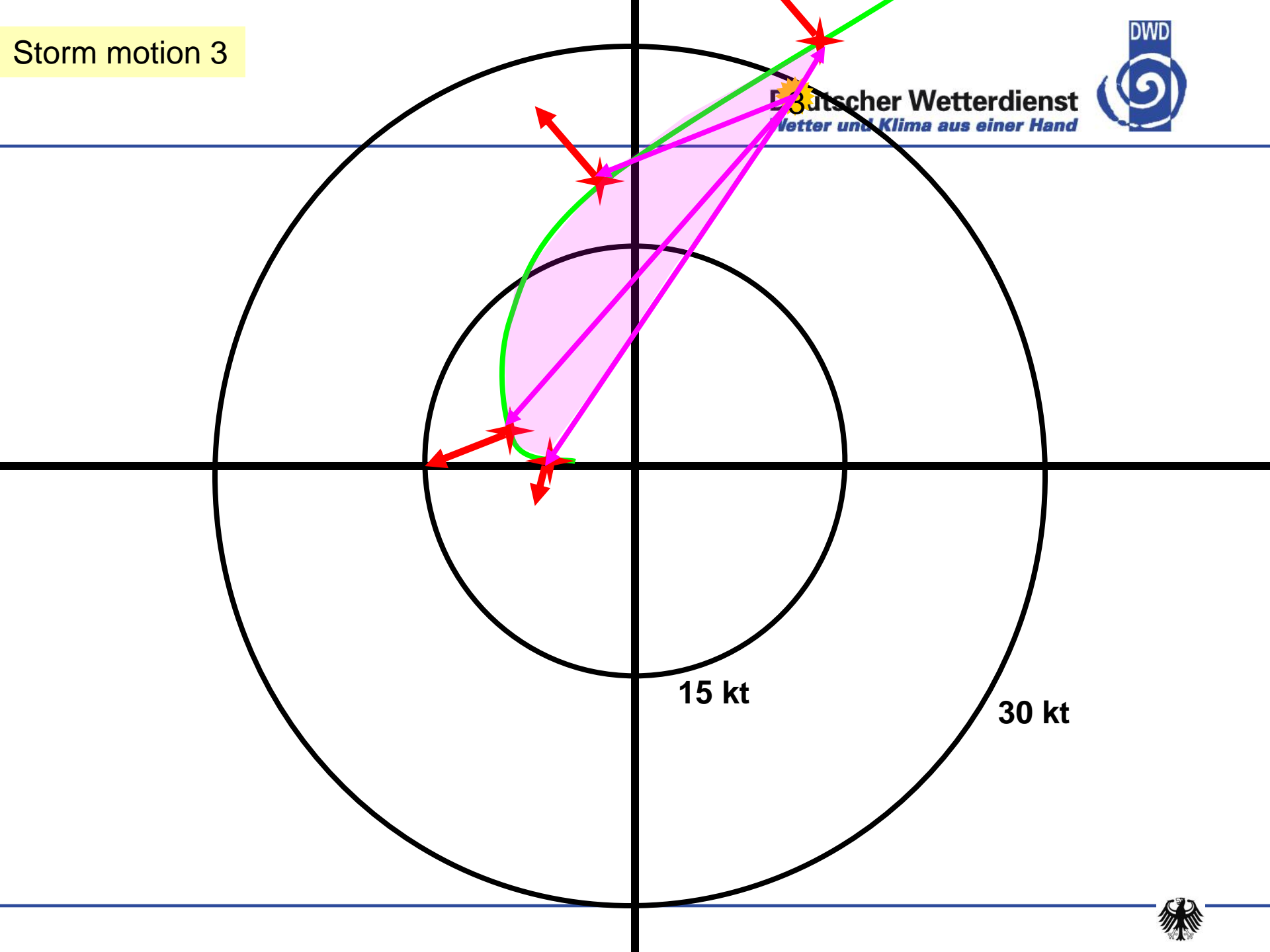


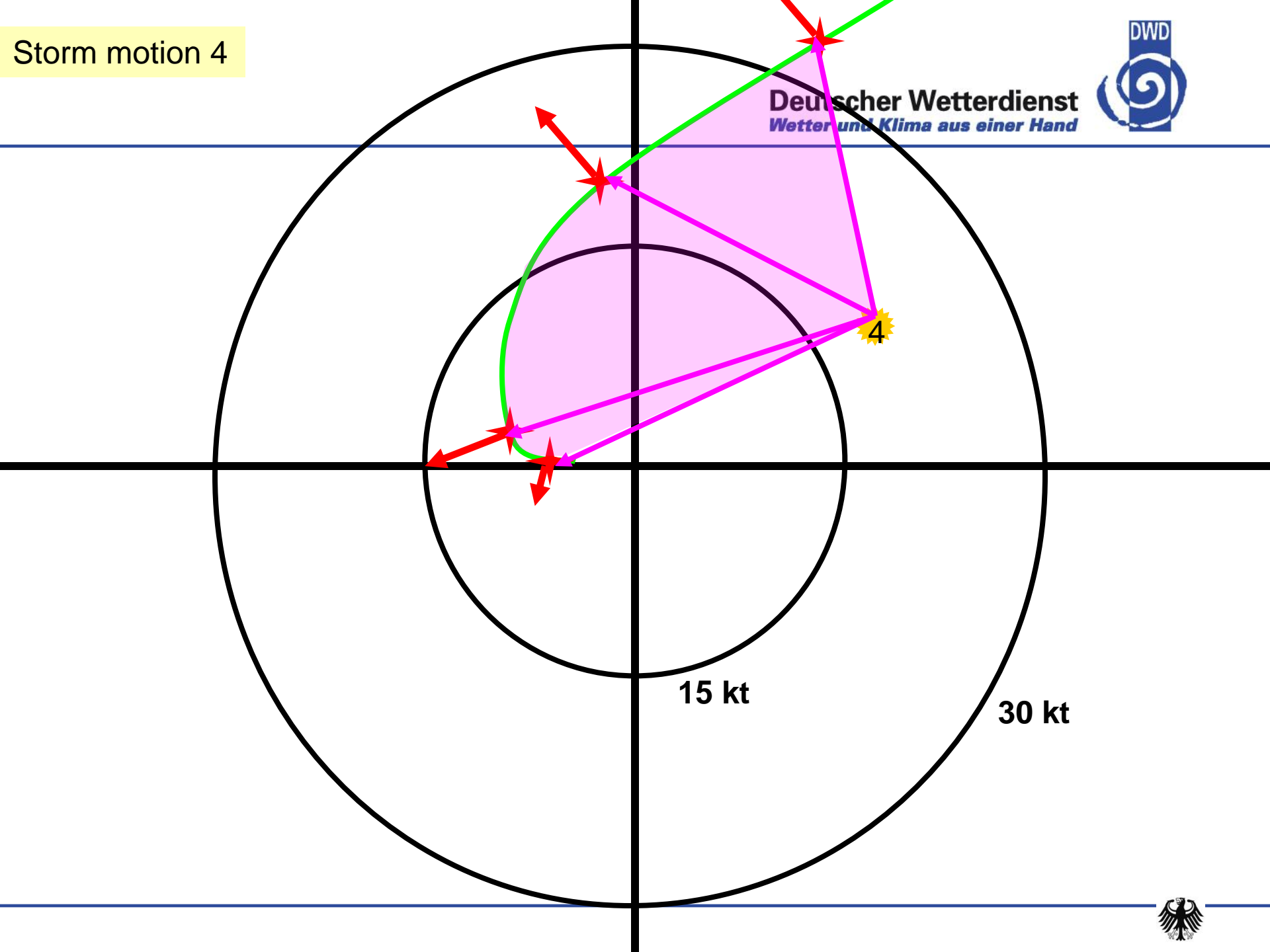


Which storm motion would be the best one for well organized mesocyclones / supercells ? 1, 2, 3 or 4 ?









Flash flood producing storms

The heaviest convective rainfall usually occurs in regions of high moisture, maximum ambient or elevated instability, best mesoscale lift and slow system movement.

@ Ted Funk, WFO Louisville, KY

The heaviest precipitation occurs where the rainfall rate is highest for the longest time

