BASIC SATELLITE METEOROLOGY COURSE



Introduction to Remote Sensing

Marianne König, EUMETSAT marianne.koenig@eumetsat.int



Definition

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Example: Aerial, satellite, spacecraft observations of the earth (and other planets / stars / galaxies!)





Opposite

In-situ measurements – the measurement device is in physical contact with the probe



Example: Thermometer, anemometer, pitot tube, ...





Biological Evolution

You are very familiar with the concept:

- Your eyes detect electromagnetic radiation in the form of visible light
- Your ears detect acoustic information
- Some animals detect the earth's magnetic field (bird migration)







Remote Sensing in Meteorology

Methods that detect and measure electromagnetic energy

Focus in this course:

- Underlying physical concepts
- Interpretation of remotely sensed images
- Important image characteristics from different sensors



Electromagnetic Waves



Characteristics:

- Wavelength λ (m or cm or μ m)
- Propagation velocity c (m/s)
- Frequency $v = c / \lambda (s^{-1})$
- Wavenumber =1/ λ (cm⁻¹)



Electromagnetic Spectrum





Electromagnetic Spectrum



Sources of Radiation (Meteorological Applications)

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Fundamental Radiation Law: Planck's Law

$$B(\nu,T) = \frac{2h\nu^{3}}{c^{2}} \frac{1}{e^{h\nu/kT} - 1}$$

Units: Watts per unit area, per unit angle, per unit wavelength (or wavenumber), e.g. W/m²/ster/cm⁻¹

= Boltzmann's constant k Т = Temperature

h

= Planck's constant

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Fundamental Radiation Law: Planck's Law

Concept of a Blackbody – Emissivity and Brightness Temperature

A "Blackbody" is an object of temperature T which radiates energy according to Planck's Law.

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A "Blackbody" is an object of temperature T which radiates energy according to Planck's Law. Nature does not have perfect blackbodies – physical property of matter is its emissivity (0-1):

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Remote Sensing of the Atmosphere

- What do we measure?
- Solar radiation: reflected by the surface, scattered by molecules, cloud droplets, ice crystals, aerosols, absorbed by the atmosphere
- Thermal radiation: emitted by the earth / clouds / atmosphere,, absorbed by the atmosphere, clouds, aerosols
- What about thermal radiation from the sun???

Explanation

At our Earth's distance from the sun, the radiation received from the sun is approximately on the same energy level as the radiation emitted from the earth/atmosphere

Interaction Processes

Absorption – Emission - Scattering

- Absorption: Energy of the electromagnetic wave is taken up by matter (e.g. change in atomic state)
- Emission: Energy change in the matter (e.g. change in the atomic state) releases electromagnetic radiation
- Emission = Absorption!!
- Absorption coefficient = property of matter
- Scattering: Radiation is "geometrically' forced to deviate from a straight line (reflection = specific scattering)

Processes for Solar Radiation

Processes for Solar Radiation

Why is grass green?

Vegetation Mapping

Vegetation cover is often mapped using near-infrared radiation (~0.8 µm wavelength) because of the high reflectivity here

Processes of Thermal Radiation

Processes of Thermal Radiation

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Earth Spectrum and Planck Curves

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Illustration: Beam at 11 µm wavelength ("Window")

Earth Surface

Illustration: Beam at 6.5 µm wavelength (WV Absorption)

Earth Surface

Weighting Functions

To Remember

Each radiance measured in the VIS or IR or MW part of the spectrum is the result of a number of processes, e.g.

- Source position
- Illumination geometry
- Surface materials
- Passage of energy through the atmosphere

Scattering

- Scattering by particles which are much smaller than the electromagnetic wavelength ("Rayleigh Scattering")
- Scattering by particles which are of same size and larger than the electromagnetic wavelength ("Mie Scattering")

Distribution for all angles: phase function

Rayleigh Scattering

- **Rayleigh scattering**, named after the British physicist Lord Rayleigh, is the elastic scattering of light or other electromagnetic radiation by particles much smaller than the wavelength of the light. The particles may be individual atoms or molecules.
- Scattering is ~ λ^{-4} , i.e. scattering occurs for shorter wavelengths!

Wavelength 0.87 µm Cloud droplets 1- 10 µm

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Practical Example: MODIS Imagery, 03 April 2011 Solar Bands

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Practical Example: MODIS Imagery, 03 April 2011 Thermal Bands

To Remember

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- Source position
- Illumination geometry
- Surface materials
- Passage of energy through the atmosphere

Radiances at a given wavelength region carry the information of these processes (but not of processes which are not important in this specific wavelength!) – measurement has to be appropriate for what you want to measure!!!!

Example: VIS data carry no temperature information

The equation of radiative transfer simply says that as a beam of radiation travels, it loses energy to absorption, gains energy by emission, and redistributes energy by scattering.

The equation is a differential equation, numerical models exist which provide a solution (Radiative Transfer Models, RTMs).

Image Acquisition

The remote sensing system must first detect and measure the energy signal:

- Light-sensitive film (classical photography)
- Electronic sensors (electrical signal is proportional to the amount of energy received) (digital photography, imaging satellite instruments)

What Is Really Measured?

- The actual measurement on the sensor level is not a radiance (or reflectance or brightness temperature), but some measure of the electric signal (voltage, current) – which is digitized for transmission to the ground station. These digital values are often called "counts"
- Some type of calibration can convert this count to a radiance (and then e.g. to a brightness temperature via Planck's Law)
- Calibration: on-board (blackbody, solar diffuser), vicarious methods (objects of known temperature/reflectivity)

Spatial Resolution

Counts and Brightness Temperatures

387	<mark>352</mark>	339
270.76	264.90	262.62
<mark>340</mark>	<mark>333</mark>	333
262.79	261.54	261.44
<mark>276</mark>	<mark>297</mark>	305
250.53	254.77	256.32

A Pixel is not a Square!

Display systems usually display one measured pixel as one pixel on a monitor – when zoomed this one pixel becomes a square.

The actually measured energy does not equally come from all areas within this square but mostly from the centre (point spread function)

The digital recording of the energy received by a sensor leads to discrete levels ("counts") that are recorded as integer values (e.g. 8 bits or 10 bits or 16 bits) – i.e. values between two counts cannot be radiometrically resolved:

Example (for IR):

Count 333 – 261.54 K Count 332 - 261.36 K

Temperatures "in between" cannot be measured

Spectral Resolution

The spectral resolution describes the ability of the system to distinguish between different parts of the range of measured wavelengths – important here are the number of wavelength intervals ("bands" or "channels") and how wide or narrow a band is.

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"Sounding" type of instruments:

Many narrow bands, each also described by a filter function – but not really used to produce "images" but to obtain information on the vertical structure of the atmosphere.

Most advanced sounders: Hyperspectral instruments with such narrow "bands" that concept of a response function does not apply

Black: hyperspectral instrument (IASI)

IASI Example

Thank you very much!

- Thank you for your attention
- I hope I could give you a basic introduction into what remote sensing is about
- Enjoy the quiz!