@ Charles Chappell

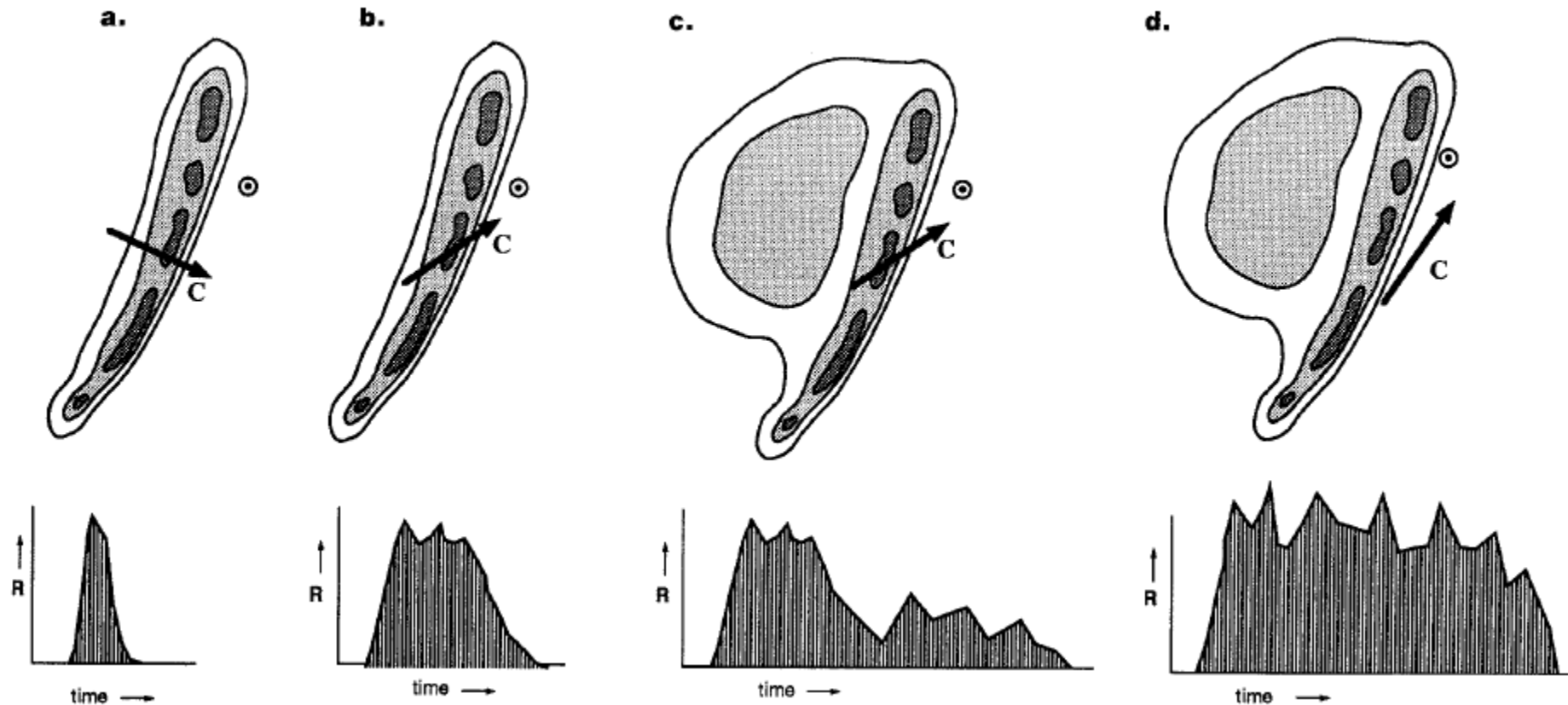
Hydrology

Orography

Long-time moisture

Land use

What causes storms to become flash flood producers?



Extreme rainfall with supercells



Most supercells **don't** offer **long-duration** excessive rain. Instead they can produce extreme rainfall amounts on a shorter time-frame. An example of the Dallas-Fort-Worth supercell in 1995*:

5 – **15** – **60** min time-frame

231 – **210** – **115** mm/h respectively // 120 mm

and Orlando, Florida supercell in 1996* :

1 – **5** min time frame

330 – **222** mm/h (mesonet measurements) // 50 mm

- Final rainfall amount not exceptional for Texas, but rainfall rates were (**urban regions!**)
- Heaviest rain was observed during the weakening phase (collapsing) of supercell thunderstorms.

* Extreme rainfall and flooding from supercell thunderstorms

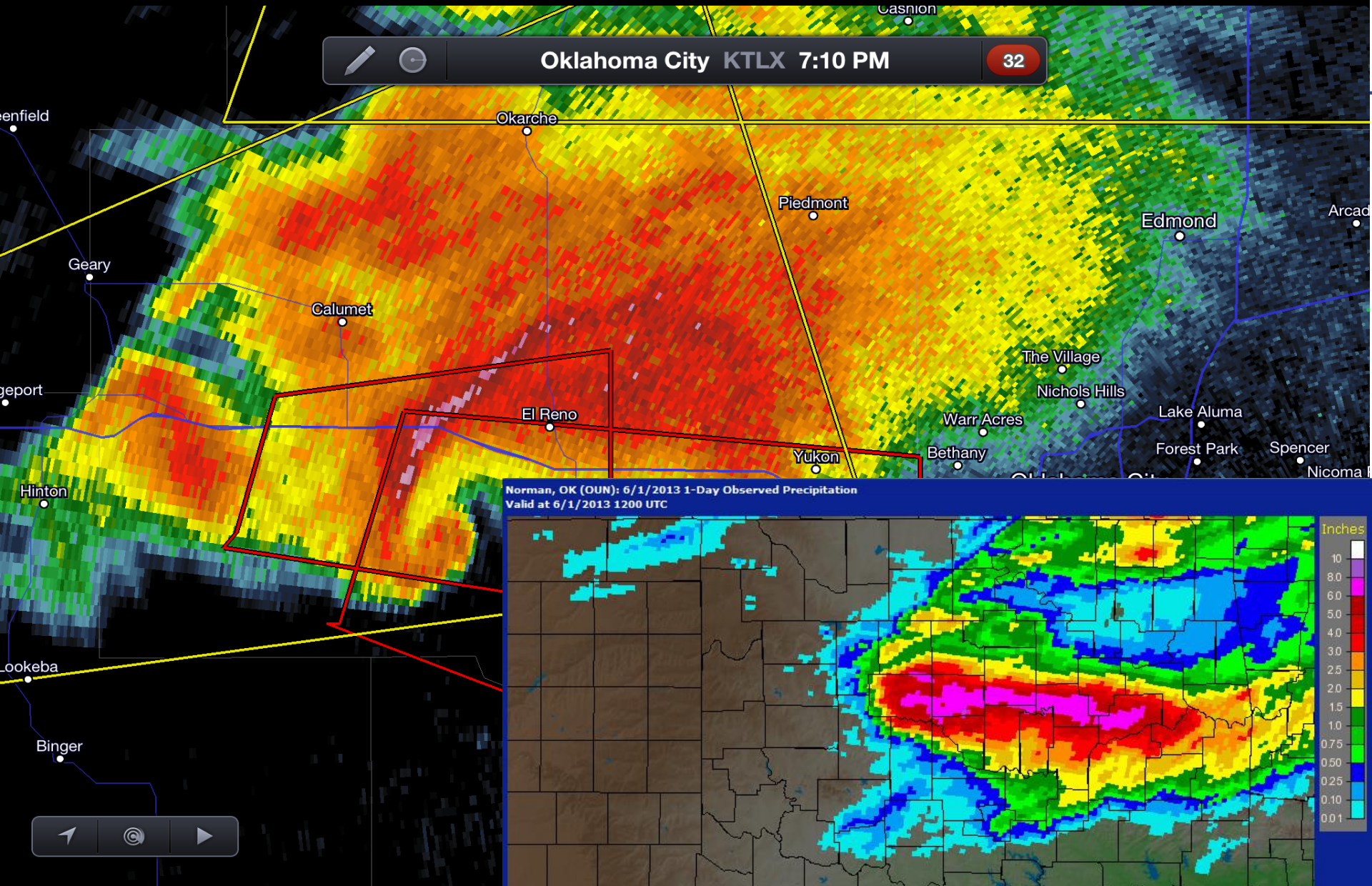
James A. Smith, Mary Lynn Baeck, Yu Zhang and Charles A. Doswell III

- ☺ Supercells have strong updrafts which ingest lots of moisture
- ☺ Supercell interaction with squall line (5th May 1995 Fort-Worth Dallas / 120 mm) , mesoscale boundaries (outflow) or orography -> influencing heavily storm motion
- ☺ Numerous supercells crossing a certain region with each preceeding storm worsening the flash flood problem

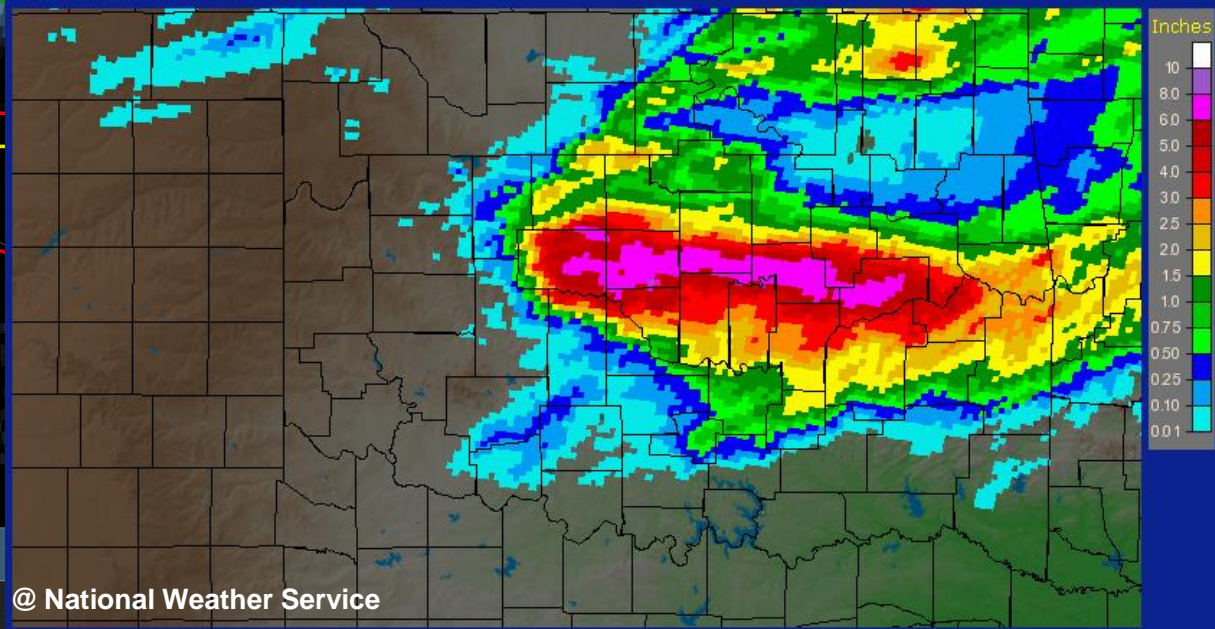
Radar measurements like rain sum not always helpful in such events due to the variable Z (radar reflectivity)-R(rainfall rate) relationship (also: **hail contamination**)

Hence rainfall maximum in radar representing hail with real rainfall maximum displaced!

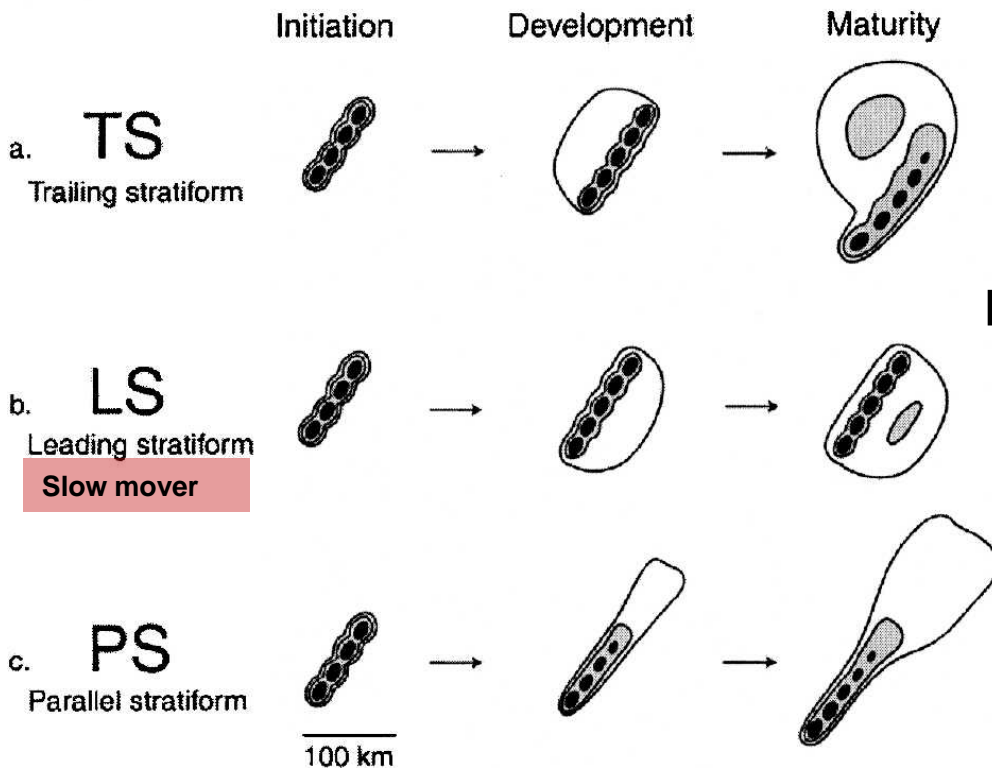
Polarimetric radar data could/should lower that problem



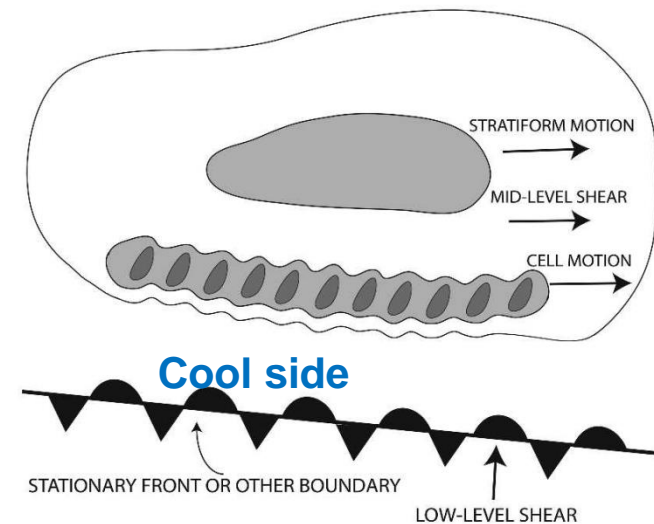
Norman, OK (OUN): 6/1/2013 1-Day Observed Precipitation
Valid at 6/1/2013 1200 UTC



Linear MCS archetypes



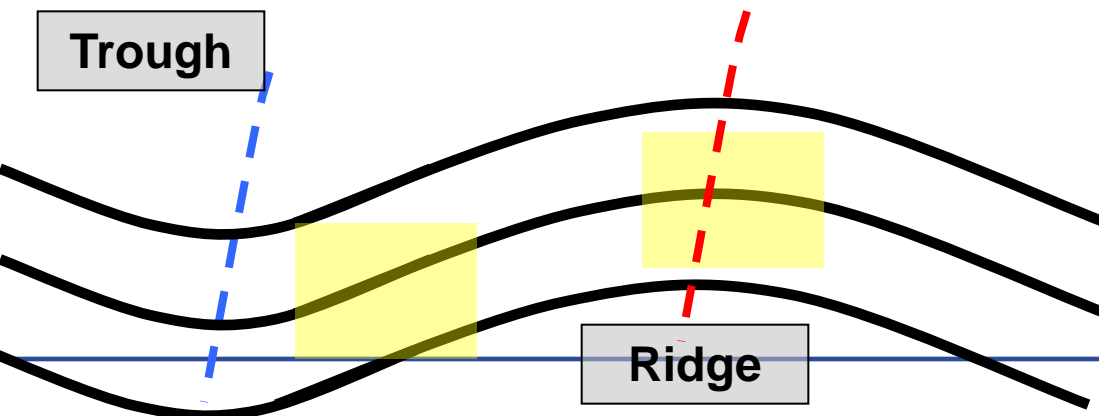
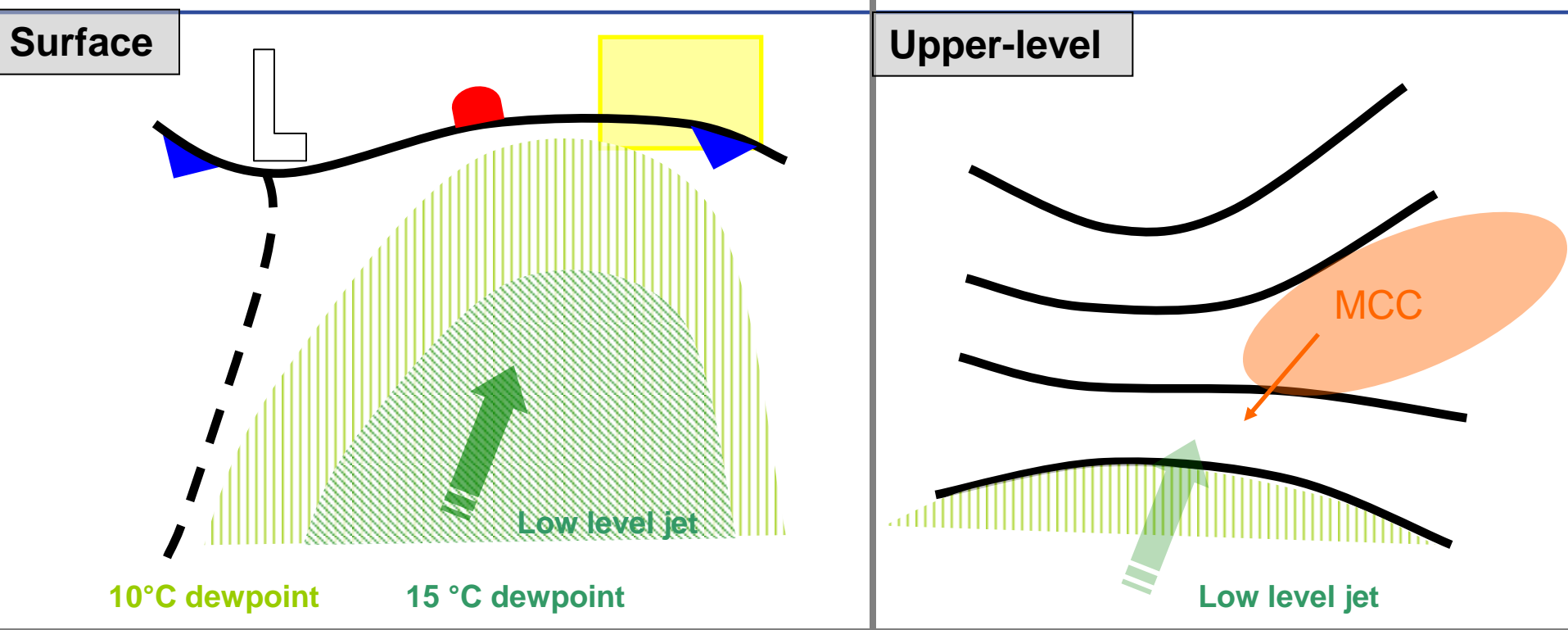
A) TRAINING LINE -- ADJOINING STRATIFORM (TL/AS)



- Large angle of low/mid-level shear vectors
- line-parallel mid-level shear **vs.** line-perpendicular shear of TS/LS

Frontal type heavy rainfall event Maddox et al. 79

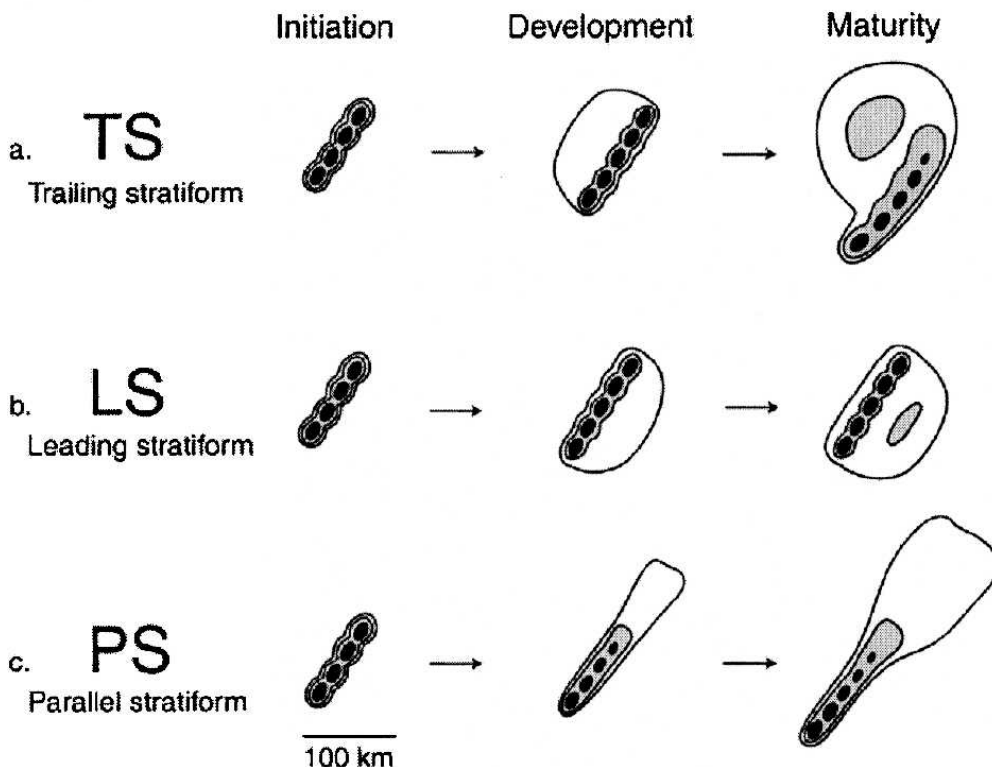
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Wetter und Klima aus einer Hand



- weak flow at mid-levels
- Propagation offset by cell motion → training



Linear MCS archetypes

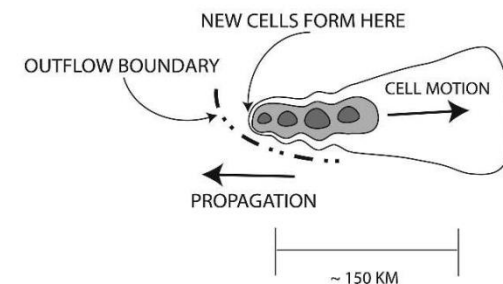


- Covering a smaller area but exceptionally high point rainfall totals e.g.

309 mm/24h 6-7th May 2000 St. Louis, Missouri

- Occur in a variety of surface wx. patterns
(**mesoscale/storm-scale!**)

B) BACKBUILDING / QUASI-STATIONARY (BB)



Excessive rain can also be produced by convection with warm cloud tops (shallow convection), therefore not offering signatures which indicate intense convection (with potential significant rain) like:

Cold-U ring
Jumping anvil cirrus

E.g. storms in the Mediterranean sometimes show warm cloud tops but are prolific lightning producers with heavy rainfall amounts.

Look for ingredients!

*Thank you very much for
your attention !*



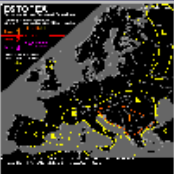
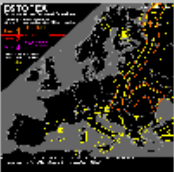
The background of the slide is a photograph of a dramatic sky. It features large, dark, billowing clouds that are illuminated from below by a bright light source, likely the sun, creating a strong orange and yellow glow. The sun itself is visible as a bright, hazy orb in the upper left corner.

Overview about the European Storm Forecast Experiment, ESTOFEX

European Storm Forecast Experiment

forecasts
verification
news
ESSL's severe
weather reports
team
FAQ
research &
education
links
sponsors
disclaimer
donate

Forecasts

forecast	map	threat level	period	forecaster
Storm forecasts				
Storm Forecast issued: Thu 02 Jun 2011 12:49		1	Fri 03 Jun 2011 06:00 - Sat 04 Jun 2011 06:00 UTC	TUSCHY
Storm Forecast issued: Wed 01 Jun 2011 13:33		1	Thu 02 Jun 2011 06:00 - Fri 03 Jun 2011 06:00 UTC	GATZEN

The forecast archive can be found [here](#).

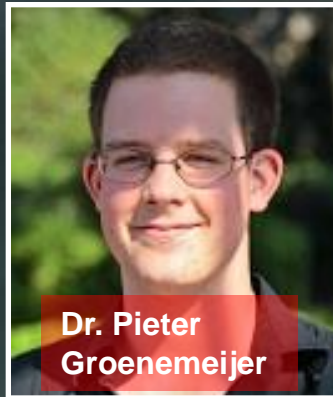
Please note that forecasts produced by ESTOFEX are not warnings of imminent severe weather. Your national meteorological service carries the responsibility to warn you for severe weather. These warnings can be found on meteoalarm.eu.

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Forecasters of ESTOFEX

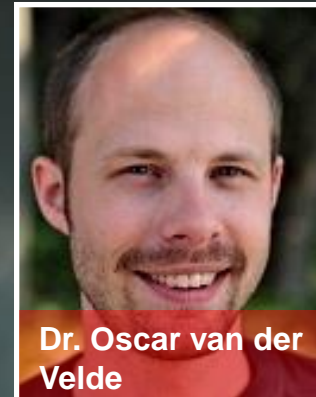
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Dr. Pieter
Groenemeijer



Dr. Johannes
Dahl



Dr. Oscar van der
Velde

Christoph
Gatzen



Helge Tuschy



Oliver
Schlenczek



Tomas Pucik



Georg Pistotnik



European Storm Forecast Experiment (ESTOFEX)

What is ESTOFEX?

An initiative of a team of European **meteorologists** and **students in meteorology** and serves as a platform for exchange of knowledge about severe convective storms in Europe and elsewhere.



ESTOFEX offers a GUIDANCE where organized convection is forecast. This guidance can be used by local forecast offices for daily warning preparation.

Which risk of (extremely) severe convective weather phenomena does ESTOFEX forecast ?

Severe convective weather phenomena:

- **tornado** (waterspouts)
- **hail** with a diameter of at least 2.0 cm
- **wind gusts** with a speed of at least 25 m/s (92 km/h or about 49 knots)
- **excessive rainfall** of at least 60 mm

Extremely severe convective weather phenomena:

- tornado; (E)F 2 or stronger
- hail with a diameter of at least 5.0 cm
- wind gusts with a speed of at least 33 m/s (119 km/h or about 65 knots)

The meaning of the level areas

- 1) E.g. probability of 15 % means a 15 % chance for that event to happen within a radius of 41 km around each point. May seem low, but compared to climatology, it is significant
- 2) **Level 0** expected probability of severe convection appears insignificant
- 3) **Level 1** most common threat level; low threat of severe weather
- 4) **Level 2** large confidence of severe storm occurrence and slight risk for extreme severe
- 5) **Level 3** major severe thunderstorm outbreak (e.g. 15th August 2008)

Issued: Wed 13 Aug 2008 21:18 Z

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